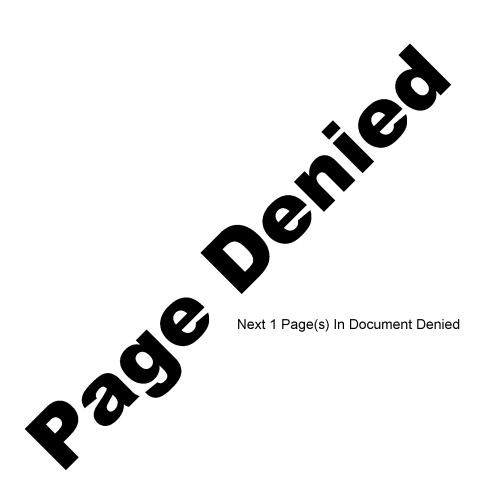
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#### **STAT**



## AIR FORCE SURVEYS IN GEOPHYSICS

**STAT** 

No. 86

THE ARDS MODEL ATMOSPHERE, 1956

R.A. MINZNER W.S. RIPLEY

DECEMBER 1956

GEOPHYS 'S RESEARCH DIRECTORATE

AIR FORCE CAMBRIDGE RESEARCH CENTER

AIR RESEARCH AND DEVELOPMENT COMMAND

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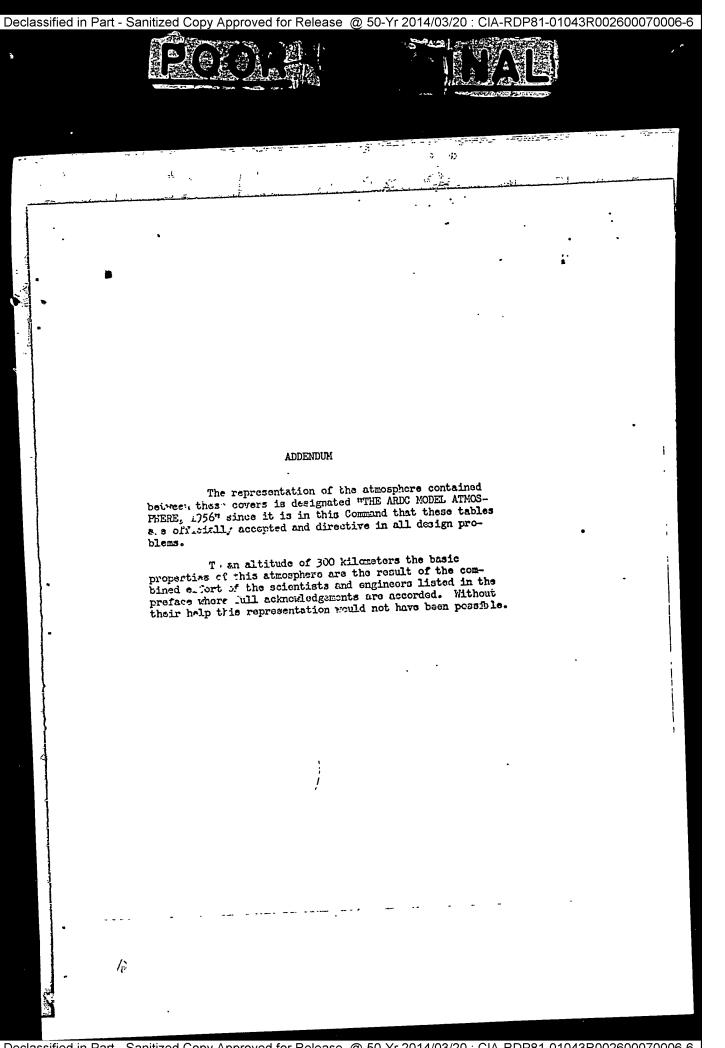
R.A. MINZNER W.S. RIPLEY

DECEMBER 1956

GEOPHYSICS RESEARCH DIRECTORATE

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BEDFORD, MASSACHUSETTS





#### PREFACE

The 1956 ARDC MODEL ATMOSPHERE, defined and tabulated to 542,248 meters or 1,850,870 feet in this Air Force Survey in Geophysics, has been prepared in partial fulfillment of ARDC Technical Requirement 140-56. This MODEL is to be used as the basis for engineering and design work performed within ARDC and by its contractors, insofar as the work requires the use of a model representing the average condition of atmospheric properties within the altitude limits of this MODEL.

This MODEL ATMOSPHERE is designed to be used for the same purposes as a standard atmosphere. For some of these purposes the MODEL should serve in the following ways:

- As a reference atmosphere to be used in calculating flight performance of aircraft.
- As the basis for calibrating barometric altimeters, where observed departures of atmospheric properties from the values of the MODEL provide the means for computing altimeter correction.
- 3. As the basis for ballistic tables where the observed departures of the atmospheric properties from the values of the MODEL provide the basis of corrections to be put into gunnery and bombing computers.
- 4. As a time average of the actual physical conditions existing at various altitudes for aircraft engineering and design purposes, and for use in solving geophysical problems.

It should be emphasized, particularly in regard to item 4, that this MODEL most probably will never completely match the actual atmosphere, and may only rarely approximate the average value at all altitudes simultaneously. While the properties at some altitude may exactly fit the values of the MODEL at any istant, the properties at other altitudes simultaneously may depart drastically from tabulated values. The greatest percentage departures probably occur at the higher altit 's. Maximum and minimum pressures at 120 km, for example, may differ by as much as a factor of 3. Neither this MODEL nor any other calculated model will accurately depict the total atmosphere at any particular moment.

The tables and graphs of this MODEL approximate the best average of available temperature, pressure, and density data, compiled and processed under Project 7603, "Atmospheric Standards." The tables are also consistent with the recently adopted Extension to the United States (ICAO) Standard Atmosphere 50,51 (1956) which were prepared concurrently under the same project. Both are consistent with the basic properties of the International Civil Aviation Organization (ICAO) Standard Atmosphere 26-28 adopted by the United States on November 20, 1952.

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The tables of this MODEL partially duplicate the tables of the ICAO Standard Atmosphere, (in the altitude region of -5,000 to +20,000 geopotential meters) although the tables of this MODEL are given in larger increments. This partial duplication is desirable and necessary, not only for the sake of continuity, but because this MODEL includes values of seven additional altitude-dependent properties not found in the ICAO Standard: Acceleration of gravity, scale height, molecular weight, particle speed, number density, mean free path, and collision

The ARDC MODEL differs from the standard atmosphere not only because of the greater altitude of the former but because the MODEL is intended to be reviewed annually and modified at any time, if necessary, to reflect significant changes in thinking brought about by more reliable atmospheric data.

We wish to acknowledge the assistance of the several members of the Geophysics Research Directorate who participated in various ways in the preparation of this survey: Dr. R. Penndorf and Mr. M. Dubin for helpful suggestions and counsel, and Mr. L. R. Shedd for his expeditious handling of many details.

We also wish to thank the members of the Working Group on Extension to the Standard Atmosphere for their helpful suggestions and encouragement. This Work-

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<sup>\*\*\*\*\*</sup> Replacement for Dr. Havens upon his departure from Naval Research Lab. \*\*\*\*\* Replacement for Dr. Brasefield upon his departure from Signal Corps Eng. Lab.

We are especially indebted to two subcommittees of this Working Group:

The first subcommittee, consisting of Dr. H. Newell. Dr. H. Kallman, and Mr. R. A. Minzner, formulated the general aspects of the temperature-altitude profile between 130 and 300 kilometers, and made recommendations concerning the degree of dissociation of  $\Omega_2$  and  $\Omega_2$  in this region.

The second subcommittee, consisting of Mr. L. P. Harrisor, Mr. W. J. O'Sullivan, Mr. W. Scholl, and Mr. R. A. Minzner, studied some of the aspects of the following atmospheric properties: coefficient of viscosity, kinematic viscosity, and the speed of sound. This subcommittee recommended departures from the ICAO values of these properties and thereupon suggested values of constants, empirical expressions, and maximum altitude of tabulation for these properties.

We are particularly grateful to Dr. F. L. Whipple whose efficient chairmanship expedited the accomplishment of the Working Group, and to Mr. N. Sissenwine who in the capacity of Executive Secretary handled a flood of detail.

Finally we wish to thank Dr. H. Werler of the U. S. Weather Bureau. Dr. Wexler served with Mr. Sissenwine as Co-chairman of the Parent Committee on Extension to the Standard Atmosphere, and though not an official member of the WGESA, was over in the background to lend his advice and support wherever needed.

R. A. MINZNER

W. S. RIPLEY

Geophysics Research Directorate

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## ABBREVIATIONS AND SYMBOLS

8.	acceleration .
ъ	subscript indicating base or reference level
_	degrees, in thermodynamic Celsius scale
°C	
Cg	speed of sound
$(c_s)_o$	sea-level value of C <sub>s</sub>
c <sub>p</sub>	specific heat of dry air at constant pressure
$\mathtt{c}_{\mathbf{v}}$	specific heat of dry air at constant volume
cgs	centimeter-gram-second system of units
cm	centimeter
đ	differential symbol
e	base of natural logarithms
•F	degrees, in thermodynamic Fahrenheit scale
F	force
f(H)	undefined function of H representing $\mathbf{T}_{\mathbf{M}}$
fps	foot-pound-second system of units
ft	foot
ft'	standard geopotential foot
G	dimensional constant in the geometric-geopotential relationship
g	effective value of acceleration of gravity
g <sub>o</sub>	sea-level value of g
gø	sea-level value of g at latitude $\phi$
gm	grem
gm-mol	gram mole

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## ADBPEVIATIONS AND SYMBOLS (Contd.)

H altitude in geopotential measure  $H_{o}$ sea-level value of H, (zero) НЪ altitude at base of layer, or reference level in geopotential measure Hg mercury  $\mathbf{H}_{\mathbf{s}}$ scale height H<sub>s</sub> ' geopotential scale height  $(H_{\rm g})_{\rm o}$ sea-level value of Hg in inch international nautical mile i n mi •K degrees, in thermodynamic Kelvin scale kg kilogram kgf kilogram force kg-mol kilogram mole km geometric kilometer km' standard geopotential kilometer L mean free path sea-level value of L molecular-scale-temperature gradient  $\partial T_{M}/\partial H$ L<sub>M</sub> L length 1b pound lbf pound force ln

natural logarithm

## ABBREVIATIONS AND SYMBOLS (Contd.) log logarithm М apparent molecular weight of air Mo sea-level value of M mass numerically equal to the molecular weight (a mole) (geometric) meter standard geopotential meter millibar шþ mks meter-kilogram-second system of units mass (1) Avogadro's number (standard) N atmospheric number density n newton $n_o$ sea-level value of n number density of a gas at temperature $T_i$ and pressure $P_o$ (Loschmid's number) P atmospheric pressure sea-level value of P ro pdl poundal value of P at base of layer or reference level constant, $\frac{GM_0}{R^*}$ •R degrees, in thermodynamic Rankine scale R\* universal gas constant effective radius of earth (at 45° 32' 40" N. lat.)

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## ABBREVIATIONS AND SYMBOLS (Contd.)

<sup>r</sup> ø	radius of earth at latitude Ø
S	Sutherland's constant
sec	second
T	temperature (real kinetic) in the absolute thermodynamic scales
To	sca-level value of T
T <sub>i</sub>	temperature of the ice point in the absolute thermodynamic scales
T <sub>M</sub>	molecular-scale temperature in the absolute thermodynamic scales
$(T_M)_o$	sea-level value of $T_{M}$
$(\tau_{M})_{b}$	value of T <sub>M</sub> at base of layer or reference
t	temperature in nonabsolute thermodynamic scales, also signifies time
t <sub>o</sub>	sea-level value of t
<sup>t</sup> i	temperature of the melting point of ice at 1013.250 mb air pressure in the nonabsolute thermodynamic scales
$^{t}_{M}$	molecular-scale temperature in the nonabsolute scales
$\overline{\mathbf{v}}$	particle speed (arithmetic average)
$\overline{v}_{o}$	sea-level value of $\overline{\mathbf{V}}$
Y	volume of one mole of air at existing conditions of T and P
v <sub>o</sub>	sea-level value of v
v <sub>i</sub>	volume of one mole of air at a temperature $\mathbf{T}_{\mathbf{i}}$ and pressure $\mathbf{P}_{\mathbf{o}}$
Z	altitude in geometric measure
α	real temperature gradient $\partial T/\partial Z$

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## ABBREVIATIONS AND SYMBOLS (Contd.)

αι	real temperature gradient $\partial$ T/ $\partial$ H
β	constant used in the empirical expression for the coefficient of viscosity
7	ratio of specific heats, $c_{ m p}/c_{ m v}$
8	partial differential symbol
η	kinematic viscosity
$\eta_{o}$	sea-level value of $\eta$
μ	coefficient of viscosity
$\mu_{\mathbf{o}}$	sea-level value of $\mu$
ν	collision frequency
٧ <sub>o</sub>	sea-level value of v
π	ratio of circumference to the diameter of a circle
ρ	atmospheric density .
ρο	sea-level value of $\rho$
ρ <sub>i</sub>	ice-point value of $\rho$
б	effective collision diameter of a mean air molecule (standard)
ø	latitude of the earth
ω	specific weight
ω <sub>o</sub>	sea-level value of $\omega$

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#### ABSTRACT

A realistic model of atmospheric properties based on reliable observations and current theories is presented.

Fifteen atmospheric properties are discussed and tabulated, thirteen to 500 km and two to only 90 km. The values of these properties are internally consistent through classical equations, and are dependent upon (1), a defined, linear, segmented, molecular-scale temperature function, (2) a molecular weight function, and (3) an acceleration of gravity function. Values of twelve physical constants required in the computations are adopted as exact. Internationally agreed-upon, exact transformation factors are employed in converting from Metric to English units. Both Metric and English tables are presented, and computational procedure is discussed. A thorough discussion of geopotential altitude, effective radius of the earth, and molecular-scale temperature is given. The relative virtues and validity of two methods for computing the acceleration of gravity are discussed. The concept and validity of the various properties as applied to high altitudes are considered briefly.

#### THE ARDC MODEL ATMOSPHERE

1956

(Tables and Graphs for Altitudes to 542,686 Meters or 1,850,870 Feet)

#### 1. Introduction

#### 1.1 Background and Early History of Standard Atmospheres

Standard atmospheres have been used for nearly a hundred years for altimetry purposes. The earliest of these were very simple and were based on an isothermal atmosphere. With the development of aircraft and precision artillery during the First World War, 1914-1918, the need for more extensive atmospheric tables for aeronautical and hallistic purposes became apparent. Atmospheric temperatures were measured at various locations in southern and western Europe. Several functions approximately fitting these temperature data were proposed and used in various countries for deriving an analytical expression for atmospheric pressure and density. No generally agreeable function was proposed, however, until 1919 when Toussaint49 suggested a segmented straight-line function as the basis for an international standard. Toussaint's temperature function was defined by a value of 15 degrees Celsius (°C) at sea level, a constant gradient of -.0065CC per meter from sea level to 11,000 meters, (m), (yielding -56.5°C for 11,000 m), and a constant gradient of zero degrees per meter from 11,000 m to 20,000 m altitude.

### 1.2 First U. S. Aeronautical Standard Atmosphere

The Toussaint formula with minor variations has remained the basis for all major aeronautical standards prepared for the 0 - 20 km altitude region. These include the first United States Standard Atmosphere prepared by Gregg21 in 1922, and the modification, extension, and amplification of the Gregg standard prepared by Diehll4 in 1925. Neither of these agreed exactly with the Toussaint proposal, however: Gregg terminated his analytically derived atmosphere at 10 km altitude although he presented observed data to 20 km; Diehl extended the analytical atmosphere to 20 km but established the tropopause at an altitude of 10,769.23 m (65,000 ft) with a temperature of -55°C, instead of at 11,000 m and -56.5°C, as suggested by Toussaint. Thus Diehl's stratosphere, 10,769.23 m to 20,000 m, was warmer by 1.5°C than that used by Toussaint.

Brombacher 4,5 amplified the Gregg Standard Atmosphere in 1926 and again in 1935 by adding tables of altitude as a function of pressure for altimetry purposes.

## 1.3 First International Standard

In 1924 the International Committee on Air Navigation (ICAN)<sup>29</sup> prepared an international standard atmosphere based exactly on Toussaint's temperature-altitude function. This standard was adopted throughout most of Europe. It was never adopted formally by the United States, however, because of two small but basic differences between this and the Diehl-U. S. Standard.

In addition to using different altitudes and temperatures for the tropopause, the ICAN and U. S. Standard also used different values for the acceleration of gravity at sea level, 9.8 and 9.80665 respectively. These differences prevented United States and European agreement on a standard atmosphere until 1952 when a new international organization, ICAO, reached a compromise.

# 1.4 ICAO Standard Atmosphere<sup>28</sup> -- New U. S. Standard<sup>26</sup>,27

Between June 1950 and November 1952 the International Civil Aviation Organization (ICAO), of which the United States was a member, proposed and adopted a compromise standard atmosphere in which the United States standard sealevel value of gravity, and the ICAN values of tropopause altitude and troppause temperature were employed. This ICAO Standard Atmosphere was formally adopted as the United States Standard Atmosphere by NACA vote on 20 November 1952.

## 1.5 High Altitude Models -- Warfield, Grimminger

The activities of ICAO emphasized international agreement and refinement of atmospheric tables within the altitude range of existing standards; i.e., sea level to 20,000 meters altitude. The ICAO did not concern itself with high altitude tables. The advances in aeronautics and ballistics during and since World War II resulted in demands for atmospheric tables to much greater altitudes. In 1947 these demands were met in part by Warfield's "Tentative Tables for the Properties of the Upper Atmosphere"52 which depicted the atmosphere to 120,000 meters altitude and which were designed to be a continuous extension of the tables of the Diehl-U. S. Standard at 20,000 m altitude. The Warfield tables were based on the best 1946 estimates of atmospheric temperature, and considered the variations of molecular weight of air and the acceleration of gravity with increasing altitude.

The 120 km altitude upper limit of the Warfield tables was inadequate, however, even before the publication of the report, and Grimminger in 1948 published tables of atmospheric properties to altitudes of over 8,800 km. These tables were essentially in agreement with the Warfield tables up to 120 km and were based on the best 1947 theoretical and experimental data.

## 1.6 New Data from Rocket-Borne Experiments

Simultaneously with the preparation of the Warfield and Grimminger tables, a new research tool, the upper air sounding rocket, was beginning to be exploited. This new device permitted making measurements of the atmosphere by

direct probing methods not previously possible. The new data compiled in 1952 as the Rocket Panel Atmosphere45 indicated that pressures in the Warfield and Grimminger tables were 2 times higher than observed at 70 km, 5 times higher than observed at 90 km, and over 10 times higher than observed at 120 km. These discrepancies, plus the fact that the Warfield tables were not continuous with the newly adopted ICAO Standard, initiated the preparation of this extension of the ICAO Standard to high altitudes.

#### 1.7 Extension to the Standard Atmosphere

In November 1953 the Geophysics Research Directorate, Air Force Cambridge Research Center, of ARDC, USAF, together with the U. S. Weather Bureau sponsored a three-day "Open Mccting on Extensions to the Standard Atmosphere." Standard atmosphere requirements and scientific data supporting various models were presented. Brombacher presented a Standard Atmosphere proposal which was not accepted because of an unrealistic stratosphere and because the constant gravity assumption employed was inconsistent with the ICAO Standard and this assumption introduced errors in the analysis. A Working Group on Extension to the Standard Atmosphere (MCESA) was appointed to recommend the temperature-altitude profile and other constants necessary for the preparation of the desired extension.

The discussions of the first meeting 18 of the Working Group dealt principally with the temperature-altitude profile in the 20 to 53 kilometer region. Temperatures were also recommended for the region between 53 and 83 km, although these were replaced by slightly different values at a later meeting. Recommendations were also made at this first meeting regarding the atmospheric properties to be included in the standard. Differences of opinion existed on the manner of accounting for variable gravity, and some conflicting recommendations resulted from this meeting.

The task of preparing the text and tables for the extension to the Standard Atmosphere was assigned to GRD (Geophysics Research Directorate). The recommendations were studied, and Minzner 40 prepared a paper, "Three Proposals for U. S. High Altitude Standard Atmosphere," which was presented at the second meeting 19 of the Working Group. Each of the three proposals suggested a different method for handling the acceleration of gravity and molecular weight as variables in the hydrostatic equation. Only one of these three proposals was consistent with the ICAO Standard Atmosphere and that one, using geopotential to account for variable gravity, and molecular-scale temperature to account for variable molecular weight, was adopted by the Working Group.

Preliminary tables of atmospheric properties to 130 km, 41 prepared at CRD, were tentatively adopted at this meeting. These tables were consistent with the temperature-altitude function to 83 km recommended by the Working Group and consistent with the temperatures of the Rocket Panel Atmosphere above this altitude. A subcommittee was appointed, however, to make recommendations concerning molecular weight and temperatures for extending the Standard Atmosphere to 300 km altitude.

This subcommittee met with several consultants and then agreed upon certain boundary conditions for oxygen and nitrogen dissociation, as well as for atmospheric temperature. Using these boundary conditions and all the available atmospheric pressure, temperature, and density data above balloon altitudes, two separate proposals were prepared, one at Rand Corporation 34,35 and the other at GRD.42

The Rand proposal assumed a density-altitude function and a molecular weight gradient arbitrarily related to this density function. From these, there was derived a nonlinear temperature-altitude profile with no discontinuous first or second derivatives.

The GRD proposal, in keeping with previous Working Group recommendations, assumed several constant gradients of molecular-scale temperature for as many altitude regions. These gradients were chosen to yield values of pressure and density consistent with the average of observed values of these properties below 160 km altitude, and consistent with current estimates of these properties at higher altitudes. Molecular weights 39 were computed from diffusion theory and the agreed-upon boundary conditions. The GRD proposal was adopted at the third and final meeting 0 of the Working Group.

A summary of the adjusted recommendations 17 resulting from the three meetings of the WGESA was prepared. A supplemental set of recommendations 43 on previously unresolved questions was also prepared. Within the framework of these recommendations, this ARDC MODEL ATMOSPHERE and the Extension to the U. S. Standard Atmosphere have been prepared.

## 2. Systems of Altitude Measure and Related Parameters

In accordance with agreements concerning publication of international aero-logical tables 30 and in keeping with the existing United States (ICAO) Standard Atmosphere, the basic altitude parameter of this MODEL is taken to be geopotential H, expressed in standard geopotential meters, m¹. Supplemental to the existing (ICAO) United States Standard, this MODEL has been prepared with parallel tabulations in integral values of both geopotential and geometric altitude measure so that the values of tabulated properties are given for both integral geopotential and integral geometric kilometers.

The relationship between geopotential and geometric altitude depends directly upon the value of the acceleration of gravity at sea level at a particular altitude and upon the variation of the acceleration of gravity with altitude and latitude. The definition of the special unit of geopotential used in this MODEL is also related to the specific sea-level value of gravity, adopted by ICAO and used in this MODEL. Therefore, a digression is made to present a detailed discussion of the acceleration of gravity before geopotential is discussed further.

#### 2.1 Acceleration of Gravity

#### 2.1.1 Sea-level value

The sea-level value of the acceleration of gravity used in this MODEL is defined to be 9.80665 m sec-2 exact. This value was originally announced by Defforges and Lubanskil3 at the 1891 meetings of the International Committee on Weights and Measures as the best value for 15° latitude. Since then, it has been used by physicists and others as an arbitrary standard and was recently adopted as an international standard in the ICAO Standard Atmosphere. It has long been recognized, however, that this value of g is not correct for 15° latitude but rather is the value for 15° 32' 10° latitude. This corrected latitude is the one to which all tables in this MODEL apply.

#### 2.1.2 Altitude variation - classical expression

The variation of the acceleration of gravity with geometric altitude is classically expressed by the equation

$$g = g \left[ \frac{r \beta}{r \beta + Z} \right]^2 , \qquad (1)$$

where

g = the acceleration of gravity of a point (in m sec-2),

Z = the geometric altitude of the point (in m),

 $g_{\not p}$  \* the sea-level value of g at the latitude  $\not p$  of the point (in m sec<sup>-2</sup>), and

 $\mathbf{r}_{\mathbf{0}}$  = the radius of the earth at latitude  $\mathbf{0}$  .

In its fundamental form this equation applies rigorously only for a nonrotating sphere composed of spherical shells of equal density. The earth, however, is definitely not spherical; furthermore, its rotation introduces centrifugal acceleration which varies with latitude and which increases with altitude. The sea-level value of the centrifugal acceleration at any selected latitude may be accounted for, in equation (1), by the proper choice of an effective value of gg. The increase of centrifugal acceleration with increasing altitude is not accounted for in the simple unadjusted inverse square law, which describes only the decreasing Newtonian component of the effective value of g. Hence, values of g computed from equation (1) become increasingly inaccurate as altitude increases. An adjustment of the value of rg to an effective

<sup>#</sup> Basic constant

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radius, however, was found to greatly improve the validity of that equation even at altitudes as great as 500 km.

#### 2.1.3 Effective earth's radius

Harrison  $^{23}$ , using a suggestion by Lambert  $^{37}$ , developed an expression for an effective earth's radius as a function of latitude. This effective radius is derived in a manner consistent with the effective sea-level value of g at latitude  $\beta$ , and consistent with the vertical gradient of g at the given latitude (neglecting local anomalies), assuming the International Ellipsoid represents the figure of the earth. The value of effective earth's radius at  $45^{\circ}$  32'  $40^{\circ}$ , computed from Harrison's equation (given in Appendix M) is

$$r = 6,356,766$$
 meters

which, for purposes of this MODEL, will be considered as an exact constant.

#### . 2.1.4 Computational equation

The exact form of the equation used to compute the acceleration of gravity and to relate geopotential to geometric altitude in this MODEL is

$$g = g_0 \left[ \frac{r}{r+Z} \right]^2 , \qquad (1a)$$

where

OV W

- g \* the acceleration of gravity in meters per second squared, (m sec-2) at altitude Z and at latitude 45° 32' 40", hereafter,
- go = 9.80665 m sec-2(exact), the sea-level value of g at 45° 32° 40° latitude, and
- r = 6,356,766 m (exact), the effective earth's radius at latitude 45° 32' 40".

(For purposes of this MODEL, this equation is assumed to apply in free air below sea level as well as above sea level.)

#### 2.1.5 Best available analytical expression

A more exact equation for g as a function of Z and  $\emptyset$  in free air, based directly on the International Ellipsoid and the International Gravity Formula, was developed by Lambert 36,38 in the form of an infinite,

<sup>≠</sup> Basic constant

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alternating power series (see Appendix.N). The values of g computed from equation (la) are in good agreement with those computed from Lambert's more exact equation. For an altitude of 500 km the value of g from the two methods differs only by 3 parts in the fifth significant figure, or less than 1/1000 of 1 per cent. For lower altitudes the agreement is much better. Values of geopotential computed for specific values of Z on the basis of equation (la) are also in good agreement with corresponding values of geopotential computed on the basis of the more exact equation for g. The percentage departures are similar. The more exact expression for g was not employed in this MODEL because of its much greater complexity. In the U. S. Standard Atmosphere, the tables will be recomputed by machine and will be based on the more exact equation.

#### 2.2 Relation of Geopotential to Geometric Altitude

### 2.2.1 Basic definition of geopotential

The geopotential of a point is defined as the increase in potential energy per unit mass lifted from mean sea level to that point against the force of gravity.

#### 2.2.2 Analytical development

The increase in potential energy of a body lifted against the force of gravity, from sea level, through a vertical distance to a given point is:

$$\Delta E = \int mgdZ, \qquad (2)$$

where

 $\Delta E$  = increase of potential energy over the sea-level value, in joules,

m = mass of the body in kilograms, kg.

The geopotential of that point  $\Delta E/m$  is therefore:

$$\frac{\Delta E}{m} = \int g dZ. \tag{2a}$$

If geopotential is given a special designation, H, with special units, we have:

$$GH = \frac{\Delta E}{m} = \int g dZ, \qquad (2b)$$

$$GdH = gdZ,$$
 (2c)

or

$$H = \frac{1}{G} \int g dZ, \qquad (2d)$$

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where

H = geopotential (in unspecified units), and

G = a proportionality factor depending upon the units of H.

When H is in units of joules kg-1 or equivalently in m<sup>2</sup> sec-2, G is nondimensional and unity. If H is expressed in some other units, standard geopotential meters for example, the value and dimensions of G must be correspondingly changed.

2.2.3 The standard geopotential meter 26-28

The basic unit of geopotential employed in this MODEL is the standard geopotential meter where one standard geopotential meter, m', is defined to be an increment of potential energy per unit mass equal exactly to

It is evident from equation, (2b) that if H is expressed in m; G is equal to 9.80665 m² sec-2 m;-1. // One standard geopotential meter is therefore the vertical distance through which one kilogram mass must be lifted against the force of gravity to increase its potential energy by 9.80665 joules. If a region existed where the value of the acceleration of gravity were constant at 9.80665 m sec-2 over an altitude interval of one geometric meter, in this region one geometric meter and one geopotential meter would then be exactly equal. This condition is very closely approximated at sea level at 15° 32° 10° latitude. Since g normally does decrease with altitude, however, even over a one meter interval, an altitude of one geometric meter at this latitude has a geopotential altitude of slightly less than 1 m; (see table in Section 2.2.5). Above sea level, at all points where the altitude gradient of g is continuously negative from sea level, the altitude in standard geopotential meters is always numerically less than the altitude in geometric meters, and the numerical difference increases with increasing altitude.

2.2.4 Standard geopotential kilometer and standard geopotential centimeter

The basic concept of the metric system of units leads directly to the conclusion that one geopotential kilometer, km', is equal to one thousand geopotential meters; i.e.,

$$1 \text{ km}^{\dagger} = 1 \times 10^3 \text{ m}^{\dagger}$$
. (3a)

<sup>\*</sup> Basic conversion of units

<sup>#</sup> Derived constant, inferred from transformation of units

Also, it follow that one geopotential centimeter, cm; is equal to one one-hundredth of a geopotential meter; i.e.,

$$1 \text{ cm}! = 1 \times 10^{-2} \text{ m}!.$$
 (3b)

One cm' may also be defined in cgs units directly by analogy with equation (3),

$$1 \text{ cm}^2 = 980.665 \text{ ergs gm}^{-1} = 980.665 \text{ cm}^2 \text{ sec}^{-2} = .01 \text{ m}^2,$$
 (3c)

where

980.665 is the numerical value of g in the cgs units.

2.2.5 Conversion of standard geopotential meters to geometric meters

The replacement of g in equation (2b) by equation (la) results

in

$$H = \frac{g_0}{G} \int \left[ \frac{\mathbf{r}}{\mathbf{r} + \mathbf{Z}} \right]^2 d\mathbf{Z}, \tag{4}$$

where

H = geopotential in standard geopotential meters, m:,

Z = geometric altitude in m,

 $Q = 9.80665 \text{ m}^2 \text{ sec}^{-2} \text{ m}^{1-1} (\text{exact})^{1/2}$ 

 $g_0 = 9.80665 \text{ m sec}^{-2} (\text{exact})^{4}$ ,

r = 6,356,766 m (exact).

Performing the indicated integration leads to

$$H = \left[\frac{g_0}{G}\right] \frac{rZ}{r + Z'} \tag{5}$$

or

$$Z = \frac{rH}{\left[\frac{g_0}{G}\right] r - H} \tag{6}$$

# Basic constant

The ratio  $g_0/G$  appearing in equations (4), (5), and (6) is numerically unity while its dimensions are m'/m. Hence while the ratio  $g_0/G$  may be ignored for numerical purposes, in this MODEL it must be retained in a dimensional analysis. (The definition of the standard geopotential meter was in fact chosen to make the ratio  $g_0/G$  numerically unity for the case when  $g_0 = 9.80665$  m sec<sup>-2</sup>, the standard sea-level value of gravity in the ICAO Standard Atmosphere and in this MODEL.)

Using equation (5), the following tables of geopotential in  $m^2$  sec<sup>-2</sup>, as well as in standard geopotential meters, have been prepared for specified geometric altitudes.

Geometric	Geopoten	tial	Differences in Values of H
Altitude Z	ΔE/m	н	values of h
m	m <sup>2</sup> sec <sup>-2</sup>	m' by equation (5)	m.º
x 10 <sup>0</sup>	9.806,648,45 x 10 <sup>0</sup>	.999,999,839 x 10 <sup>0</sup>	.000,000,0
x 10 <sup>1</sup>	9.806,634,56 x 10 <sup>1</sup>	.999,998,423 x 10 <sup>1</sup>	.000,000,0
x 10 <sup>2</sup>	9.806,495,72 x 10 <sup>2</sup>	.999,984,265 x 10 <sup>2</sup>	.000,000,0
x 10 <sup>3</sup>	9.805,107,53 x 10 <sup>3</sup>	.999,842,719 x 10 <sup>3</sup>	.000,000,0
x 10 <sup>4</sup>	9.791,247,11 x 10 <sup>1</sup>	.998,429,339 x 10 <sup>4</sup>	.000,07
L x 10 <sup>5</sup>	9.654,768,23 x 10 <sup>5</sup>	.984,512,367 x 10 <sup>5</sup>	•088
5 x 10 <sup>6</sup>	4.545,771,23 x 10 <sup>6</sup>	.463,539,663 x 10 <sup>6</sup>	9•9
L <b>x</b> 10 <sup>6</sup>	8.473,638,99 x 10 <sup>6</sup>	.864,070,707 x 10 <sup>7</sup>	70.6

Equations (4) through (6) do not represent the only possible equations for converting geometric measure to geopotential measure. While equation (2d) is the fundamental and rigorously correct equation for converting geopotential measure to geometric measure, equations (4) through (6) are only as good as the expression for g introduced into equation (2d). A more precise expression for g is discussed in Appendix N. This expression is an alternating infinite-power series in terms of latitude and altitude. Evaluating this expression for latitude 45° 32' 40" and introducing it into equation (2d) yields another alternating power series as the expression for H in terms of Z. The departures of the result of equation (5) from the results of this more exact

method are small. The differences in the values of H computed by both methods for  $45^{\circ}$  32'  $40^{\circ}$  latitude are given in the above table. For altitudes of  $1 \times 10^{3}$  meters and below, the number of significant figures limits difference determinations. For altitudes above  $8 \times 10^{0}$  meters, the number of available terms in the series limits the difference determinations. From these results it is obvious, however, that for practical applications, at least, equations (h) and (5) are quite adequate. (See appendix P)

#### 2.2.6 Other special units of geopotential

Two other special units of geopotential, neither of which is employed in this MODEL, preceded the standard geopotential meter. The geodynamic meter, the first of such units to be used, was defined by Bjerknes<sup>3</sup> to be equal to 10 joules kg<sup>-1</sup>. Thus a geodynamic meter differed in magnitude from a geometric meter by about 2% at sea level.

The second special unit of geopotential to be introduced, and the one generally used by meteorologists, is the geopotential meter<sup>23,32</sup> equal to 9.8 joules kg<sup>-1</sup> or 9.8 m<sup>2</sup> sec<sup>-2</sup>. This latter unit was defined on the basis of a sea-level value of g equal to 9.8 m sec<sup>-2</sup>. The numerical differences between altitudes measured in geopotential meters and the same altitudes expressed in standard geopotential meters are small, of the order of 1/10 of 1 per cent, and in many instances may be neglected.

#### 2.2.7 Analytical usage

Geopotential has its greatest appeal, for use in this MODEL, from an analytical point of view, because it is a parameter involving both g and Z, and hence its use reduced by one the number of variables in the differential form of the barometric equation relating the basic atmospheric properties of this MODEL. This reduction in the number of variables comes without requiring the erroneous assumption of constant acceleration of gravity, used in some of the earlier standards. (The constant gravity assumption would result in a computed pressure which, at 500 km, is 40 per cent lower than one finds when variations in gravity are accounted for.) This pressure discrepancy is equivalent to an altitude discrepancy of 42.6 km at 500 km. If variable gravity is retained in the hydrostatic equation explicitly, rather than being concealed in the geopotential altitude, the algebraic expression resulting from the integration of the hydrostatic equation is excessively complicated.

#### 3. Basic Atmospheric Properties of the MODEL

The basic properties of this ARDC MODEL are those properties rigorously related by the hydrostatic equation and the equation of state (perfect gas law). These are pressure, density, and the ratio of temperature to molecular weight of air (which will be expressed in terms of molecular-scale temperature). Defining the altitude function of any one of these properties specifies the remainder of these basic properties in any model. In this MODEL, according to custom, the temperature function is the defining property.

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## 3.1 Molecular-Scale Temperature and Its Development

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## 3.1.1 Ratic of temperature to molecular weight, T/M

The property, T/M, is a composite of two variables which are conveniently handled as an entity because of the frequent occurrence of this ratio in atmospheric equations. In fact, the occurrence is so frequent and so fundamental that all so-called atmospheric temperature measuring experiments successfully used in rockets to date measure T/M, rather than T independently.

The combining of the two variables into a single parameter is of particular convenience in the computation of atmospheric tables to great altitudes because:

a. The values of T and M have not been independently measured above 90 km with any degree of reliability; and

b. The introduction of T/M, as a single function of H, into the differential form of the barometric equation greatly simplifies the integration and resulting algebraic computational equations over the case when two irdependent functional relationships are used.

Until recently, aerologists have not been concerned with relating pressure-altitude gradients or speed of sound etc., to the ratio T/M, since within the altitude region of their concern (below about 90 km), the molecular weight of air, M, is known to remain essentially constant at its sealevel value, M<sub>o</sub>. For the same reason, the preparation of tables of atmospheric models and standards did not require the consideration of M as a variable; and hence the increased complexity of equations resulting from considering M a variable was not a problem. Defining the atmosphere in terms of T/M instead of in terms of T alone solves the problem of complexity but introduces the problem of consistency with existing standards. This consistency problem is solved by defining a new property, the molecular-scale temperature, such that it is a function of T/M and is equal to T at all altitudes where M is equal to M<sub>o</sub>.

## 3.1.2 Molecular-scale temperature concept

The molecular-scale temperature,  $T_M$ , which Minzner 10,41 suggested as the basic parameter for the Standard Atmosphere, is a parameter which combines the ratio of two fundamental variables T/M with a constant in such a manner that  $T_M$  is equal to T wherever  $M = M_O$ , and simultaneously accounts for variations in M without specifying its functional variation. Molecular-scale temperature is that temperature derived from essentially all rocket experiments when variations in molecular weight from its sea-level value are unknown and hence neglected. Molecular-scale temperature is an amplification and redefinition of Whipple's  $T_{29}$  in the Rocket Panel Atmosphere. 45 Analytically  $T_M$  is defined by the following equation:

$$\mathbf{T}_{\mathbf{H}} = \left(\frac{\mathbf{T}}{\mathbf{H}}\right) \mathbf{H}_{\mathbf{0}} , \qquad (7)$$

where

- T = temperature (kinetic) in the absolute thermodynamic scales,
- T<sub>M</sub> = molecular-scale temperature in the absolute thermodynamic scales,
- M = molecular weight (nondimensional),
- Mo = sea-level value of molecular weight equal to 28.966 (nondimensional, exact)?6-28,31,47 (See section 5.1.)

The use of  $T_M$  in the ARDC MODEL retains consistency with the existing United States Standard Atmosphere, since over the altitude region of the Standard (0 to 20,000 m<sup>1</sup>) as well as to considerably greater altitudes, the ratio of  $M_O/M$  is unity; and hence  $T_M = T$  for these altitudes.

3.1.3 Form of altitude function of molecular-scale temperature

Molecular-scale temperature is the key or defining property of this MODEL, in that the specification of the variation of  $T_M$  with altitude simultaneously and completely establishes the altitude variation of more than half of the fifteen properties of this MODEL. (The determination of the remaining properties requires a definition of the altitude variation of molecular weight above 90 km in addition to the altitude variation of the molecular-scale temperature.)

In accordance with precedent 26-28 and by agreement of the Working Group on Extension to the Standard Atmosphere, 18 the temperature parameter of this MODEL is defined to be a continuous function of altitude consisting of a consecutive series of functions linear in geopotential H, whose first derivatives are discontinuous at the intersections of the linear segments. The use of such a function implies that the atmosphere is made up of a finite number of concentric layers, each layer characterized by a specific constant value of the slope of the temperature parameter with respect to altitude. This slope will hereinafter be referred to as the gradient. The following is the general form of each segment of the function:

$$T_{M} = (T_{M})_{b} + I_{M}(H - H_{b}),$$
 (8)

where

H = geopotential altitude in m!,

T<sub>M</sub> = the molecular-scale temperature in <sup>O</sup>K at altitude H,

Im the gradient of the molecular-scale temperature in terms of geopotential altitude; i.e., \( \partial T\_{M} \sigma \text{B} \), in ok m:-1, constant for a particular layer,

H<sub>b</sub> = geometric altitude in m; at the base of a particular layer characterized by a specific value of I<sub>p(j)</sub> and

 $(T_{\text{M}})_{\text{b}}$  = the value of  $T_{\text{M}}$  at altitude  $H_{\text{b}}$ .

3.1.4 Kelvin or absolute temperature scale

In agreement with Resolution 164 of the 1947 meeting of the International Meteorological Organization, 31 and consistent with the ICAO Standard Atmosphere, the absolute temperature in degrees Kelvin of the melting point of ice subjected to atmospheric pressure of 1013.25 mb (or 101,325. newtons m<sup>-2</sup>) is taken\* to be  $T_i = 273.16^{\circ}K$ . Temperatures on the absolute Kelvin scale are related to temperatures on the Celsius scale\*\* by the relationship:

$$T(^{\circ}K) = T_{i} + t(^{\circ}C),$$
 (9)

where

T<sub>i</sub> = ice-point temperature, 273.16°K (exact),

t(°C) = temperature in the thermodynamic Celsius scale.

The magnitude of Kelvin degree and the Celsius degree are equal and hence temperature gradients are numerically the same in both systems.\*\*

- The Tenth General Conference on Weights and Measures has adopted 273.15°K for t<sub>i</sub> but this value will not be used in this MODEL.
- \*\* For relations between the two metric and two English temperature scales commonly used in scientific and engineering fields refer to Appendix C.

Basic constant

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## 3.1.5 Specific altitude function of molecular-scale temperature

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In accordance with the ICAO Standard Atmosphere,  $(T_M)_o$ , the sea-level value of  $T_M$ , is taken to be  $15^{\circ}C$  (exact) or  $288.16^{\circ}K$  (exact) by equation (9). This sea-level temperature plus the values of  $I_M$ , and the extent of the respectively associated layers completely define the profile of molecular-scale temperature with respect to altitude. The following are the values of  $I_M$  and their respectively associated altitude layers employed in this MODEL.

Table of Molecular-Scale Temperature Gradients Versus Altitude

L <sub>M</sub> in °K m;-1	Atmospheric Layers in m
-0.0065 exact	-5,000 to 0
-0.0065 exact	0 to 11,000
0.0 exact	11,000 to 25,000
+0.003 exact	25,000 to 47,000
0.0 exact	47,000 to 53,000
-0.0039 exact	53,000 to 75,000
0.0 exact	75,000 to 90,000
40.0035 exact	90,000 to 126,000
40:0100 exact	126,000 to 175,000
40.0058 exact	175,000 to 500,000

These values of  $L_{\rm H}$ , together with equation (8), imply ten specific functions of H to define  $T_{\rm H}$  over the desired altitude intervals. This molecular-scale temperature profile results in the following values of molecular-scale temperature  $(T_{\rm H})_{\rm b}$  associated with the base of the respective layers,  $H_{\rm b}$ :

Base Altitudes and the Respective Base Values of Molecular-Scale Temperatures

H <sub>b</sub> in m <sup>s</sup>	(T <sub>M</sub> ) <sub>b</sub> in <sup>o</sup> l
0	288.16
11,000	216.66
25,000	216.66
47,000	282.66
53,000	282.66
75,000	196.86
90,000	196.86
126,000	322.86
175,000	812.86

<sup>/</sup> Entire table consists of basic constants.

# 3.1.6 Basis for selecting the temperature-altitude function

The temperature-altitude function of this MODEL was selected to be in exact agreement with the present ICAO Standard Atmosphere which extends from -5,000 m¹ to 20,000 m¹. (The temperature-altitude function is also in agreement with the recently adopted Extension of the Standard Atmosphere to 300,000 m¹ which was prepared concurrently with this MODEL.) The values of the function between 20,000 m¹ and 53,000 m¹ were suggested by Whipple and adopted at the First Meeting18 of the WGESA. Between 53,000 m¹ and 500,000 m¹, the temperature-altitude function is that presented by Minzner20,42 and adopted to 300,000 m¹ for the Standard Atmosphere at the Third Meeting of the WGESA.

The linearized temperature-altitude function of this MODEL follows approximately along the average of observed temperatures up to about 90 or 100 km, the highest altitude for which "direct" temperature observations have been reliably made. The pressures and densities inferred by this linear-lized temperature-altitude function at the various altitudes agree very well with the average of all measured pressures and densities up to 160 km, the maximum altitude of such observations. Agreement between the inferred pressures or densities and the average of observed values was, in fact, the primary criterion for choosing the temperature-altitude function between 70 and 160 km.

Above 160 km, only theoretical approaches are presently available for estimating temperatures, pressures, or densities. Between 160 and 300 km, this MODEL represents an approximate mean value of the recent theoretical estimates of these properties.

For the region above 300 km, there are two basic theories on which to base a temperature-altitude profile. This MODEL follows that theory which results in the higher atmospheric densities at 500 km.

One of these theories, fostered principally by Bates, 1,2 assumes an upward conduction of energy from layers of high solar energy absorptivity, between 100 and 250 km. The proponents of this theory generally deduce an essentially isothermal atmosphere at a temperature between 850° and 1100°K extending upward from 250 or 300 km.

A second theory, proposed by Chapman, 8-10 suggests that the earth is bathed in the solar corona which extends outward from the sun beyond the earth's orbit around the sun. Some of the energy of the very high-temperature (high-velocity) particles comprising the corona, through which the earth is said to move in its orbin, is conducted downward toward the earth's surface. Thus a temperature of the order of 2 x 105 °K, a few earth's radii away from the earth, drops to the order of 1000°K at 300 km altitude as the conducted energy is shared by increasing numbers of particles. This theory, therefore, implies a positive real-temperature gradient which Chapman suggests might be of the order of 2.5°K per kilometer, in the 300 to 500 km region. This value corresponds closely with the molecular-scal? temperature gradient of 5.8°K/km used in that region of this MODEL.

Neither theory has any strong experimental support at present. The positive temperature-altitude gradient above 300 km was selected for this MODEL, however, because it inferred a higher atmospheric density at 500 km than is inferred by an isothermal atmosphere above 300 km. Higher densities in the vicinity of 500 km altitude are conservative from the point of view of satellite design.

#### 3.2 Pressure

3.2.1 Development of the general pressure-altitude equation

Atmospheric pressure is expressed as a function of altitude through the hydrostatic equation,

$$dP = -g \rho dZ, \qquad (10)$$

where

P = atmospheric pressure in newtons m<sup>-2</sup>,

g = acceleration of gravity in m scc-2,

 $\rho$  = atmospheric density in kg m<sup>-3</sup>, and

Z = altitude in m.

The density,  $\rho$ , may be eliminated by replacing it with its equivalent in terms of pressure and temperature in the form of the perfect gas law,

$$\rho = \frac{P!!}{R^{tk}T},\tag{11}$$

where

T = atmospheric temperature in OK, and

 $R^*$  = universal gas constant; i.e., 8.31439 x 103 joules ( ${}^{\circ}K$ )-1 kg-1 (exact).

The value of R\* was chosen to be in agreement with recent determinations of its value and consistent with the ICAO Standard Atmosphere.

The substitution of equation (11) into equation (10) plus some manipulation, leads to the differential form of the barometric equation,

# Basic constant

$$d \ln P = \frac{-gM}{R^{3}T} dZ . \qquad (12)$$

It is to be noted that the pressure is now expressed as a function of T/M. The introduction of molecular-scale temperature from equation (7) and geopotential from equation (2c) changes equation (12) in five variables to the following equation in only three variables:

$$d \ln P = \frac{-GM_0}{R^*} \frac{dH}{T_M}. \tag{13}$$

Equation (13) in turn leads to

$$\ln \frac{P}{P_b} = -Q \int_{H_b}^{H} \frac{dH}{f(H)}, \qquad (14)$$

where

 $P_b$  = pressure at altitude  $H_b$ ,  $Q = G M_o/R^*$ , a constant equal to 0.03h,16h,79h,2°K m;-1 # f(H) = a functional representation of  $T_{M^*}$ .

3.2.2 Pressure-altitude equations for linear temperature functions

For purposes of this MODEL, f(H) is defined by equation (8). Thus the integration of equation (14) yields two different forms of the barometric equation, depending on whether  $I_{\underline{M}}$  of equation (8) is equal to zero or equal to a non-zero constant:

For  $I_{H} = 0$ ,  $P = P_{b} \text{ exponential } \frac{-Q(H - H_{b})}{(T_{M})_{b}}; \qquad (15)$ 

For L not equal to zero,

$$P = P_{b} \left[ \frac{(T_{M})_{b}}{(T_{H})_{b} + I_{M}(H - H_{b})} \right]^{\frac{Q}{I_{M}}};$$
 (16)

# Derived constant

(T<sub>M</sub>)<sub>b</sub> = the value of molecular-scale temperature in <sup>O</sup>K at the base of a layer characterized by a constant value of I<sub>M</sub>,

 $I_M$  = the value of  $T_M/H$  in  ${}^{O}K$  m'-1 for a particular altitude region.

The forms of equations (15) and (16) are such that pressure may be computed in any units merely by introducing  $P_{\rm b}$  in terms of the desired units. For numerical computation purposes equation (15) is more usable in the form

$$P = \frac{P_b}{\text{antilog}_{10} \frac{\log_{10} e^Q}{(T_M)_b} (H - H_b)}$$
 (17)

where

 $log_{10}e = .434,294,482$ //, the modulus of common logarithms.

### 3.2.3 Sea-level value of pressure

Pressures at all altitudes computed from equation (15) or (16) depend directly on the sea-level value of pressure. In keeping with the ICAO Standard Atmosphere26-28 and implicit in the Resolution of the Proceedings of the International Committee on Weights and Measures, 14 the sea-level value of pressure, Po, is taken to be 101,325 newtons m-2 or 1,013.25 mb. This pressure corresponds to the pressure exerted by a column of mercury 760 mm high having a density of 13.595,1... gm cm-3 and subject to a gravitational acceleration of 9.80665 m sec-2.

#### 3.2.4 Base pressures for various layers

With  $P_0$  used for  $P_b$  in equation (16) and using suitable values of  $(T_M)_b$  and  $I_M$ , the value of P is computed for 11,000 m', the top of the troposphere, the first atmospheric layer above sea level. This value of P, designated by  $P_{11}$ , in turn becomes the value of  $P_b$  for use in computing the pressure within and at the top of the next layer. In this way the values of  $P_b$  for each successive layer are determined. The value adopted in this MODEL for  $P_0$ , i.e., 1,013.250 mb or 101,325.0 newtons  $m^{-2}$  (exact) is identical to that adopted by ICAO and other prominent groups.31,46

7 Basic constant

fff Numerical constant

# 3.2.5 Specific computational equations

The specific equations for computing pressure for each of ten atmospheric layers (determined by ten molecular-scale temperature functions) are as follows:

For -5,000.0 
$$m^{1} \le H \le 0.0 m^{1}$$
,

$$P = P_0 \left[ \frac{288.160 - 6.500,00 \times 10^{-3} H}{288.160} \right]^{5.256,122,18}$$
 (16a)

where

Po= atmospheric pressure at sea level, defined to be 101,325.0 newtons m-2, or 1,013.25 mb (exact).

For 0.0 m' \( \frac{1}{2} \) \( \frac{1}{2} \) 11,000 m',

$$P = \frac{P_0}{\left[\frac{288.160}{288.160 - 6.500, \infty \times 10^{-3}H}\right]}$$
 (16b)

For 11,000 m; \$\lefth{\lefth}{4} \lefth{\lefth}{25,000 m},

$$P = \frac{P_{11}}{\text{antilog}_{10} \left[ (0.068, 483, 253, 0 \times 10^{-3}) (H - 11,000.0) \right]}, \quad (17a)$$

where

P<sub>11</sub> = the pressure at 11 km' computed from equation (16b).

For 25,000 m; ≤ H ≤ 47,000 m;,

$$P = \frac{P_{25}}{\left[\frac{11.660 + 3.000,00 \times 10^{-3}H}{216.660}\right]^{11.388,264,73}},$$
 (16c)

# Basic constant

P<sub>25</sub> = the pressure at 25 km<sup>1</sup> computed from equation (17a).

For 47,000 m' ≤ H ≤ 53,000 m',

$$P = \frac{P_{47}}{\text{antilog}_{10} \left[ (0.052,492,682,3 \times 10^{-3})(H - 47,000.0) \right]},$$
 (17b)

where

 $P_{47}$  = the pressure at 47 km computed from equation (16c).

For 53,000 m' ≤ H ≤ 75,000 m',

$$P = \frac{P_{53}}{\left[\frac{282.660}{489.360 - 3.900,00 \times 10^{-3}H}\right]^{8.760,203,64}},$$
 (16d)

where

P<sub>53</sub> = pressure at 53 km computed from equation (17b).

For 75,000 m' ≤ H ≤ 90,000 m',

$$P = \frac{P_{75}}{\text{antilog}_{10} \left[ (0.075,371,236,4 \times 10^{-3})(H - 75,000.0) \right]}$$
(17c)

where

P = the pressure at 75 km' computed from equation (16d).

For 90,000 m; \( \frac{1}{2} \) \( \frac{1}{2} \) (000 m;

$$P = \frac{P_{90}}{\left[\frac{3.500,00 \times 10^{-3} H - 118.140}{196.860}\right]^{9.761,369,77}},$$
 (16e)

 $P_{90}$  = the pressure at 90 km² computed from equation (17c).

For 126,000 m<sup>2</sup>  $\leq H \leq$  175,000 m<sup>2</sup>,

$$P = \frac{P_{126}}{\left[\frac{10.000,0 \times 10^{-3}H - 937.140}{322.860}\right]^{3.416,479,42}},$$
(16f)

where

P = the pressure at 126 km computed from equation (16e).

For 175,000 m: ≤ H ≤ 500,000m:,

$$P = \frac{P_{175}}{\left[\frac{5.800,00 \times 10^{-3}H - 202.140}{812.860}\right]},$$
 (16g)

where

P<sub>175</sub> = the pressure at 175 km computed from equation (16f).

3.3 Density

3.3.1 Computational equation

Atmospheric density at altitude H is readily computed from the perfect gas law, equation (11), implicit in the barometric equation. With the introduction of the molecular-scale temperature concept, equation (11) for density in kg m-3 becomes,

$$\rho = \frac{H_0}{R^8} \cdot \frac{P}{T_M} = 3.483,839,46 \times 10^{-3} \cdot \frac{P}{T_M}. \tag{18}$$

 $\nabla$ 

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where

P = atmospheric pressure in newtons m<sup>-2</sup> (or mb x 10<sup>2</sup>), expressed by equations (16a - 16g) and (17a - 17c),

T<sub>M</sub> = molecular scale temperature in <sup>O</sup>K expressed by equation (8) with its various values of I<sub>M</sub>.

The computational equation for  $\rho$  is left in terms of P and T<sub>M</sub> instead of in terms of H, for to convert to the latter would require ten different functions, as in the case of T<sub>M</sub> and P. The computational equations of all other properties of this MODEL will be similarly expressed in terms of P or T<sub>M</sub>, rather than in terms of H.

3.3.2 Sea-level value - ratio equation

Evaluating equation (18) at sea level yields the sea-level value of density:

$$P_{o} = \frac{H_{o}}{R^{*}} \cdot \frac{P_{o}}{(T_{M})_{o}} = 1.225,013,998 \text{ kg m}^{-3}, \ \mathcal{H}$$
 (18a)

where

Po = sea-level value of P, 101,325.0 newtons m-2 (exact), and

 $(T_{\rm M})_{\rm o}$  = sea-level value of  $T_{\rm M}$ , 288.16°K (exact).

Dividing equation (18) by equation (18a) yields

$$\frac{\rho}{\rho_o} = \frac{P}{P_o} \cdot \frac{(T_M)_o}{T_M} . \tag{18b}$$

3.4 Validity of the Basic Properties

The three basic properties of this atmospheric MODEL are rigorously self-consistent through the perfect gas law and the hydrostatic equation, which accounts for the variations of the effective acceleration of gravity with altitude,

<sup>#</sup> Basic constant

<sup>//</sup> Derived constant

through the use of geopotential. The user of these tables is warned that the validity of the hydrostatic equation as well as some of the other classical equations, in their simple forms, may decrease considerably at great altitudes. The uncertainties at high altitudes in most equations relating the various atmospheric properties, however, are perhaps small compared with the present uncertainties at these altitudes in the defining property of this MODEL, T/M.

# 4. Secondary Properties Defined as Functions of T/M

This section is devoted to all those atmospheric properties of the ARDC MODEL ATMOSPHERE, except P and  $\rho$ , which are classically defined as functions of the ratio T/M and which are, therefore, conveniently redefined in terms of molecular-scale temperature without otherwise involving M or T explicitly. (Some of the properties of this group depend also upon the acceleration of gravity.) Properties which depend also upon P or  $\rho$ , or combinations of these, are implicitly in this group. The properties of this group tabulated in this MODEL are scale height, speed of sound, air-particle speed (arithmetic average), and specific weight.

### 4.1 Scale Height

4.1.1 Definition

If both sides of equation (12) are divided by dZ, we have

$$\frac{d \ln P}{dZ} = \frac{-gM}{R^{2}T} \qquad (12a)$$

A dimensional analysis of the quantities in the right-hand side of this equation show that the net dimensions are reciprocal meters. The reciprocal of the right-hand side of equation (12a), by virtue of its dimensions has been given the name "scale height." Thus scale height as tabulated in this MODEL is defined as

$$H_{\mathbf{S}} = \frac{R^*T}{gM} , \qquad (19)$$

where

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(G)

H = scale height in m (not m'),

g = acceleration of gravity in m sec<sup>-2</sup>,

and R\*T and M have their usual significance.

### 4.1.2 Concepts

Using equation (19), equation (12a) may now be rewritten as

$$\frac{d \ln P}{dZ} = \frac{-1}{H_B} \quad , \tag{12b}$$

and scale height is seen to be the negative reciprocal of the slope of the ln P versus Z curve.

The geometric-altitude-pressure equation for an isothermal atmospheric layer may be manipulated to show that when gravity is considered to be constant, the scale height at any altitude represents the vertical distance above the reference altitude at which the atmospheric pressure has dropped to a value of 1/o of its value at the reference altitude. This concept for scale height is often erroneously thought to apply to an atmosphere in which temperature and gravity vary. A check of pressures and scale heights in the troposphere of this MODEL shows the scale height at sea level to be 8.4344 km. The pressure, however, has dropped to 1/e of its sea-level value at an altitude of 7.68 km, where the scale height is 7.0 km. Since this concept of scale height is developed from the equation for an isothermal constant-gravity atmosphere, the concept will not hold for other conditions.

From the same basic, isothermal, pressure-altitude equation one may demonstrate that the scale height at any altitude is the length to which the total of a unit cross-section column of the atmosphere above that point would be compressed, if subjected to the pressure and gravity of that altitude. That is, the reduced thickness of the residual, isothermal, constant-gravity atmosphere above a given altitude, when subjected to the pressure of that altitude, is equal to the scale height. Again this concept does not apply rigorously anywhere in this MODEL since the atmosphere is not indefinitely isothermal above any point, neither is the gravity constant.

## 4.1.3 Definition of geopotential scale height

The limitations imposed by constant gravity in the latter two concepts of scale height can be eliminated through the use of a geopotential scale height. If both sides of equation (13) are divided by dH, we obtain

$$\frac{\mathrm{d} \ln P}{\mathrm{dH}} = \frac{-\mathrm{G} M_0}{\mathrm{R}^{\mathrm{H}} T_{\mathrm{M}}} , \qquad (13a)$$

A dimensional analysis of the right-hand side of this equation shows the net dimensions to be reciprocal geopotential meters. Thus the reciprocal of this equation serves to define geopotential scale height:

$$H_{\mathbf{S}}^{\dagger} = \frac{R^{\mathsf{H}}T_{\mathsf{H}}}{G H_{\mathsf{A}}} , \qquad (13b)$$

 $F_s$  = geopotential scale height in m', and  $F_s$  = 9.80665 m<sup>2</sup> sec<sup>-2</sup> m'<sup>-1</sup>.

4.1.4 Concept of geopotential scale height

The combining of equations (13a) and (13b) yields

$$\frac{d \ln P}{dH} = \frac{-1}{H_g!}, \qquad (13c)$$

and the geopotential scale height is seen to be the negative reciprocal of the slope of the ln P versus H curve.

The manipulation of equation (15) (for a variable-gravity, isothermal atmosphere) leads to the conclusion that for a variable-gravity, isothermal atmosphere, the geopotential scale height at any altitude represents the increment in geopotential above the reference altitude at which the atmospheric pressure has dropped to a value of 1/2 of its value at the reference altitude. This concept does apply rigorously to isothermal regions of this MODEL. Equation (15) also leads to the conclusion that the geopotential scale height at any altitude is the reduced thickness in geopotential of the residual, isothermal, variable-gravity atmosphere above a given altitude when subjected to the pressure of that altitude. Even though this concept accounts for variable gravity, it still is not rigorously applicable to the MODEL since no indefinite isothermal atmosphere to great altitudes is speculated in this MODEL.

The geopotential scale height at any altitude is readily transformed to a geometric length by adding the geopotential scale height to the reference geopotential altitude and converting the resulting geopotential measure to geometric altitude, by means of equation (6). Then the reference geopotential altitude is converted to geometric altitude with the same equation. Finally, the smaller geometric altitude is subtracted from the larger. The difference is the equivalent geometric length for the geopotential scale height at the reference altitude.

While geopotential scale height is obviously the preferable parameter from the point of view of using the several concepts in a variable-gravity atmosphere, only geometric scale height from equation (19) will be tabulated in this edition of the ARDC MODEL.

4.1.5 Computational equation for (geometric) scale height

Introducing  $T_{H}$  from equation (7) into equation (19) leads to the computational equation for  $H_{g}$ :

$$H_3 = \frac{R^{*}T_{H}}{M_{0}g} = 287.039,632,6 \left[\frac{T_{H}}{g}\right].$$
 (19a)

4.1.6 Sea-level value and ratio equation

The sea-level value of  ${\rm H_g}$  is obtained by evaluating equation (19a) at sea level, such that

$$(H_s)_o = \frac{R^*(T_M)_o}{H_o g_o} = 8.434,413,43 \times 10^3 \text{ m} \text{ //}$$
 (19b)

where

$$(H_s)_o$$
 = sea-level value of  $H_s$ ,  
 $(T_M)_o$  = sea-level value of  $T_M$ , 288.16°K (exact),  
 $g_o$  = sea-level value of  $g_s$ , 9.806,65 m sec<sup>-2</sup> (exact).

Dividing equation (19a) by (19b) yields

$$\frac{H_s}{(H_s)_o} = \frac{T_M}{(T_M)_o} \frac{g_o}{g} , \qquad (19c)$$

which is an alternate form for computing values of Hs.

#### 4.1.7 Validity

Because the analytical expression for scale height is implicit in the barometric equation, as is evident from equation (12), the validity of the value of  $H_{\rm S}$  at various altitudes depends directly on the validity of the barometric equation. (Scale height from this consideration might also be considered one of the basic properties along with pressure and density.) The use

<sup>≠</sup> Basic constant

<sup>//</sup> Derived constant

of the tabulated values of scale height, however, in connection with several commonly accepted concepts of scale height is to be avoided except for rough approximations.

## 4.2 Speed of Sound

4.2.1 Defining equation

The square of the speed of sound propagation is defined in

this MODEL to be

$$c_{s}^{2} = \frac{\gamma_{P}}{\rho}, \qquad (20)$$

where

C = speed of sound in m sec-1,

P = pressure in newtons n=2,

 $\rho$  = density in kg m<sup>-3</sup>, and

ratio of specific heat of air at constant pressure to the specific heat of air at constant yolume, defined to be 1.4 (dimensionless, exact.)

## 4.2.2 Computational equation

Eliminating  $\rho$  between equations (18) and (20) and extracting the square root results in:

$$c_s = \left[\frac{\gamma_R^*}{M_A}T_M\right]^{\frac{1}{2}} = 20.046,333,47 \left(T_H\right)^{\frac{1}{2}}.$$
 (20a)

4.2.3 Sea-level value and ratio equation

Evaluating equation (20a) at sea level yields

$$(C_s)_o = \left[\frac{\gamma_R^*}{M_o} \cdot (T_M)_o\right]^{\frac{1}{2}} = 340.292,046 \text{ m sec}^{-1},$$
 (20b)

where

A Basic constant

<sup>//</sup> Derived constant

Dividing equation (20a) by equation (20b) reduces the number of constants so that:

$$\frac{C_s}{(C_s)_o} = \left[\frac{T_M}{(T_M)_o}\right]^{\frac{1}{2}}.$$
 (20c)

## 4.2.4 Validity

These equations for computing the velocity of sound apply only when the sound wave is a small perturbation on the ambient condition. Harrison24 has shown that even when this condition is met, the above definition for the velocity of sound is not quite correct for two reasons: First, Y is not really a constant, but rather, varies with pressure and temperature over a small region around the value 1.4; second, the form of the above relationship is not completely correct, since even if the best value of Y is used for a given set of conditions, computed values of C<sub>S</sub> differ slightly from experimentally determined values. In spite of these discrepancies, however, the stated relationships are adopted in accordance with Subcommittee recommendations43 which are in conformity with established aerodynamic practice but at variance with the present United States Standard Atmosphere.

The limitations of the concept of velocity of sound due to extreme attenuation are also of concern. This situation exists for high frequencies at sea-level pressures and applies to successively lower frequencies as atmospheric pressure decreases, or as mean free path increases. For this reason the concept of speed of sound progressively loses its meaning at high altitudes, except for frequencies approaching zero and for very short distances. To call attention to this limitation, it was agreed to terminate at 90 km² the tabulation of the velocity of sound, in the Extension to the United States Standard Atmosphere. In conformity with this agreement, tabulations in this MODEL are also similarly terminated. Because of the relationship between sound velocity and air particle speed (Section 4.3), sound velocities for altitudes above 90 km² may readily be obtained for use with suitable caution.

#### 4.3 Air Farticle Speed (Arithmetic Average)

### 4.3.1 Concept

The mean air particle speed is the arithmetic average of the distribution of speeds of all air particles within a given elemental volume. This quantity has significance provided that the volume considered contains a sufficiently large number of particles so that their velocities follow a Maxwellian distribution, and provided that variations of  $\rho$  and T/M in any direction are negligible within the volume element.

#### 4.3.2 Defining equation

Arithmetic average of air particle speed is defined to be:

$$\vec{\nabla} = \left[ \frac{8 R^*}{\pi} \frac{T}{H} \right]^{\frac{1}{2}}, \qquad (21)$$

1

where

 $\overline{V}$  = air particle speed (arithmetic average) in m sec<sup>-1</sup>,  $\pi$  = 3.141,592,654 (dimensionless).

4.3.3 Computational equation

The introduction of  $T_M$  from equation (7) into equation (21) yields the computation equation for  $V_1$ 

$$\vec{v} = \left[\frac{8R^*}{\pi M_{\bullet}} T_{M}\right]^{\frac{1}{2}} = 27.035,909,86 (T_{M})^{\frac{1}{2}}$$
 (21a)

4.3.4 Sea-level value and ratio equation

Evaluating equation (21a) at sea level leads to

$$\vec{v}_{o} = \left[\frac{8 R^{*}}{\pi M_{o}} (T_{M})_{o}\right]^{\frac{1}{2}} = 458.942,035 \text{ m sec}^{-1}$$
 (21b)

where

Equation (2la) divided by equation (2lb) yields

$$\frac{\vec{v}}{\vec{v}_{\bullet}} = \left[\frac{T_{\rm H}}{(T_{\rm H})_{\bullet}}\right]^{\frac{1}{2}} \tag{21c}$$

4.3.5 Validity

On considering the restrictions applied to the volume element for which we desire the value of  $\vec{v}$ , it is evident that these restrictions come

// Derived constant

/// Numerical constant

into conflict with each other at high altitudes and the validity of the concept of V decreases with altitude. It is uncertain whether or not the concept retains reasonable significance at altitudes as great as 500 km. Nevertheless, as in the case of pressures and densities, etc., values have been tabulated to this altitude, on the basis that with suitable caution, such values are better than no values.

4.3.6 Relationship to sound velocity

From a comparison of equation (20c) and equation (21c) it is evident that

$$\frac{C_s}{(C_s)_{\bullet}} = \frac{\bar{v}}{\bar{v}_{\bullet}} . \tag{22}$$

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Since values of  $\overline{V}/\overline{V}_0$  are tabulated to 500 km², values of  $C_0/(C_8)$  and hence values of  $C_0$  are readily available to the same altitude, even though their significance is extremely questionable.

4.4 Specific Weight

4.4.1 Concept

The specific weight  $\omega$  of a body of uniform density at any particular point in space is the weight per unit volume of that body at that point. The weight per unit volume is equal to the mass per unit volume times the acceleration of gravity, which in turn is equal to the density of the body times the acceleration of gravity, g. Since g is assumed to vary in this MODEL in accordance with equation (la), the specific weight of a body will vary proportionately.

The density of the air mass also varies with altitude and hence  $\omega$  is dependent upon two variables,  $\rho$  and g. This is at variance with the procedure in the ICAO Standard Atmosphere in which specific weight is defined to vary only with  $\rho$ .

4.4.2 Defining and computational equation

In this MODEL specific weight is defined by

$$\omega = \rho g$$

where

- $\omega$  = specific weight in kg m<sup>-2</sup> sec<sup>-2</sup> or newtons m<sup>-3</sup>(at any point),
- $\rho$  = density in kg m<sup>-3</sup> (at the point),
- g = acceleration of gravity in m sec<sup>-2</sup> (at the point).

Eliminating p by means of equation (18) results in

$$\omega = \frac{g M_0 P}{R^{8}T_{M}} = 3.483,839,46 \times 10^{-3} \frac{g P}{T_{M}} . \tag{23a}$$

4.4.3 Sea-level value and ratio equation

The evaluation of equation (23) and (23a) at sea level yields

$$\omega_{o} = \rho_{o} g_{o} = \frac{M_{o}P_{o}g_{o}}{R^{*}(T_{M})_{o}} = 12.013,283,5 \text{ kg m}^{-2}\text{sec}^{-2},$$
 (23b)

where

$$ω_o$$
 = sea-level value of  $ω$  ,

 $ρ_o$  = sea-level value of  $ρ$  , 1.225,014,00 kg m<sup>-3</sup>,

 $g_o$  = sea-level value of g, 9.806,65 (exact).

Dividing equations (23) and (23a) by the appropriate portions of equation (23b) results in:

$$\frac{\omega}{\omega_o} = \frac{\rho}{\rho_o} \frac{g}{g_o} = \frac{P}{P_o} \frac{(T_M)_o}{T_M} \frac{g}{g_o}. \tag{23c}$$

Introducing H<sub>s</sub> from equation (19a) into the right-hand member of equation (23c) leads to:

$$\omega = \frac{PH}{R^2T} \cdot g = \frac{P}{H_g} \cdot \tag{23d}$$

4.4.4 Validity

The validity of the values of  $\omega$  depends only upon the validity of the values of g and  $\rho$  which have already been discussed.

<sup>#</sup> Basic constant

<sup>//</sup> Derived constant

### 5. Other Secondary Properties

The last group of properties of this ARDC MODEL ATMOSPHERE includes all those properties considered in this MODEL which are defined by functions of T and M, in forms different from T/M, so that these functions cannot be redefined in terms of molecular-scale temperature without the additional use of either M or T in its independent form. This group includes molar volume, number density, mean free path, collision frequency, coefficient of viscosity, and kinematic viscosity, as well as temperature and molecular weight. Either molecular weight or temperature must now be defined in terms of altitude before any of these remaining secondary properties can be computed. The molecular weight is the one specifically defined in this MODEL.

### 5.1 Molecular Weight

### 5.1.1 General definition

Molecular weight is defined to be dimensionless. On the chemical scale\* mclecular weight (of a compound) is defined to be 16 times the ratio of the average mass of a molecule of the compound to the average mass of an oxygen atom, where both the oxygen and the compound are assumed to have their natural distribution of isotopes, and where average is to be construed as the arithmetic mean.

### 5.1.2 Concept applied to air

The definition of molecular weight includes the concept of a mixture of the several isotopes of an atomic species and the resulting mixture of similar molecules of different masses. Therefore, it is not unreasonable to extend the definition of molecular weight to include mixtures of different kinds of molecules as in the atmosphere. Such an extension of the basic definition is employed in this MODEL in establishing the concept of the molecular weight of air.

The definitions of atomic or molecular weights on the physical scale are more specific than the equivalent definitions on the chemical scale, in that on the physical scale, the ratios are established with reference to the mass of an atom of a specific oxygen isotope, 016. Because the mass of an 016 atom is less than the mass of an average oxygen atom, the atomic or molecular weights on the physical scale are greater than on the chemical scale by approximately the ratio 32.0087/32.0000. When the physical scale is used for expressing molecular weight, values of the universal gas constant,  $R^*$ , and other constants must be proportionately changed.

# 5.1.3 Molecular weight of air and mole defined

Molecular weight of air, M, is defined as 16 times the ratio of the arithmetic mean mass of a single molecule of the air mixture to the arithmetic mean mass of a single atom of oxygen in a natural mixture of the several oxygen isotopes.

A kilogram mole of air is defined as a quantity of air having a mass in kilograms numerically equal to the molecular weight of the air.

5.1.4 Sea-level and low-altitude value of molecular weight of air

The value of M at sea level is determined from an assumed distribution of the several atmospheric constituents at sea level. In accordance with the ICAO agreements the atmosphere of this ARDC MODEL is assumed to be dry and to have the following composition at sea level and at all altitudes up to and including 20 km. This model has assumed a continuation of this composition up to 90 km.

Constituent Gas	Mol. Fraction Per Cent	$\frac{\text{Molecular Weight}}{(0 = 16.000)}$
Nitrogen (N <sub>2</sub> )	78.09	28.016
Oxygen (O <sub>2</sub> )	20.95	32.0000
Argon (A)	0.93	39 <b>.</b> 944
Carbon dioxide (CO <sub>2</sub> )	0.03	կ <b>4.</b> 010
Neon (Ne)	1.8 x 10 <sup>-3</sup>	20.183
Helium (He)	5.24 x 10 <sup>-4</sup>	4.003
Krypton (Kr)	1.0 x 10 <sup>-4</sup>	83.7
Hydrogen (H <sub>2</sub> )	5.0 x 10 <sup>-5</sup>	2.0160
Xenon (Xe)	8.0 x 10 <sup>-6</sup>	131.3
Ozone (0 <sub>3</sub> )	1.0 x 10 <sup>-6</sup>	48.0000
Radon (Rn)	$6.0 \times 10^{-18}$	222.

The above data yield a value of 28.966 (nondimensional) for the molecular weight of air. In this MODEL the molecular weight of air at sea level, and for

Y)

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a considerable altitude above and below sea level, is defined as a constant. Thus for  $-5,000 \text{ m}^{\frac{1}{2}} \text{ H} \leq 90.000 \text{ m}^{\frac{1}{2}}$ .

$$M = 28.966$$
 . (24)

5.1.5 Molecular weight of air at high altitudes and validity of the values

Atmospheric composition at high altitudes is thought to vary considerably from that near sea level. The variation in composition may result from dissociation of various molecules of the atmosphere as well as from diffusive separation of molecules of various masses in a gravitational field. While several theories describing these phenomena exist, there are only a few data to support or disprove these theories. The choice of 90,000 m as the top of the region of constant composition is quite arbitrary but is as good as any other current choice.

It is thought that the dissociation of O2 is the principal factor in producing a change in molecular weight between 90,000 and 175,000 m's. Rocket measurements of O2 concentration obtained by Byram, Chubb, and Friedman provide partial support to this contension. Diffusive separation and the dissociation of N2 is thought to dominate the variation of molecular weight of the mixture of atmospheric gases above 175,000 m²s.

Miller<sup>39</sup> combined these theories, assumptions, and data with scale height gradients of this MODEL and computed molecular weights for specific altitudes between 90,000 and 500,000 m. A plot of these data versus altitude suggested the possibility of approximating the graph with two analytical functions. Campen of GMD developed the desired functions in the form of the following two equilateral hyperbolae which for this MODEL define molecular weight from 90 to 500 km.

For 90,000 m;  $\leq H \leq 175,000 \text{ m}$ ,

$$M = \frac{23.160,126,7 \text{ H} - 1,757,856.05}{\text{H} - 78,726.25}.$$
 (24a)

For 175,000 m<sup>1</sup>  $\leq$  H  $\leq$  500,000 m<sup>2</sup>,

$$M = \frac{13.139,119,0 \text{ H} + 514,492.02}{\text{H} - 56,969.89}$$
 (24b)

For purposes of defining other atmospheric properties, it is convenient to

establish the following relationships:

$$M = |M^t|$$
, and (24c)

$$\frac{M}{M_0} = \frac{M^t}{M_0^t} , \qquad (24d)$$

where

M' is a kilogram mole of air, a mass in kg numerically equal to the molecular weight, and

M' is the sea-level value of M'.

Using equation (5), relating geopotential and geometric altitude, equations (21,), (24a) and (24b) are converted to the following in terms of Z:

For  $-4,996.070,27 \text{ m} \leq Z \leq 91,292.532,7 \text{ m}$ 

$$M = 28.966.$$
 (25)

For 91,292.532,7 m  $\leq$  Z  $\leq$  179,954.085 m,

$$M = \frac{23.170,552,5 \ Z - 1,779,899.46}{Z - 79,713.475,7}.$$
 (25a)

For 179,954.085  $m \le Z \le 542,685.673$ ,

$$M = \frac{13.339.605,8 \text{ Z} + 519.1 \text{h} \cdot 6 \text{h}}{\text{Z} - 57.1 \cdot 85.075.2}$$
 (25b)

These equations yield results within  $\pm$  1% of Miller's values at all altitudes except for a small region around 105 km where the analytical results are about 3% higher than Miller's values.

### 5.2 Mol Volume

#### 5.2.1 Concept and definition

Density of the air at any altitude is expressed as the mass per unit volume at that altitude. If the mass is that of a mole of air, the related volume is that of a mole of air. Thus the mol volume of air is given by

$$\mathbf{v} = \frac{\mathbf{H}^{t}}{\mathbf{Q}}, \tag{26}$$

- v = the volume (in m<sup>3</sup>) of a mole of atmospheric gas at a particular altitude,
- $\rho$  = the density (in kg m<sup>-3</sup>) of air at the same altitude, and
- M' = the kilogram molecular weight, the mass in kg of a kilogram mole of air having the composition of this altitude. (This mass is numerically equal to the molecular weight defined by equations (24), (24a), and (24b).)
- 5.2.2 Computational equation

Eliminating  $\rho$  between equations (18) and (26) yields a computational\* expression for v in terms of basic properties and constants:

$$v = \frac{R^{3}M^{1}T_{M}}{M_{0}P} = 287.039,632,6 \frac{M^{1}T_{M}}{P}$$
, (26a)

where

- $R^*$  = universal gas constant, 8.314,39 x 10<sup>3</sup> joules (°K)<sup>-1</sup> kg<sup>-1</sup> (exact)<sup>2</sup>,
- Mo = sea-level value of molecular weight, 28.966 (dimensionless, exact)
- T<sub>M</sub> = molecular scale temperature, in <sup>O</sup>K, at the altitude in question, and
- P = atmospheric pressure in newtons  $m^{-2}$  (or mb x  $10^2$ ).

<sup>\*</sup> Values of v are not tabulated for various altitudes in this edition of the MODEL but the equations are developed for use in the expressions for number density and implicitly mean free path. It will be noted from a comparison of equations (26c) and (28c) that  $v/v_0 = L/L_0$ . Thus values of v for any altitude are readily available from these tables.

<sup>#</sup> Basic constant

## 5.2.3 Sea-level value and ratio equation

Equations (26) and (26a) evaluated at sea level yield:

$$v_o = \frac{M_o^i}{\rho_o} = \frac{R^*M_o^i (T_M)_o}{M_o P_o} = 23.645, \text{ hhh}, 1 \text{ m}^3,$$
 (26b)

where

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 $\mathbf{v}_{\mathbf{a}}$  = the sea-level value of  $\mathbf{v}_{\mathbf{i}}$ 

M'o = a mole of air at sea level, 28.966 kg (exact)//,

 $\rho_0$  = sea-level value of  $\rho$  , 1.225,013,998 kg m<sup>-3</sup>, H

 $(T_H)_e$  = the sea-level value of  $T_M$ , 288.16°K (exact), and

 $P_0$  = the sea-level value of P, 101,325.0 newtons  $m^{-2}$ (exact).

From equations (24d), (26), (26a), and (26b) it is obvious that

$$\frac{\mathbf{v}}{\mathbf{v_o}} = \frac{\mathbf{H}^{\dagger}}{\mathbf{H}_o^{\dagger}} \cdot \frac{\rho_o}{\rho} = \frac{\mathbf{H}}{\mathbf{H}_o} \cdot \frac{\mathbf{T}_{\mathbf{M}}}{(\mathbf{T}_{\mathbf{M}})_o} \cdot \frac{\mathbf{P_o}}{\mathbf{P}} \cdot \tag{26c}$$

5.2.4 Ice-point value

The (standard) ice-point value\* of the volume of a mole of gas is considered to be one of the basic physical constants. This value may be computed by evaluating equation (27) at the ice point, i.e., at a temperature of 273.16° K and a pressure of 101,325.0 newtons m-2 (1013.250 mb),

$$v_{i} = \frac{M'_{o}}{P_{i}} = \frac{R^{*}H'_{o}(T_{M})_{1}}{M_{o}P_{o}} = 22.414,594,3 \text{ m}^{3}, \text{ } (26d)$$

// Derived constant

<sup>∠</sup> Basic constant

<sup>\*</sup> These conditions referred to as standard conditions by chemists are not to be confused with the standard sea-level values of the standard atmosphere where the  $T_0 = (T_M)_0 = 288.16$ .

 $v_i$  = the ice-point value of v, and

 $(T_{\rm M})_{\rm i}$  = the ice-point value of  $T_{\rm M}$  = 273.16° K (exact),

 $\rho_1$  = the ice-point value of  $\rho$  , 1.292,283,037 from the left-hand members of equation (26d).

The above value of  $v_i$  for a kilogram mole is in keeping with 22.4146 m<sup>3</sup>, the value currently accepted outside of the realm of this standard. (The latter is equivalent to 22,414.6 cm<sup>3</sup> for a gram mole.)

## 5.2.5 Validity

The validity of the concept of molar volume at great altitudes becomes vague because the volume becomes so large that density and molecular weight cannot be assumed to remain constant throughout the volume and hence the specified volume will most probably not contain exactly one mole of atmospheric gases.

#### 5.3 Number Density

## 5.3.1 Concept and definition

The number density of air is defined to be the number of atmospheric particles per unit volume, considering only neutral or ionized atoms or molecules. (Electrons and other subatomic particles are ignored.) The number of particles contained in a mole of air is by definition Avogadro's number. Thus Avogadro's number divided by the mol volume yields number density, i.e.:

$$n = \frac{N}{V}, \qquad (27)$$

where

n = atmospheric-particle, number density, at a specified altitude, in m=3,

v = mol volume at that altitude in m<sup>3</sup>, and

N = Avogadro's number, 6.023,80 x  $10^{26}$  (dimensionless, exact)  $\frac{16,46}{1}$ 

A more recent value of N might have been used but that would not be consistent with the current values adopted by the National Research Council.46

<sup>/</sup> Basic constant .

## 5.3.2 Computational equation

Introducing equation (26a) into equation (27) leads to that computational form of the expression for number density in terms of basic properties and constants:

$$n = \frac{N M_0 P}{R^2 M^2 T_M} = 2.098,595,21 \times 10^{24} \frac{P}{M^2 T_M}$$
 (27a)

## 5.3.3 Sea-level value and ratio equation

Upon evaluation of equation (27) and (27a) at sea level, one

obtains:

$$n_0 = \frac{N}{V_0} = \frac{NM_0P_0}{E^{8}M_0^{1}(T_{H})_0} = 2.547,552,07 \times 10^{25} \text{ m}^{-3},$$
 (27b)

where

no = the sea-level value of n,

vo = the sea-level value of v.

The manipulation of equations (27), (27a), and (27b) and reference to equations (26c) and (24d) show the following relationships to exist:

$$\frac{\mathbf{n}}{\mathbf{n}_{0}} = \frac{\mathbf{v}_{0}}{\mathbf{v}} = \frac{\rho}{\rho_{0}} \cdot \frac{\mathbf{M}_{0}}{\mathbf{M}} = \frac{\mathbf{M}_{0}}{\mathbf{M}} \cdot \frac{(\mathbf{T}_{\mathbf{M}})_{0}}{\mathbf{T}_{\mathbf{M}}} \cdot \frac{\mathbf{p}}{\mathbf{P}_{0}} \cdot \tag{27c}$$

## 5.3.4 Validity

In the form of equation (27) the validity of n would be open to considerable question at high altitudes. In terms of equation (27a), however, where all the parameters are defined at a point or within a volume considerably smaller than v, the validity of n is probably limited principally by the validity of the values of T<sub>M</sub> and M.

## 5.4 Hean Free Path

## 5.4.1 Concept and definition

Mean free path is the mean value of the distances traveled by each of the molecules of a given volume between successive collisions with other molecules of that volume, provided that a sufficiently large number of molecules are contained within the volume. It is usually considered necessary that the volume be the cube of a length many orders of magnitude greater than the mean free path. From kinetic theory and assuming a gas of uniform temperature and density, the following expression for mean free path is developed:

$$L = \frac{1}{\sqrt{2}\pi\sigma^2 n},$$
 (28)

where

L = mean free path in m at a particular altitude,

n = number density in m<sup>-3</sup> at the same altitude,

 $\pi$  = a numerical constant, 3.141,592,654 /

σ = average effective collision diameter, taken to be exactly 3.65 x 10-10 m for this MODEL.

This value of  $\sigma$  is an arbitrarily adopted average of several published values.

5.4.2 Computational equation

Eliminating n between equation (27a) and equation (28) yields:

$$L = \frac{R^{8}H^{1}T_{M}}{\sqrt{2}\pi\sigma^{2}NN_{0}P} = 8.050,460,475 \times 10^{-5} \frac{M^{1}T_{M}}{P} . \qquad (28a)$$

5.4.3 Sea-level value and ratio equation

The evaluation of equations (28) and (28a) at sea level

results in:

$$L_o = \frac{1}{\sqrt{2} \pi \sigma^2 n_o} = \frac{R^2 H_o^1 (T_M)_o}{\sqrt{2} \pi \sigma^2 N M_o P_o} = 6.631,722,3 \times 10^{-8} m, \quad (28b)$$

where

L = sea-level value of L,

 $n_0$  = sea-level value of number density, 2.547,552,07 x  $10^{25}$  m<sup>-3</sup>.

/ Basic constant

Equation (28a) divided by the right-hand member of equation (28b) and the use of equation (2hd) leads to the following ratio equation:

$$\frac{L}{I_D} = \frac{M}{M_O} \cdot \frac{(T_M)_O}{T_M} \cdot \frac{P}{P_O} . \tag{28c}$$

A comparison of equations (26c), (27c), and (28c) shows that:

$$\frac{L}{L_0} = \frac{v}{v_0} = \frac{n_0}{n} = \frac{\rho_0}{\rho} \cdot \frac{M}{M_0} = \frac{M}{M_0} \cdot \frac{(T_M)_0}{T_M} \cdot \frac{P}{P_0} . \tag{28d}$$

5.4.4 Validity

Equation (28) for mean free path is based on the concept that temperature and density are uniform throughout a volume equal to the cube of a length many orders of magnitude greater than the mean free path. At 90,000 m<sup>2</sup> the mean free path is 2.5 cm. A length two orders of magnitude greater than L would be 2.5 meters and a cube of this dimension is perhaps approaching the smallest size cube which contains a sufficient number of molecules at this altitude to rigorously apply the derivation of equation (28). Temperatures and densities within this volume may certainly be considered constant. At higher altitudes, however, this may no longer be true for the necessary size cube.

In this MODEL, the value of L from equation (28) becomes 1 meter at 114,000 m. A cube of length two orders of magnitude larger, a 100-meter cube, would have a change in density from top to bottom of about 1%. This amount is considerably more than should be tolerated for the conditions of rigorous validity of the equation for L. At an altitude of 210,000 m', the value of L is 1 kilometer; while at 390,000 m', the value of L is 100 kilometers. Certainly at these altitudes the density is not uniform throughout a sufficiently large cube and the distance through which a molecule will travel between successive collisions depends on its direction of motion. The value of L from equation (28) for a given altitude requires that conditions along the path of the molecule remain equal to those at the particular altitude. At high altitudes this condition can only be met for those molecules moving in a horizontal direction. For molecules moving vertically downward, the distance traveled between collisions will be less than L, because the motion is into a region of exponentially increasing density. For molecules moving vertically upward, the distance traveled between collisions will be greater than L because the motion is into a region of exponentially decreasing density. Some kind of average of these directional mean free path lengths, considering all possible directions, is suggested as a more general concept of mean free path at these altitudes. An unpublished study at GRD shows that the horizontal mean free path, obtained from equation (28), yields values which agree well with this newly suggested mean free path concept to altitudes of about 220,000 m. Above this altitude,

equation (28) should only apply to a horizontal mean free path.

5.5 Collision Freq .ency

5.5.1 Concept and definition

The average velocity of the molecules or atoms within any given volume of air, divided by the mean free path of the molecules within that volume yields the mean collision frequency of the molecules of that volume. That is, any particular molecule in that volume will collide successively with other molecules at a mean rate given by the collision frequency. Analytically collision frequency is defined by

$$\nu = \frac{\overline{V}}{L} , \qquad (29)$$

where

 $\nu$  = the collision frequency in sec<sup>-1</sup>,

 $\overline{V}$  = the average particle velocity in m sec-1, and

L = the mean free path in m.

5.5.2 Computational equation

Equation (21a) for  $\overline{V}$  divided by equation (28a) for L leads

to:

$$\nu = 4\sigma^2 \text{ N.} \left[ \frac{\pi M_0}{R^*} \right]^{\frac{1}{2}} \cdot \frac{P}{M'(T_M)^{\frac{1}{2}}} = 3.358,306,019 \times 10^7 \frac{P}{M'(T_M)^{\frac{1}{2}}} \cdot (29a)$$

5.5.3 Sea-level value and ratio equation

From the evaluation of equations (29) or (29a) at sea level

one obtains:

$$\nu_{\rm o} = \frac{\bar{\rm v}_{\rm o}}{L_{\rm o}} = 4\sigma^2 \ {\rm N} \cdot \left[\frac{\pi \,{\rm M}_{\rm o}}{{\rm R}^*}\right]^{\frac{1}{2}} \cdot \frac{{\rm P}_{\rm o}}{{\rm M}_{\rm o}^{\rm i} \left({\rm T}_{\rm M_{\rm o}}\right)^{\frac{1}{2}}} = 6.920,404,9 \times 10^9 \ {\rm sec}^{-1}, \quad (29b)$$

where

$$\overline{V}_0 = 458.942,034 \text{ m sec}^{-1},$$
 $L_0 = 6.631,722,29 \times 10^{-8} \text{ m}.$ 

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Equations (29), (29a) and (29b) permit the following ratio expressions:

$$\frac{\nu}{\nu_{o}} = \frac{\vec{v}}{\vec{V}_{o}} \cdot \frac{I_{o}}{I} = \frac{P}{P_{o}} \cdot \frac{M_{o}}{M} \cdot \left[\frac{(T_{M})_{o}}{T_{M}}\right]^{\frac{1}{2}}.$$
(29c)

## 5.5.4 Validity

The validity of the value of  $\nu$  is limited principally by the validity of L. Even with the broader concept of L suggested in Section 5.4.4, the value of L should not apply without restrictions above 220 to 250 km. Similarly, values of  $\nu$  must not be used without caution above this altitude.

# 5.6 Temperature (Real Kinetic)

## 5.6.1 Concept and validity

Temperature in this MODEL is a measure of the kinetic energy of the molecules and atoms comprising the atmosphere at any specified altitude. Tabulated values most probably will not indicate the temperature of any body suspended in or passing through the region.

The determination of the value of atmospheric temperature, T, at any given altitude, from conventional measuring techniques requires a knowledge of molecular weight M of the air at that altitude. Without this knowledge of molecular weight, the measurement yields only the value of T/M. Because values of M have not been measured at high altitudes, the so-called temperature measurements from rockets yield only the ratio T/M. This ratio, however, was shown to relate the basic atmospheric properties of pressure, density, specific weight, scale height, particle speed and sound speed. The altitude function of this ratio, T/M, in the form of molecular scale temperature, T<sub>M</sub>, defines the altitude functions of these properties.

With the establishment of the independent assumption regarding the altitude function of molecular weight in Section 5.1, it is now possible to specify values of T with the same degree of reliability as exists in the values of M. These values of T will then permit the determination of the coefficient of viscosity and kinematic viscosity from empirical expressions involving T.

## 5.6.2 Computational equation

The computational equation for real temperature follows directly from the definition of molecular-scale temperature in equation (7).

$$T = T_{M} - \frac{M}{M_{O}} = .03\mu,523,23\mu,1 M \cdot T_{M}$$
, (30)

T temperature (real kinetic, absolute scale) at any specified altitude, and

T<sub>M</sub> = molecular scale temperature (absolute scale) at that altitude.

5.6.3 Sea-level value and ratio equation

Equation (30) evaluated at sea level yields:

$$T_o = (T_M)_o \frac{M_o}{M_o} = (T_M)_o = 288.16^\circ \text{ K (exact)},$$
 (30a)

where

To = sea-level value of T, and

 $(T_M)_0$  = sea-level value of  $T_M$  defined to be 288.16° K (exact).

From the division of equation (30) by (30a), one obtains:

$$\frac{\mathbf{T}}{\mathbf{T}_{o}} = \frac{\mathbf{T}_{M}}{(\mathbf{T}_{M})_{o}} \cdot \frac{\mathbf{M}}{\mathbf{M}_{o}} \tag{30b}$$

5.7 Coefficient of Viscosity

5.7.1 Concept

Viscosity of a fluid (or gas) is a kind of internal friction which resists the relative motion between adjacent regions of a fluid. If two very large parallel plates surrounded by a gas (at normal pressures) are moving relative to each other so that their separation remains constant, experiments show that the layer of gas directly at the surface of each plate is at rest with respect to that plate. It has also been shown that each layer of gas exerts a

<sup>//</sup> Derived constant

drag on the neighboring layers so that there exists a velocity gradient normal to the surface of the plates. If the plates are sufficiently close, the velocity gradient is constant. The relative motion of the plates is resisted by a drag force proportional to the product of the area of the plates times the normal velocity gradient of the fluid. The proportionality factor in this relationship is known as the coefficient of viscosity  $\mu$ . This proportionality factor has been found to vary with the temperature of the gas, but to be independent of the gas pressure within limited ranges of pressure. Various people have contributed to the development of a theoretical expression for  $\mu$  from kinetic theory and Chapman? has recently derived cumbersome formulas which accurately represent the dependence of  $\mu$  on the temperature, at least over the range of 100—1500° K. Because of the complexity of the Chapman equations, however, the values for coefficient of viscosity in this MODEL are computed from the well-known empirical Sutherland's equation, with coefficients as used by the National Bureau of Standards. The plates are sufficients as used by the National Bureau of Standards.

5.7.2 Computational equation

Sutherland's empirical equation for computing viscosity is

$$\mu = \frac{\beta T^{3/2}}{T + S}$$
, (31)

where

5.7.3 Sea-level value and ratio equation

The sea-level value of  $\mu$  is

$$\mu_{o} = \frac{\beta T_{o}^{3/2}}{T_{o} + S} = 1.789,428,53 \times 10^{-5} \text{ kg m}^{-1} \text{ sec}^{-1}$$

$$= 1.789,428,53 \times 10^{-4} \text{ poise,}$$
(31a)

where

 $\mu_o$  = the sea-level value of  $\mu$  ,

 $T_o$  = the sea-level value of T.

≠ Basic constant

Equation (31) divided by equation (31a) yields the ratio equation:

$$\frac{\mu}{\mu_{o}} = \left[\frac{T}{T_{o}}\right]^{3/2} \left[\frac{T_{o} + S}{T + S}\right] . \tag{31b}$$

## 5.7.4 Validity

The users of this MODEL are cautioned that the value of the coefficient of viscosity determined by equation (31) is open to question for conditions of very high and very low values of pressure and density. While equation (31) suggests that the coefficient of viscosity is independent of pressure and depends only on temperature, the measurement of  $\mu$  with an oscillating disk viscometer indicates this situation to be true only within certain limits of pressure, of the order of 2 to .1 atmospheres.

As the pressure decreases below .1 atmosphere, a point is reached where  $\mu$  begins to fall off with further decrease in pressure in a manner which depends upon the size of the viscometer. This change in the dependence of  $\mu$  first occurs when the mean free path of air molecules becomes some small fraction of a linear dimension characteristic of the apparatus or other body. Such a dimension in the case of the viscometer would be the distance between plates.

As the pressure is decreased still further, a point is reached when the mean free path becomes equal to or greater than this characteristic dimension. At this point the viscous stress (drag force per unit area) becomes directly proportional to the quadruple product of density of the gas, velocity of the moving plates or other body, one-fourth the mean speed of the molecules, and a function indicating the reflective properties of the surfaces. This situation characterizes the "free-molecule region" of the gas.

For pressures in between the free-molecule region and the region characterized by viscosity independent of pressure, there exists for any particular viscometer a transition region where the coefficient of viscosity is neither independent of pressure nor directly proportional to it, and the relationship is rather difficult to treat theoretically. Studies indicate, however, that as the dimensions of the viscometer are made larger, both the high and low pressure boundaries of the transition region are moved to smaller values of pressure. Thus by greatly increasing the size and plate separation of the viscometer, the pressure region for which equation (31) yields satisfactory values of  $\mu$  is extended to very low values of pressure.

It may well be that this procedure can be extended until the characteristic dimension becomes so great that appreciable differences in density or temperature exist over a vertical distance equal to this dimension. At this point, equation (31) would begin to become inaccurate regardless of further increase in viscometer size. By dividing atmospheric density by the density gradient at various altitudes, it may be shown that 0.1 per cent variation in density occurs over a vertical distance of 5 to 10 meters at all altitudes below 130 km. Viscometers with plate separations of 10 meters would be expected to yield values of  $\mu$  consistent with equation(31) for pressures as low as those found at 90 kilometers altitude.

Thus values of  $\mu$  tabulated in this MODEL only from-5,000 m to 90,000 m are probably reliable for suitable conditions over this entire range of altitudes, but only when these conditions include body dimensions which are sufficiently large. For altitudes above 40 km, each case ought to be examined with caution before using the tabulated values of  $\mu$ .

## 5.8 Kinematic Viscosity

5.8.1 Definition and computational equation

Kinematic viscosity is defined as the ratio of the coefficient of viscosity of a gas to the density of the gas. Analytically it is expressed as:

$$\eta = \frac{\mu}{\rho} , \qquad (32)$$

where

 $\eta$  = kinematic viscosity of air in  $n^2$  sec<sup>-1</sup>,

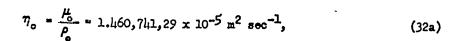
μ = coefficient of viscosity of air in kg sec<sup>-1</sup> m<sup>-1</sup>, and

 $\rho$  = atmospheric density in kg m<sup>-3</sup>.

Because of the empirical nature of the expression for  $\mu$  and since no other atmospheric properties of this MODEL depend upon  $\eta$ , the expression for  $\eta$  has not been transformed to an expression in terms of the three properties, pressure, molecular-scale temperature, and molecular weight. Computations of  $\eta$  have been made directly from equation (32).

5.8.2 Sea-level value and ratio equation

Equation (32) evaluated at sea-level yields:



 $\eta_o$  = sea-level value of  $\eta$ ,

 $\mu_{\rm o}$  = sea-level value of  $\mu$ , 1.789,428,53 x 10<sup>-5</sup> kg m<sup>-1</sup> sec<sup>-1</sup>,  $\mu$ 

From the division of equation (32) by equation (32a) and from equations (7), (18b), and (31b), one obtains:

$$\frac{\eta}{\eta_{o}} = \frac{\mu}{\mu_{o}} \cdot \frac{\rho_{o}}{\rho} = \frac{P}{P_{O}} \cdot \frac{M}{M_{O}} \cdot \left[\frac{T}{T_{O}}\right]^{\frac{1}{2}} \left[\frac{T_{O} + S}{T + S}\right]. \tag{32b}$$

### 5.8.3 Validity

The validity of the tabulated values of  $\eta$  is no better than the validity of either  $\mu$  or  $\rho$ . Within the altitude range of tabulation of  $\eta$ , values of  $\mu$  are the more uncertain and the use of values of  $\eta$  should be subject to the same restrictions applied to the use of  $\mu$ .

// Derived constant

### 5.9 Summary of Ratio Equations

Because of the common relationship of molecular-scale temperature or real temperature and molecular weight to all the properties of this MODEL, the ratio of these properties to their sea-level values are all interrelated in the following multiple equation:

$$\frac{T_{M}}{\left(T_{M}\right)_{O}} \cdot \frac{P_{O}}{P} = \frac{\rho_{O}}{\rho} = \frac{H_{S}}{\left(H_{S}\right)_{O}} \cdot \frac{g}{g_{O}} \cdot \frac{P_{O}}{P} = \left[\frac{C_{S}}{\left(C_{S}\right)_{O}}\right]^{2} \cdot \frac{P_{O}}{P} = \left[\frac{\overline{v}}{\overline{v}_{O}}\right]^{2} \cdot \frac{P_{O}}{P}$$

$$\frac{\omega_o}{\omega} \quad \frac{g}{g_o} \quad = \frac{v}{v_o} \quad \frac{M_o}{M} = \frac{n_o}{n} \quad \frac{M_o}{M} \quad = \frac{L}{L_o} \quad \frac{M_o}{M} = \frac{v_o}{v} \quad \frac{\overline{v}}{\overline{v}_o} \quad \frac{M_o}{M} = \frac{T}{T_o} \quad \cdot \quad \frac{M_o}{M} \quad \cdot \quad \frac{P_o}{P} = \frac{1}{2} \quad \cdot \quad \frac{M_o}{M} \quad \cdot \quad \frac{P_o}{P} = \frac{1}{2} \quad \cdot \quad \frac{M_o}{M} \quad \cdot \quad \frac{P_o}{M} = \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} = \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} = \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} = \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} = \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} = \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} = \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} = \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} = \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} = \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} = \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} = \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} = \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} = \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} = \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} = \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} = \frac{P_o}{M} \quad \cdot \quad \frac{P_o}{M} = \frac{P_o}{M} \quad \cdot \quad \frac{$$

$$\frac{\mu}{\mu} - \frac{\eta_0}{\eta} \quad . \tag{33}$$

## 6. Metric Gravitational System of Units

#### 6.1 Unconventional Form

In this MODEL, as in the ICAO Standard Atmosphere, the system of units employing the dimensions of the Type I gravitational system is not strictly a gravitational system; rather, it is a form of absolute system employing the names of gravitational units, (see Appendix J). In order that there be no confusion between the kilogram force as used in this MODEL and the kilogram force as used in a pure gravitational system of units, the rollowing development is presented.

#### 6.2 Basic Concepts

All properties in this MODEL may be expressed in terms of mass m, length  $\ell$ , time t, and temperature T. The metric absolute system of mechanical units, which has been employed throughout the discussion to this point, uses the kilogram as the unit of mass, the meter as the unit of length, and the second as the unit of time. The unit of acceleration a, therefore, has the dimension of m sec<sup>-2</sup>, while the unit of force F, expressed by Newton's second law as F = ma, has the dimensions of kg m sec<sup>-2</sup> and has been named the "newton."

The metric gravitational system of units is based on the kilogram force kgf, meter, and second. These units through Newton's law imply a unit of mass equal to the unit of force divided by the unit of acceleration, and having the dimensions of kgf sec<sup>2</sup> m<sup>-1</sup>, for which there is no specific, commonly used name. The English counterpart of this unit of mass is the slug or lbf sec<sup>2</sup> ft<sup>-1</sup>.

In its fundamental concept, the kilogram force is the force which gravity exerts on a kilogram mass at the particular altitude and latitude under consideration, and the relationship between the absolute and the gravicational system of units thus depends upon the location. For any fixed latitude, as applied to this MODEL, the variations of gravity with altitude could be used to rigorously relate the kilogram mass and the kilogram force at various altitudes.

### 6.3 Modified Definition of the Kilogram Force

The drafters of the ICAO Standard Atmosphere, on which this MODEL is based, have chosen not to follow the fundamental concept of the gravitational system of units. They have in effect defined the kilogram force as the force which gravity exerts on a kilogram mass at a location where g is equal to go, i.e., at sea level and at 15° 32' 40" latitude. This definition makes the kilogram force an absolute unit, and makes the resulting system of units an absolute system, employing only the dimensions of a gravitational system. The system might therefore be called an absolute-force, gravitational system of units. In equation form, the definition of this absolute kilogram force in terms of the kilogram mass is:

$$1 \text{ kgf} = 9.80665 \text{ m sec}^{-2} \times 1 \text{ kg},$$
 (34)

or conversely,

$$1 \text{ kg} = \frac{1}{9.80665} \text{ kgf sec}^2 \text{ m}^{-1}. \tag{35}$$

The dimensions of the right-hand side of equation (35) are these previously associated with mass in the metric gravitational system. Thus it appears that the metric units of mass in this absolute-force, gravitational system is always exactly \$.80665 times as great as the kilogram mass.

### 6.4 Conversion from Absolute System

Since units of length, time, and temperature are the same in both absolute and gravitational systems of units, only those properties of the MODEL which inherently involve the dimensions of mass have different magnitudes in the two systems. Thus solving equation (35) for unity provides the necessary factor for converting in either direction between the absolute system and the absolute-force gravitational system of units:

$$1 = 9.80665 \text{ m sec}^{-2} \text{ kg kgf}^{-1}(\text{exact}).$$
 (36)

The factor required for converting from the absolute system to the pure gravitational system of units varies according to the geographic location and is expressed by:

$$1 = g kg kgf^{-1}$$
 (36a)

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where g is the acceleration of gravity in  $m \, sec^{-2}$  at the particular altitude and latitude in question.

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## 6.5 Properties Requiring Conversion

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A dimensional analysis of the various properties of this MODEL in terms of mass, length, and time indicates that only pressure, density, specific weight, and coefficient of viscosity involve the dimensions of mass. Hence, only these properties are expressed differently in the two systems of units. For each of these properties the conversion from the metric, absolute system to the metric, absolute-force, gravitational system at any altitude is accomplished by dividing the magnitude and dimensions of the property in the former system by the right-hand side of equation (36), (which is equal to unity).

### 6.6 Converted Sea-Level Values

The sea-level values of atmospheric pressure, density, specific weight, and coefficient of viscosity in units of the metric, absolute-force, gravitational system are obtained by dividing the defined value of  $P_0$  in newtons-2 and the right-hand members of each of equations (18a), (23b), and (31a) respectively by the right-hand side of equation (36). Thus:

$$P_{o} = \frac{101,325. \text{ nt m}^{-2}}{9.80665 \text{ m sec}^{-2} \text{ kg kgf}^{-1}} = 10,332.2745 \text{ kgf m}^{-2},$$
 (37)

$$P_0 = \frac{1.225,013,998 \text{ kg m}^{-3}}{9.80665 \text{ m sec}^{-2} \text{ kg kgf}^{-1}} = .12\text{h},916,663 \text{ kgf sec}^2 \text{ m}^{-4},$$
(38)

$$\omega_{\rm o} = \frac{12.013,283,5 \text{ kg m}^{-2} \text{ sec}^{-2}}{9.80665 \text{ m sec}^{-2} \text{ kg kgf}^{-1}} = 1.225,013,990 \text{ kgf m}^{-3},$$
(39)

$$\mu_{o} = \frac{1.789,428,53 \times 10^{-5} \text{ kg m}^{-1} \text{ sec}^{-1}}{9.80665 \text{ m sec}^{-2} \text{ kg kgr}^{-1}} = 1.824,709,28 \times 10^{-6} \text{ kgf sec m}^{-2}$$
 (40)

#### 6.7 Conversion for All Altitudes

The ratios P/P<sub>o</sub>,  $\rho/\rho_o$ ,  $\omega/\omega_o$ , and  $\mu/\mu_o$  in the absolute system of units, when multiplied by the respective sea-level values given above, yield the values of P,  $\rho$ ,  $\omega$ , and  $\mu$  in the absolute-force, gravitational

system of units.\*

## 7. Preparation of the Metric Tables

## 7.1 Computation of the Tables

The acceleration of gravity, molecular-scale temperature, pressure, and molecular weight are the only properties which were computed directly as functions of H alone, g in terms of a single function for all altitudes, TM and P in terms of ten different functions for ten altitude regions respectively, and M in terms of three different functions for three altitude regions respectively. The remaining properties were computed from expressions in terms of g, TM, P, and M, or in terms of T derived from TM and M. To have computed each of the properties in terms of H alone would have required the development of ten functions for each property, each function applying to a specific altitude region.\*\* Such a procedure would have been unwieldy, and would not have added to the accuracy or validity of the tables. Even the stated computational equations for each of the properties, while serving well for isolated calculations, do not necessarily represent the best approach for development of the tables.

From the multiple equation (33) it is evident that if the ratios of certain basic atmospheric properties to their sea-level values are determined, the remaining ratios are readily computed from products or quotients of not more than two previously determined ratios. The tabulated ratios, when multiplied by the sea-level values of the respective properties in any desired absolute system of units, then yield the required absolute tables.\*\*\*

### 7.2 Detailed Computational Procedure

The following procedure is suggested as one of the better methods for use in any expansion or revision of these tables by desk calculator

<sup>\*</sup> For conversion to the pure gravitational system, these values in the absolute-force, gravitational system of units would have to be multiplied by  $g_0/g$ .

<sup>\*\*</sup> A single function of altitude, closely approximating the densities of this MODEL, particularly above 100 km, was developed by L. Jacchia<sup>33</sup> of the Astrophysical Observatory, Smithsonian Institute and is presented in Appendix L.

<sup>\*\*\*</sup> The tabulation of properties in the absolute-force, gravitational system employed in this MODEL is also made in this manner, although this procedure would not apply to the pure gravitational units.

#### techniques:

- A. List all integral multiples of the desired increment of geometric altitude for which atmospheric properties are to be computed and determine the corresponding values of geopotential altitude to nine significant figures by means of equation (5).
- B. List all integral multiples of the same increment of geopotential altitude for which atmospheric properties are to be computed and determine the corresponding values of geometric altitude to nine significant figures.
- C. Combine the entries of lists compiled in steps A and B into a single list arranged in numerically ascending values of geopotential.
- D. Compute values of  $g/g_0$  to nine significant figures for all tabulated values of H by means of equation (la).
- E. Compute values of  $T_M$  in  ${}^{O}K$  to nine significant figures for all tabulated values of H, using equation (8) and the values of H tabulated in Section 3.1.5 .
- F. Compute values of  $T_{\rm M}/(T_{\rm M})_{\rm o}$  to nine significant figures for all tabulated values of H, using the defined value of  $(T_{\rm H})_{\rm o}$ , 288.16°K.
- G. Compute values of  $\left[T_{\rm H}/\left(T_{\rm H}\right)_{\rm O}\right]^{\frac{1}{2}}$  to nine significant figures for all tabulated values of H.
- $H_{\bullet}$  Compute values of  $P/P_{0}$  to nine significant figures for all tabulated values of H from equations (17a) through (17c), as each applies to its respective altitude range.
- I. Compute value of M to nine significant figures for all tabulated values of H, using equations (24), (24a), and (24b) as each applies to its respective altitude region.
- J. Compute values of  $\rm M/M_{\odot}$  to nine significant figures, using the defined value of  $\rm M_{\odot},\ 28.966$  .
- K. Compute values of T in  ${}^{\circ}$ K to nine significant figures, and  $T/T_0$  for all tabulated values of H above 90,000 m², using equations (30) and (30b), in terms of previously determined quantities. (Below 90,000 m², T =  $T_M$ , and  $T/T_0 = T_M/(T_M)_0$ ; hence T and  $T/T_0$  need not be computed for this altitude region.)
- L. Compute values of  $(T/T_0)^{3/2}$  to nine significant figures for all tabulated values of H up to and including 90,000 m<sup>1</sup> only. For this

altitude region,

$$(T/T_o)^{3/2} = [T_M/(T_M)_o] \cdot [T_M/(T_M)_o]^{\frac{1}{2}}$$
.

- M. Compute values of  $\frac{T_0 + S}{T + S}$  to nine significant figures for all tabulated values of H up to and including 90,000 m<sup>t</sup> only, using S = 110.4 °K from equation (31).
- N. Using the previously established ratios and the following equations, compute to nine significant figures the values of the eleven ratios of atmospheric properties to their respective sea-level values, for all tabulated values of H, except in the case of  $C_g/(C_g)_0$ ,  $\mu/\mu_0$ , and  $\eta/\eta_0$ , which are computed only to 90,000 m\* inclusively:

$$\frac{\rho}{\rho_{o}} = \frac{(T_{H})_{o}}{T_{M}} \cdot \frac{P}{P_{o}} \tag{18b}$$

$$\frac{H_{S}}{(H_{S})_{o}} = \frac{T_{M}}{(T_{M})_{o}} \cdot \frac{g_{o}}{g}$$
 (19c)

$$\frac{\mathbf{c_s}}{\left(\mathbf{c_s}\right)_o} = \left[\frac{\mathbf{T_M}}{\left(\mathbf{T_M}\right)_o}\right]^{\frac{1}{2}} \tag{20c}$$

$$\frac{\vec{v}}{\vec{v}_o} = \left[ \frac{T_M}{(T_M)_o} \right]^{\frac{1}{2}}$$
 (21c)

$$\frac{\omega}{\omega_o} = \frac{\rho}{\rho_o} \cdot \frac{g}{g_o}$$
 (23c)

$$\frac{\mathbf{v}}{\mathbf{v}_0} = \frac{\mathbf{M}}{\mathbf{M}_0} \cdot \frac{\mathbf{g}}{\mathbf{p}} \tag{26c}$$

$$\frac{n}{n_0} = \frac{\rho^{\prime}}{\rho_0} \cdot \frac{M_0}{M} \tag{27e}$$

$$\frac{L}{L_0} = \frac{n_0}{n} \tag{28d}$$

$$\frac{\nu}{\nu_0} = \frac{\bar{v}}{\bar{v}_0} \cdot \frac{I_0}{I} \tag{29c}$$

$$\frac{\mu}{\mu_o} = \left[ \frac{T_o + S}{T + S} \cdot \frac{T}{T_o} \right]^{3/2} \tag{31b}$$

$$\frac{\eta}{\eta_0} = \frac{\rho_0}{\rho} \cdot \frac{\mu}{\mu_0} \tag{32b}$$

O. Compute the mks values of g, P,  $\rho$ , H<sub>g</sub>, C<sub>g</sub>,  $\overline{V}$ ,  $\omega$ ,  $\overline{V}$ , n, L,  $\nu$ ,  $\mu$ , and  $\eta$  to nine significant figures in the mks absolute units by multiplying the tabulated values of g/g<sub>o</sub>, P/P<sub>o</sub> and the tabulated values of each of the eleven ratios listed under step N respectively, by the following corresponding, sea-level values, as they are basically defined or as they are derived by the several equations, sing the mks system of units.

$$g_0 = 9.80665 \text{ m sec}^{-2}$$
, defined (Section 2.1.1)

$$P_0 = 101,325 \text{ newtonsm}^{-2}, \text{ defined (Section 3.2.3)}$$

$$P_0 = .76 \text{ m Hg}$$
, defined (Section 3.2.3)

$$\rho_{\rm o} = 1.225,013,998 \text{ kg m}^{-3}$$
 from equation (18a)

$$(H_s)_0 = 8.434,413,43 m$$
 (196)

$$(C_s)_0 = 340.292,046 \text{ m sec}^{-1}$$

<sup>\*</sup> These properties are listed here only for completeness and are not used in step 0 of the computational procedure since values of TM, M, and T have already been tabulated.

•				
	V <sub>o</sub> = 458.942,035 m sec <sup>-1</sup>	from	equation	(21b)
	$\omega_0 = 12.013,283,5 \text{ kg m}^{-2} \text{ sec}^{-2}$	×	u	<b>(</b> 23b <b>)</b>
π̈́	M <sub>o</sub> = 29.966, defined	Ħ	Ħ	(24)
	v <sub>o</sub> = 23.6կ5,կկկ,1 m <sup>3</sup>	Ħ	Ħ	(26b)
	n <sub>o</sub> = 2.547,552,07 x 10 <sup>25</sup> m <sup>-3</sup>	п	Ħ	(27b)
	I <sub>o</sub> = 6.631,722,3 x 10 <sup>-8</sup> m	Ħ	n	(28b)
	$\nu_{\rm o} = 6.920, \mu_0 \mu_{\rm s}, 9 \times 10^9  {\rm sec}^{-1}$	n	n	(29b)
*	T <sub>o</sub> = 288.16	n	π	(30a)
	$\mu_{\rm o} = 1.789,428,53 \times 10^{-5}  {\rm kg m^{-1}  sec^{-1}}$	.n	18	(3la)
	$\eta_{o} = 1.460,741,29 \times 10^{-5} \text{ m}^{2} \text{ sec}^{-1}$	н	tt	(32a)

P. Compute the values of P,  $\rho$ ,  $\omega$ , and  $\mu$  in the mks, absolute-force, gravitational units. to nine significant figures by dividing the tabulated mks absolute values of these four properties by 9.80665 m sec-2 kg kgf-1 (exact) from equation (36). In principle this procedure is equivalent to multiplying the tabulated values of P/P<sub>0</sub>,  $\rho/\rho_{\rm o}$ ,  $\omega/\omega_{\rm o}$ , and  $\mu/\mu_{\rm o}$  by the following sea-level values in gravitational units:

$$P_o = 10,332.274,5 \text{ kgf m}^{-2},$$
 from equation (37)  
 $\rho_o = .124,916,663 \text{ kgf sec}^2 \text{ m}^{-4},$   $\text{m}$   $\text{m}$  (38)  
 $\omega_o = 1.225,013,998 \text{ kgf m}^{-3},$   $\text{m}$   $\text{m}$  (39)  
 $\mu_o = 1.824,709,28 \times 10^{-6} \text{ kgf sec m}^{-2},$   $\text{m}$   $\text{m}$  (40)

Q. Independently repeat the entire procedure of steps A through P, compare the two results, and account for any discrepancies.

<sup>\*</sup> See footnote on page 56.

<sup>\*\*</sup> The remaining atmospheric properties of this MODEL are numerically and dimensionally equal in both mks systems tabulated.

R. Tabulate the corrected results to any desired number of significant figures less than nine, with values of the ratios always given to one more significant figure than the values of the property itself.

# 7.3 Tabulations Presented

Of the sixteen properties discussed, only one, the mol-volume, is not tabulated for other than sea-level values. In the present edition of the metric tables, the values of pressure, density, specific weight, and coefficient of viscosity are given only in the absolute system of units.

# 7.4 Significant Figures

The number of significant figures to which these tables might be computed is limited only by the capabilities of the machine. The constants, the defining properties, and the functional relationships are all specified as being exact, and thus they do not limit the number of significant figures of the tables. Such a procedure makes for internal consistency to any degree desired. The choice of the number of significant figures tabulated in this MODEL resulted from arbitrary decisions and does not in the slightest amount indicate the validity of the values in depicting the actual atmosphere.

The sca-level values of the various properties are given to eight or nine significant figures depending on whether the first significant figure is greater than or less than 5. Tabulated values of geopotential and geometric altitude are listed to the nearest meter or standard geopotential meter. Tabulated values of g are given in six significant figures\* and values of TM to five significant figures for all altitudes. The values of the remaining properties are given to five significant figures from -5,000 m' to +75,000 m'. Above 75,000 m, the values of these properties are given to only four significant figures. The ratios of the various properties to their respective sea-level values are given to one more significant figure than the corresponding value of the property.

# 7.5 Accuracy of Tabulations

The metric tables were prepared with the aid of desk calculators from the equations developed above. The values of the atmospheric properties discussed in Sections 3 and 4 were computed independently by two people and any discrepancies in results were resolved. Any errors which may appear in the tabulated values of these properties will be due to inaccurate copying. The tables of properties in Section 5 have been computed only once and here some possibility of computational error exists.

A comparison with a more accurate method for computing g indicates that the sixth significant figure is not meaningful for indicating the actual effective gravity above about 40 km.

#### 8. Preparation of the English Tables

#### 8.1 Conversion of Basic Units

The English tables of THE ARDC MODEL ATMOSPHERE are given in terms of the foot (ft), pound (lb), second (sec), and degree Rankine (OR), each of which is defined exactly in terms of the corresponding units employed in the metric tables. The second, of course, is common to both the English and metric systems of measurement. The foot and the pound are defined as follows:

The magnitude of the degree Rankine in terms of the degree Kelvin is derived from the defined relationship of the two temperature scales:

$$T(^{\circ}R) = 1.8 T(^{\circ}K)$$
 (Ref. 60) (43)

where T(OR) is the absolute temperature in the thermodynamic Rankine scale.

From equation (43) one infers that

$$1^{\circ}K = 1.8^{\circ}R \text{ (exact)},$$
 (432)

and from equations (41), (42), and (43a) respectively, one determines the following three conversion factors:

$$1 = 0.3048 \text{ m ft}^{-1} \text{ (exact)}$$
 (412)

$$1 = 0.453,592,3 \text{ kg lb}^{-1} \text{ (exact)}$$
 (42a)

$$1 = 1.8^{\circ} R (^{\circ} K)^{-1} (exact).$$
 (43b)

These three factors are sufficient to convert values of all atmospheric properties in the mks <sup>O</sup>K absolute system of units to the correct values in the fps <sup>O</sup>R absolute system of units.

<sup>\* &</sup>quot;The round value has been accepted by the U.S. National Bureau of Standards and the Commonwealth Standards Laboratory as the common basis on which the American and British representation of the 'foot' should be unified when necessary legal provision is forthcoming." 26-28

<sup>\*\* &</sup>quot;This value is based on an informal understanding between the National Bureau of Standards (Washington, D.C.) and the National Physical Laboratory (Teddington, England) that this rounded quantity would be convenient if the English-speaking nations could arrive at a uniform basis of conversion from the metric to the English system of units." 20-20

#### 8.2 Other Necessary Conversions

#### 8.2.1 English absolute to English gravitational units

As in the metric system of units, the English gravitational system employed in this MODEL is not a pure gravitational system where the unit of force varies with the location in accordance with the value of g. Rather, the unit of force, the pound force (lbf) is taken to be that force which gravity exerts on a pound mass (lb) at a point where g has the standard sea-level value of this MODEL,  $\mathbf{g}_{\mathbf{0}}$ . The definition of the pound force in equation form is

1 lbf = 
$$g_0 \times 1$$
 lb. (L4)

Dividing the defined metric value of  $\mathbf{g}_{o}$  by the conversion factor of equation (41a) yields

$$g_0 = \frac{9.80665}{.3048}$$
 ft  $sec^{-2}$  (45)

Thus,

1 lbf = 
$$\frac{9.80665}{.30h8}$$
 ft sec<sup>-2</sup>lb. (hha)

Since force has the dimension of lbf, and acceleration is in ft  $\sec^{-2}$  by Newton's second law, mass must have the dimensions of lbf  $\sec^2$  ft<sup>-1</sup>. This unit is called the slug. Solving equation (44a) for 1 lbf  $\sec^2$  ft<sup>-1</sup>, one obtains:

1 slug = 1 lbf 
$$\sec^2$$
 ft<sup>-1</sup> =  $\frac{9.80665}{.3048}$  lb. (45)

Thus we find that the slug, the unit of mass in the English (absolute-force) gravitational system of units is exactly 9.80665/.3048 times as large as 1 lb (mass). The factor for converting back and forth between the two Englsih systems of units employed in this MODEL is therefore:

$$1 = \frac{9.80665}{.3018} \text{ ft sec}^{-2} \text{ lb lbf}^{-1}$$
 (46)

or

$$1 = \frac{9.80665}{.3048}$$
 lb slug-1 (46a)

#### 8.2.2 Metric gravitational to English gravitational units

The combining of equations (35), (42a), and (45) yields the following direct relationship between the metric and English gravitational units of mass:

1 slug = 1 (lbf sec<sup>2</sup> ft<sup>-1</sup>) = 
$$\frac{.453,592,3}{.3048}$$
 (kgf sec<sup>2</sup> m<sup>-1</sup>). (47)

Dividing the two right-hand members of equation (47) respectively by the corresponding parts of equation (41a) yields

$$1 \text{ lbf} = .453,592,3 \text{ kgf}.$$
 (48)

This equation provides the factor for converting directly between the two gravitational systems of this MODEL:

$$1 = .453,592,3 \text{ kgf lbf}^{-1}. \tag{49}$$

8.2.3 Rankine-to-Fahrenheit scale and Kelvin-to Fahrenheit scale conversions

The relationship of the thermodynamic Fahrenheit temperature scale to the thermodynamic Rankine scale is established by the following definition:

$$t (^{\circ}F) - t_{i}(^{\circ}F) = T (^{\circ}R) - T_{i}(^{\circ}R),$$
 (50)

where t<sub>i</sub>(°F) is defined to be 32°F (exact), the ice-point temperature.

Using the definition of  $T_i$  in  ${}^{O}K$  (see Section 3.1.4) and equation (43), one obtains

$$T_i(^{\circ}R) = 1.8 \times 273.16 = 491.688^{\circ}R.$$
 (51)

Introducing equations (43) and (51) into equation (50) yields

$$t (^{\circ}F) = 1.8 (T^{\circ}K - 273.16) + 32.$$
 (52)

8.2.4 Standard geopotential meter to standard geopotential foot

From equation (41) it follows directly that

1 std. geopotential foot (ft:) = 0.3048 x 1 std. geopotential meter m:. (53)

Thus the factor for converting m' to ft' and vice versa becomes:

$$1 = 0.3048 \text{ m}! \text{ ft}!^{-1}(\text{exact}).$$
 (53a)

/ Basic constant

#### 8.2.5 Geometric meter to nautical mile

The defined conversion\* from meters to the international nautical mile (i n mi) in this MODEL is:

$$1 (i n mi) = 1,852 meters (exact).$$
 (54)

The conversion factor is therefore:

$$1 = 1,852 \text{ m (i n mi)}^{-1}$$
 (54a)

8.3 Sea-Level Values of Atmospheric Properties in English Units

By means of equation (43a) for  $T_M$  or by the proper application of equations (41a), (42a), and (43b) to the mks, absolute, sea-level values of the various other atmospheric properties listed under computational procedure, step 0 of Section 7.2, the following sea-level values in English absolute units\*\* are derived. The English absolute values of  $P_0$ ,  $\rho_0$ ,  $\omega_0$ , and  $\mu_0$ , when divided by the conversion factor given in equation (46) yield the sea-level values of these properties in the English (absolute-force) gravitational system.\*\*\*

$$g_0 = 32.174,048,55 \text{ ft sec}^{-2}, \text{ from equation (45a)}$$

$$(T_M)_0 = 1.8(286.16^{\circ}\text{K}) = 518.688^{\circ}\text{R}$$
(55)

$$P_0 = \frac{101,325 \times .3048}{.453,592,3} = 68,087.266,9 lb ft-1 sec-2 or poundals ft-2 (56)$$

$$P_0 = \frac{101,325 \times (.3048)^2}{.453,592,3 \times 9.80665} = 2,116.216,95 \text{ lbf ft}^{-2}$$
 (56a)

or

$$P_0 = \frac{.76 \times 12}{.3048} = 29.921,259,84 \text{ in Hg}$$
 (56b)

<sup>\*</sup> United States Department of Defense Directive 2045.1, 17 June 1954, directed the adoption of the international nautical mile (equal to 1852 meters) as a standard value with the Department of Defense effective 1 July 1954.

<sup>\*\*</sup> See Appendix J.

<sup>\*\*\*</sup> All remaining properties are numerically and dimensionally the same in both systems.

$$\rho_{\rm o} = \frac{1.225,013,998 \times (.3048)^3}{.453,592,3} = .076,475,137,4 \text{ 1b ft}^{-3}$$
 (57)

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$$\rho_{o} = \frac{1.225,013,998 \times (.3048)^{4}}{.453,592,3 \times 9.80665} = 2.376,919,99 \times 10^{-3} \text{ lbf sec}^2 \text{ ft}^{-4}$$
 or slugs ft-3 (57a)

$$(H_B)_0 = \frac{8,434.113,43}{3018} = 2.767,196,007 \times 10^{14} \text{ ft}$$
 (58)

$$(C_s)_o = \frac{340.292,046}{.3048} = 1.116,443,720 \times 10^3 \text{ ft sec}^{-1}$$
 (59)

$$\bar{v}_0 = \frac{458.912,035}{.3048} = 1.505,715,337 \times 10^3 \text{ ft sec}^{-1}$$
 (60)

$$\omega_{o} = \frac{12.013,283,5 \times (.3018)^{2}}{.1453,592,3} = 2.1460,5114,77 \text{ lb ft}^{-2} \text{ sec}^{-2}$$
 (61)

$$\omega_{o} = \frac{12.013,283,5 \times (.3048)^{3}}{.453.592.3 \times 9.80665} = 7.647,513,72 \times 10^{-2} \text{ lbf ft}^{-3}$$
 (61a)

$$M_o = 28.966$$
 (nondimensional) (unchanged) (62)

$$v_0 = \frac{23.6 \text{h}5, \text{h}\text{h}, 08}{(.3048)^3} = 835.030,977 \text{ ft}^3$$
 (63)

$$n_0 = 2.547,552,07 \times (.3048)^3 \times 10^{25} = 7.213,864,115 \times 10^{23} \text{ ft}^{-3}$$
 (64)

$$L_0 = \frac{6.631,722,29 \times 10^{-8}}{.3048} = 2.175,761,906 \times 10^{-7} \text{ ft}$$
 (65)

$$\nu_{\rm o} = 6.920, \mu_{\rm o}, 91 \times 10^9 \text{ sec}^{-1} \text{ (unchanged)}$$
 (66)

$$\mu_{\circ} = \frac{1.789, 428, 53 \times .3048 \times 10^{-5}}{.453, 592, 3} = 1.202, 440, 640 \times 10^{-5} \text{lb ft}^{-1} \text{sec}^{-1}$$
 (67)

$$\mu_{\rm o} = \frac{1.789,428,53 \times (.3048)^2 \times 10^{-5}}{.453,592,3 \times 9.80665} = 3.737,299.76 \times 10^{-7} \text{ lbf sec ft}^{-2}$$
 (67a)

$$\eta_{o} = \frac{1.160,711.29 \times 10^{-5}}{(.3018)^{2}} = 1.572,328,83 \times 10^{-1} \text{ ft}^{2} \text{ sec}^{-1}$$
 (68)

It is to be noted that only three exactly defined numerical constants Here employed in all the above conversions. Hence the English values may be reliably carried to any number of significant figures consistent with the metric absolute values

# 8.4 Calculation of the English Tables

### 8.4.1 Functions employed

This MODEL ATMOSPHERE is defined exactly in terms of various gradients of molecular-scale temperature in <sup>O</sup>K m:-l between specific exact values of altitude expressed in m', and in terms of constants defined exactly in metric units. These definitions cannot be converted exactly to English units. Thus it is preferable to compute English tables from exactly the same equations used for the metric tables, after first making the necessary conversion of the English altitudes to metric altitudes, and then obtaining the English values of the various properties by another conversion.

### 8.4.2 Altitude increments

The argument of the English tables, similar to the metric tables, is given in consecutive integral multiples of a fixed altitude increment in both geometric feet and standard geopotential feet, i.e.,

### $n \times 2500$ ft and $n \times 2500$ ft!

where n = -6, -5, -1, -3, -2, -1, 0, +1,2,3 etc. to 21. From -15,000 ft! to 60,000 ft! the increment is 2500 ft or ft!; from 60,000 ft! to 300,000 ft!, the increment is 10,000 ft or ft!; from 300,000 ft! to 500,000 ft!, the increment is 25,000 ft or ft!; from 500,000 ft! to 1,000,000 ft!, the increment is 50,000 ft or ft!; and from 1,000,000 ft! to 1,700,000 ft!, the increment is 100,000 ft or ft!.

### 8.4.3 Altitude conversions

In order to use identically the same equations for converting between geopotential and geometric altitude for the English tables as was used in the metric tables, these conversions must be made in metric units. Thus, to convert the tabulated integral multiple values of ft to m', multiply the altitudes in ft by exactly .3048 m ft<sup>-1</sup>, from equation (41a), to obtain the equivalent in meters, and then convert the results to m' by using equation (5). This value of m' is then converted to the equivalent in ft' by dividing by exactly .3048 m' ft'-1

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from equation (532). Starting with tabulated, integral, multiple values of ft, the conversion to m' is directly by means of equation (532). This value of m' is then converted to m by means of equation (6), and the corresponding value of ft is then obtained by means of equation (412). Since the conversion factors cited and the constants of equations (5) and (6) are all defined to be exact, the conversions may be carried to any desired number of significant figures.

# 8.4.4 Computational procedure

Having arranged in sequence the values of m' for each English altitude to be tabulated, the computation of the tables proceeds exactly as indicated in Section 7.2, steps D through N, but stopping short of O.

Compute the values of TM and T in OC to nine significant figures from the Kelvin values by means of equation (9). Compute the values of TM and T in OF to nine significant figures from the Kelvin values by using equation (50).

Compute the values of the remaining properties in English units from the multiplication of the ratios of the various properties determined in step N by their respective sea-level values in the desired English absolute and absolute-force units.

### 8.4.5 Tabulated values

In this edition of the MODEL, only half of the properties discussed are contained in the English tables. The properties tabulated are those designated by g, P,  $\rho$ , C<sub>0</sub>, M, T,  $\mu$ , and  $\eta$ . It should noted that  $\rho$  and  $\mu$  are given only in Type I, absolute-force, gravitational units, while P is given not only in this system (lbf ft-2) but also in mb and in inches of Hg. Temperatures in the English tables are given in  ${}^{\circ}$ C,  ${}^{\circ}$ F, and  ${}^{\circ}$ R.

These tables were prepared from a single computation using desk calculators; as the values have not been checked by independent calculations, some chance of error exists.

Above 60,000 ft the altitude increments of the English tables are considerably larger than the increments of the metric tables.

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METRIC TABLE I

# TEMPERATURES AND MOLECULAR WEIGHT AS FUNCTIONS OF GEOMETRIC AND GEOPOTENTIAL AF THUDE

ALTITUDE	м	OLECULA	TEMPERA	TURE REAL K	INETIC	MOLECUL	AR WEIGHT
Z,m H,		M, °K	$\tau_{ m M}/\tau_{ m Mo}$	T,°K	T/To	м	м/м
-5,000 -5,0 -4,966.1 -5,0 -4,000 -4,0 -3,997.5 -4,0 -3,000 -3,0 -2,998.6 -3,0 -2,000 -2,0 -1,999.4 -2,0 -1,000 -1,0	003.9 000 2 002.5 000 3 001.4 000 3 000.6	520.69 520.66 514.18 514.16 507.67 507.66 501.16 501.16	1.11287 1.11278 1.09028 1.09023 1.06770 1.06767 1.04513 1.04511			28.966	1.00000
- 999.8 -1,0 0 1,000 1,000.2 1,1 2,000 1,2,000 2,3,001.4 3,4,000 3,4,002.5 4, 5,000 4,5,003.9 5,6,000 6,005.7 7,000 6,7007.7 8,000 7,8,000 8,010.7 8,000 8,010.7 8,000 8,010.7 8,000 8,010.8 9,012.8 9	0 999.8 000 999.4 000 998.6 000 997.5 000 ,996.1 ,000 ,994.3 ,000 ,989.9 ,000 ,989.9 ,000	294.66 288.16 281.66 281.66 275.16 268.67 268.66 262.16 255.69 249.16 242.71 242.66 236.23 236.16	.84210 .81978 .81954 .79726 .79698	same as T <sub>M</sub> for altitudes up to 90 km	same as $T_{ m M}/T_{ m Mo}$ for altitudes up to 90 km.	constant at 28.966 for altitudes up to 90 km'	1,00000 for altitudes up to 90 km'
10,016 10 11,000 10 11,019 11 12,000 11 12,023 12 13,000 12 13,027 13 14,000 13	,984.3 ,000 ,981 ,000 ,977 2,000 2,973 3,000 5,979	223.26 223.16 216.78 216.66 216.66 216.66 216.66 216.66	77443 75229 75187 75187 75187 75187 75187 75187	51 DO 74 74 74 74 74		28 <b>.</b> 96	6 1.00000 

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	<del></del> !		TEMPERAT	URE		MOLECUL	AR WEIGHT
ALTITUI	DE	MOLECULA	R SCALE	REAL K	INETIC		
Z,m I	H,m'	T <sub>M</sub> , °K	$T_{M}/T_{MO}$	т,°к	T/To	М	м/м <sub>о</sub>
15,035 1 16,000 1 16,040 1 17,006 1 17,046 1 18,000 1 18,051 1 19,000 1 19,057 1 20,000 2 20,063 2 21,070 22,000 22,076	29,8 30,0 30,8 31,0 31,8 32,0 32,0 32,0	216.66 21	.751874 .751874 .751874 .751874 .751874 .751874 .751874 .751876 .75186 .75186 .75186 .75186 .75186 .75	150 00 du sabutitus 20, 15, 16, 16, 16, 16, 16, 16, 16, 16, 16, 16	E TW/TMO fo	8	

ALTI	TUDE	MOLECUL	TEMPFRA AR SCALE	TURE REAL KINETIC		MOLECUL	AR WEIGHT
Z,m	H,m'	т <sub>м</sub> , •к	$T_{\rm M}/T_{\rm Mo}$	т,°к	T/To	М	м/м <sub>о</sub>
35,000 35,194 36,000 36,205 37,000 37,217 38,000 38,229 39,000 39,241	34,808 35,000 35,797 36,000 36,786 37,000 37,774 38,000 38,762 39,000	246.09 246.66 249.05 249.66 252.02 252.66 254.98 255.66 257.95 258.66	.853988 .855983 .864283 .866394 .874575 .376805 .384865 .887215 .895151 .897626			28.966	1.00000
40,000 40,253 41,000 41,266 42,000 42,279 43,000 43,293 44,000 44,307	39,750 40,000 40,737 41,000 41,724 42,000 42,711 43,000 43,698 44,684	260.91 261.66 263.87 264.66 266.83 267.66 269.79 270.66 272.75 273.66	.905433 .908937 .915713 .918448 .925989 .928859 .936262 .939270 .946532 .949681	altitudes up to 90 km'	$T_{ m M}/T_{ m MO}$ for altitudes up to 90 km,	for altitudes up to 90 km'	itudes up to 90 km'
45,321 46,000 46,335 47,000 47,350 48,000 48,365 49,000 49,381	45,000 45,670 46,000 46,655 47,000 47,640 48,000 48,625 49,000	276.66 278.67 279.66 281.63 282.66 282.66 282.66 282.66	.960092 .967062 .970503 .977322 .980913 .980913 .980913 .980913	same as T <sub>M</sub> for a	same as T <sub>M</sub> /T <sub>Mo</sub> for	constant at 28.966 f	1.00000 for altitudes
50,000 50,396 51,000 51,412 52,000 52,429 53,000 53,446 54,000 54,463	49,610 50,000 50,594 51,578 52,000 52,562 53,000 53,545 54,000	282.66 282.66 282.66 282.66 282.66 282.66 282.66 282.66 282.66 280.53 278.76	.980913 .980913 .980913 .980913 .980913 .980913 .980913 .973535 .967379			. 28.966	1.00000

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ALTITUDE		MOLECUI	TEMPERA LAR SCALE		INETIC	MOLECUI	MOLECULAR WEIGHT	
Z,m	H,m'	T <sub>M</sub> , °K	T <sub>M</sub> /T <sub>Mo</sub>	T,°K	T/To	М	M/M <sub>o</sub>	
55,000 55,480 56,000 56,498 57,000 57,516 58,534 59,553 60,000 61,591 62,000 61,591 62,611 63,000 64,651 65,600 66,692 67,714 68,000 67,714 68,000 67,714 68,000 67,714 68,000 69,757 70,000 70,779 71,802 72,805 73,000	54,528 55,000 55,511 56,000 56,493 57,000 57,476 58,000 58,457 59,000 61,401 62,000 61,401 62,382 63,362 64,000 64,342 65,322 66,000 67,280 66,301 67,000 67,280 68,259 69,000 69,238 70,000 71,194 72,000 72,171	276.70 274.86 272.96 269.06 269.06 269.26 26	.960230 .953845 .946929 .940311 .933633 .926777 .920340 .913243 .907052 .899709 .893767 .886174 .880487 .872640 .867211 .859106 .853939 .845572 .832038 .827408 .814148 .804969 .800893 .791435 .7876901 .774395 .764367 .774395 .764367 .734678 .737299 .734678 .723765 .721448	same as $T_M$ for altitudes up to 90 km.	same as $T_{M}/T_{M_{O}}$ for altitudes up to 90 km'	constant at 28.966 for altitudes up to 90 km'	1.00000 for altitudes up to 90 km'	
73,848 74,000 74,872	73,000 73,148 74,000	204.66 204.08 200.76	.710230 .708221 .696696			28.966	1.00000	

ALTI	TUDE	MOLECUL	TEMPER	ATURE REAL KI	NETIC	MOLECUL	AR WEIGHT
Z,m	H,m'	T <sub>M</sub> , °K	T <sub>M</sub> /T <sub>Mo</sub>	T,°K	T/To	М	м/м <sub>o</sub>
75,000 75,895 76,000 76,920 77,000 77,944 78,000 78,969 79,000 81,000 81,000 81,020 82,045 85,000 83,072 84,000 84,098 85,000 85,125 86,000 86,152 87,000 87,179 88,000 89,235 90,000 90,264 91,000	74,125 75,000 75,102 76,000 77,005 77,005 77,000 77,055 78,000 79,006 79,981 80,900 81,930 82,904 83,000 84,000 85,000 85,000 85,000 86,798 87,771 88,000 88,744 89,000 89,716	200.27 196.86	.694999 .683162	same as T <sub>M</sub> for altitudes up to 90 km'	same as $T_{M}/T_{Mo}$ for altitudes up to 90 km'	constant at 28.966 for altitudes up to 90 km' 60.000 constant at 28.966 for altitudes	1.00000 for altitudes up to 90 km'
91,293 92,000 92,322 93,000 93,351 94,000 94,381	90,000 90,688 91,000 91,659 92,000 92,630 93,000	196.86 199.27 200.36 202.67 203.86 206.07 207.36	.683162 .691526 .695308 .703325 .707454 .715109 .719600	196.9 197.0 197.1 197.5 197.7 198.3 198.6	.68316 .68355 .68395 .68523 .68609 .68799 .68929	28.63 28.49 28.22 28.09 27.87	1.00000 .98848 .98367 .97429 .96980 .96208 .95787

			TEMPER	ATURE		MOTECHI	AR WEIGHT
ALTI	rude	MOLECUL	AR SCALE	RFAL K	INETIC	NOLECOL	
Z,m	H,m'	T <sub>M</sub> ,°K	$T_{M}/T_{MO}$	т, "К	T/T <sub>o</sub>	М	M/M <sub>o</sub>
95,000 95,411 96,000 96,441 97,000 97,472 98,000 98,503 99,000 99,534	93,601 94,000 94,572 95,500 95,542 96,500 96,512 97,000 97,482 98,000	209.46 210.86 212.86 214.36 216.26 217.86 219.65 221.36	.726902 .731746 .738691 .743892 .750477 .756038 .762258 .768184 .774037	204.4	.71215	26.74 26.65	.95147 .94751 .94217 .93842 .93794 .93038 .92661 .92322 .92004 .91680
100,000 100,566 101,000 101,598 102,000 102,631 103,000 103,663 104,000	98,451 99,000 99,420 100,000 100,38 101,000 101,35 102,00	226.44 228.36 229.83 231.86 233.22 235.36 236.61 238.86 240.07	.792476 .797582 .804622 .809349 .816766 .82111 .82891	208.0 208.9 210.0 210.0 212.0 212.0 214.0 214.0	72196 72481 72881 73155 73592 73852 74329	26.39 26.32 26.24 26.18 26.10 26.05 26.05 25.93 8 25.86	.91102 .90876 .90578 .90587 .90101 5 .89941 7 .89665 3 .89531 5 .89265
105,000 105,73 106,000 106,76 107,000 107,79 108,000 108,83	104,000 104,264 105,000 105,228 106,000 106,19 107,000 107,100	245.86 1 246.76 249.3 29 250.1 252.8 26 253.5 26 256.3 26 256.9	6 .85320 8 .85638 6 .86535 6 .86813 6 .87749 5 .8798 6 .8896 6 .8896	7 218. 22 219. 35 220. 31 221. 99 223. 76 223. 45 225.	6 .7584 1 .7605 8 .7663 3 .7681 1 .7743 6 .7759 .5 .7821	8 25.17 25.6 25.6 25.6 25.6 25.5 25.5 25.5 25.4 25.4	5 .88897 2 .88806 5 .88557 3 .88483 6 .88241 4 .88182 8 .87948 .6 .87903
110,00 110,90 111,00 111,91 112,00 112,91 113,00 114,0	00 108,1 02 109,0 00 109,0 37 110,0 00 110,0 73 111,0 00 111,0	29 260.3 00 263.3 95 263.0 00 266.0 061 267.0 000 270.0 026 270.0 092 273.0	9033 96 •9139 96 •9150 86 •9260 97 •9268 36 •9382 45 •9383	56 228 37 230 91 230 83 232 122 232 129 235 149 235 173 237	.1 .791 .2 .798 .5 .799 .7 .807 .8 .807 .1 .815 .2 .816	97 25. 76 25. 38 25. 89 25. 86 25. 09 25.	32 .87420 32 .87397 25 .87182 25 .87168 19 .86958 19 .86952 13 .86749

	<del></del>			<del></del>			
ALTI	rude	MOLECULA	TEMPERA R SCALE	TURE REAL K	INETIC	MOLECULA	R WEIGHT
Z,1n	H,m'	T <sub>M</sub> , °K	$T_{ m M}/T_{ m Mo}$	т,°к	T/T <sub>o</sub>	м	м/м <sub>о</sub>
115,000 115,045 116,000 116,082 117,000 117,119	112,957 113,000 113,921 114,000 114,885 115,000	277.21 277.36 280.58 280.86 283.96 284.36	.961993 .962521 .973709 .974667 .985422 .986813	239.9 240.1 242.4 242.6 244.8 245.1	.83268 .83618 .84106 .84175 .84949	25.07 25.07 25.02 25.02 24.97 24.96	.86558 .86549 .86377 .86362 .86206
118,000 118,156 119,000 119,194	115,850 116,000 116,813 117,000	287.33 287.86 290.71 291.36	.997131 .998959 1.00884 1.01110	247.2 247.6 249.7 250.2	.85796 .85929 .86648 .86814	24.92 24.92 24.87 24.87	.86043 .86019 .85889 .85860
120,000 120,232 121,000 121,270 122,000 122,309 123,000 123,348 124,000	117,777 118,000 118,740 119,000 119,703 120,000 120,665 121,000	294.08 294.86 297.45 298.36 300.82 301.86 304.19 305.36 307.56 508.86	1.02054 1.02325 1.03224 1.03540 1.04393 1.04754 1.05562 1.05969 1.06731 1.07184	252.2 252.7 254.6 255.3 257.1 257.9 259.6 260.5 262.1 263.1	.87504 .87703 .88363 .88596 .89226 .89493 .90091 .90393 .90960	24.80 24.79 24.76 24.75 24.72 24.71 24.69	.85743 .85710 .85604 .85567 .85471 .85431 .85344 .85302 .85223 .85178
124,387 125,000 125,427 126,000 126,467 127,000 127,507 128,000 128,548 129,000	126,000 126,434	310.92 312.36 314.29 315.86 317.65 319.36 321.02 322.86 327.20	1.07899 1.08398 1.09067 1.09613 1.10235 1.10827 1.11402 1.12042 1.12549 1.15512	264.6 265.7 267.1 268.3 269.7 270.9 272.2 273.6 277.1 281.7	.91831 .92204 .92704 .93113 .93580 .94025 .94458 .94940 .96168	24.65 24.64 24.62 24.61 24.59 24.57 24.56 24.54 24.53	.85108 .85060 .84997 .84947 .84892 .84840 .84736 .84693 .84637
130,000 130,630 131,000 131,672 132,000 132,774 135,000	128,000 128,355 129,000 129,315 130,000	342.86 346.41 352.86 356.01 362.86 3584.79	1.16882 1.18983 1.20214 1.22453 1.23545 1.25923 1.33532 1.43275	284.9 289.9 292.7 298.0 300.6 306.1 323.9 346.7	.98881 1.0059 1.0159 1.0341 1.0450 1.0623 1.1241 1.2031	24.51 24.49 24.48 24.46 24.45 24.45 24.38 24.32	

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ALTI	TUDE	TEMPERATURE				MOLECULA	MOLECULAR WEIGHT	
AUTI			AR SCALE		KINETIC		<del></del>	
Z,m	H,m'	T <sub>M</sub> , *K	$T_{M}/T_{Mo}$	T,°K	T/T <sub>o</sub>	М	M/M <sub>o</sub>	
210.000	376 007	1.70 60	1 50157	362.7	1.2588	24.28	.83835	
140,000 143,153	136,983 140,000	432.69 462.86	1.50157	387.2	1.3435	24.23	.83644	
145,000	141,766	480.52	1.66755	401.4	1.3931	24.20	.83541	
148,385	145,000	. 512.86	1.77978	427.6	1.4837	24.15	.85566	
150,000	146,542	528.28	1.83329	440.0	1.5269	24.13	.83288	
153,625	150,000	562.86	1.95329	467.9	1.6237	24.08	.83127	
155,000	151,311	575.97	1.99877	478.5	1.6604	24.06	.85069	
158,874	155,000	612.86	2.12680	508.2	1.7635	24.02	.82919	
160,000	156,072	623 58	2.16399	516.8	1.7935	24.01	.82878	
164,131	160,000	662.06	2.30032	548.4	1.9032	23.97	.82737	
165,000	160,826	671.12	2.32897	555.1	1.9263	23.96	.82709	
169,397	165,000	712.86	2.47383	588.6	2.0428	23.92	.82575	
170,000	165,572	718.58	2.49369	593.2	2.0587	23.91	.82558	
174.671	170,000	762.86	2.64735	628.8	2.1823	23.88	.82432	
175,000	170,311	765.97	2.65815	631.3	2.1909	23.87	.82424	
179,954	175,000	812.86	2.82086	669.0	2.3217	23.84	.82303	
180,000	175,043	813.11	2.82174	669.1	2.3220	25.84	.82290	
185,000	179,768	840.52	2.91684	679.7	2.3588	23.42	.80869	
185,245	180,000	841.86	2.92150	680.2	2.5606	23.41	.80802	
190,000	184,486	867.88	3.01179	690.4	2.3960	23.04	•79555	
190,545	185,000	870.86	3.02214	691.6	2.4001	23.00	.79418	
195,000	189,196	895.20	3.10660	701.3	2.4336	22.69	•78337	
195,854	190,000	899.86	3.12278	703.1	2.4401	22.63	.78138	
200,000	193,899	922.48	3.20127	712.2	2.4715	22.36	.77204	
201,171	195,000	928.86	3.22342	714.8	2.4804	22.29	.76951	
205,000	198,595	949.71	3.29579	723.2	2.5097	22.06	.76149	
206,497	200,000	957.86	3.32406	726.5	2.5212	21.97	-75846	
210,000	203,284	976.91	3.39016	734.3	2.5481	21.77	.75162	
211,831	205,000	986.86	3.42469	738.3	2.5622	21.67	.74816	
215,000	207,966	1004.1	3.48440	745.4	2.5867	21.50	.74238	
217,175	210,000	1015.9	3.52533	750.3	2.6036	21.39	.73854	
220,000	212,641	1031.2	3.57849	756.6		21.25	•73371	
222,526	215,000	1044.9	3.62597	762.3		21.13	.72953	
225,000	217,308	1058.2	3.67243	767.8		21.02	•72555 TOXAGE	
227,887	220,000	1073.9	3.72661	774.3		20.89	.72106	
230,000	221,969	1085.3	3.76624	779.1		20.79	.71787	
233,256	225,000	1102.9	3.82725	786.5	_		.71310	
235,000	226,622	1112.3	3.85990		_ : :	11	.71062 .70561	
235,634	230,000	1131.9	3.92789	798.6	2.7715			

ALTI	rude	MOTECIE	TEMPERA AR SCALE	TURE REAL	KINETIC	MOLECULA	MOLECULAR WEIGHT	
Z,m	H,m'	T <sub>M</sub> , °K	T <sub>M</sub> /T <sub>Mo</sub>	T,°K	T/To	М	м/м <sub>o</sub>	
240,000 244,021 245,000 249,417 250,000 254,821 255,000	231,268 235,000 235,908 240,000 240,540 245,000 245,165	1139.2 1160.9 1166.1 1189.9 1193.0 1218.9 1219.8	5.95342 4.02853 4.04680 4.12916 4.14003 4.22980 4.23313	801.7 810.9 813.1 823.2 824.5 835.5 835.9	2.7823 2.8140 2.8218 2.8567 2.8613 2.8995 2.9010	20.39 20.23 20.20 20.04 20.02 19.86 19.85	.70377 .69853 .69729 .69184 .69114 .68550	
260,000 260,235 265,000 265,657 270,000 271,088 275,000 276,528	249,784 250,000 254,395 255,000 258,999 260,000 263,597 265,000	1246.6 1247.9 1273.4 1276.9 1300.0 1305.9 1326.7 1334.9	4.32608 4.33044 4.41890 4.43108 4.51157 4.53172 4.60411 4.63236	847.4 847.9 858.8 860.3 870.3 872.8 881.8 885.3	2.9425 2.9804 2.9856 3.0202 3.0289 3.0600 3.0722	19.69 19.68 19.54 19.52 19.39 19.36 19.25	.67975 .67950 .67447 .67379 .66943 .66837 .66463	
280,000 281,977 285,000 287,435 290,000 292,902 295,000 298,377	268,187 270,000 272,771 275,000 277,347 280,000 281,917 285,000	1353.3 1363.9 1379.9 1392.9 1406.5 1421.9 1433.0 1450.9	4.69650 4.73300 4.78876 4.83363 4.88088 4.93427 4.97286 5.03491	893.3 897.8 904.8 910.4 916.3 922.9 927.8 935.5	3.0999 3.1157 3.1398 3.1592 3.1797 3.2029 3.2196	19.12 19.07 18.99 18.93 18.87 18.80 18.75 18.68	.66005 .65829 .65566 .65359 .65146 .64911 .64744 .64482	
300,000 303,862 305,000 309,356 310,000 314,859	300,000	1508.9 1512.3 1537.9		939.3 948.2 950.8 960.8 962.3	3.2905 3.2995 3.3344 3.3395 3.3783		.64359 .64072 .63989 .63679 .63634 .63302	
320,000 325,893 330,000 336,963 340,000 348,069	310,000 313,714 320,000 322,738	1595.9 1617.4 1653.9 1669.7	5.53810 5.61286 5.73938 5.79449	985.3 988.9 1008 1024 1031 1050		18.13 18.06 17.94 17.89	.62593 .62593 .62344 .61938 .61767 .61331	

ALTI	TUDE	MOLECIT	TEMPER		KINETIC	MOLECULA	R WEIGHT
Z,m	H,m'	T <sub>M</sub> , *K	T <sub>M</sub> /T <sub>Mo</sub>	т, °к	T/T <sub>o</sub>	М	и/N <sub>o</sub>
350,000 359,213 360,000	331,735 340,000 340,705	1721.9 1769.9 1773.9	5.97558 6.14194 6.15613	1,054 1,075 1,077	3.6588 3.7322 3.7385 3.8181	17.74 17.60 17.59 17.45	.61230 .60767 .60728 .60259
370,000 370,394 380,000 381,612 390,000 392,867	349,648 350,000 358,565 360,000 367,456 370,000	1825.8 1827.9 1877.5 1885.9 1929.1 1943.9	6.33614 6.34321 6.51561 6.54449 6.69456 6.74577	1,100 1,101 1,123 1,127 1,146 1,153	3.8212 3.8975 3.9103 3.9768 3.9996	17.45 17.33 17.31 17.21 17.21	.60241 .59818 .59750 .59404 .59290
400,000 404,160 410,000 415,491 420,000	376, 320 380,000 585,158 390,000 393,970	2059.9	6.87297 6.94704 7.05086 7.14832 7.22823	1,169 1,178 1,192 1,204 1,214	4.0560 4.0889 4.1351 4.1784 4.2140	16.99 16.93 16.89	.59014 .58859 .58647 .58454 .58299
426,860 430,000 438,267 440,000 449,713	400,000 402,756 410,000 411,516 420,000	2117.9 2133.8 2175.9 2184.7	7.34960 7.40507 7.55087 7.581 <i>3</i> 9 7.75215	1,230 1,237 1,256 1,260 1,282	4.2680 4.2928 4.3577 4.3713 4.4475	16.79 16.72 16.70	.58072 .57971 .57712 .57659 .57372
450,000 460,000 461,197 470,000 472,721 480,000 484,283 490,000 495,884	420,250 428,959 430,000 437,642 440,000 450,000 454,932 460,000	2295.8 2291.9 2336.2 2349.9 2386.4 2407.9 2436.5	7.93247 7.95343 8.10725 8.15471 8.28151 8.35598 8.45526	1,307 1,327 1,333 1,350 1,359	4.5280 4.5374 4.6061 4.6273 4.6840 4.7173	16.53 16.52 16.46 3 16.44 0 16.38 3 16.35 3 16.31	.57363 .57082 .57050 .56815 .56744 .56560 .56455 .56317 .56179
500,000 507,525 510,000 519,205 520,000 530,000 530,925 540,000 542,686	470,000 472,12 480,00 480,67 489,21 490,00 497,71	2523.9 2 2536.2 2 2581.9 9 2585.8 2 2635.3 0 2639.9 9 2684.	9 8.7585h 2 8.80125 9 8.95981 3 8.97346 3 9.14522 9 9.16109 6 9.31646	1,417 1,437 1,437 1,438 2 1,462 1,463 1,463	4.8976 7 4.916 7 4.987 9 4.993 1 5.078 5 5.148	6 16.20 7 16.18 7 16.12 9 16.12 9 16.06 0 16.06 9 16.01	.56085 .55918 .55864 .55668 .55651 .55448 .55430 .55266

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METRIC TABLE II

PRESSURE, DENSITY AND ACCELERATION OF GRAVITY AS FUNCTIONS OF GEOMETRIC AND GEOPOTENTIAL ALTITUDE

ALTI	muz.	FRESSURE		DENS	SITY	ACCELER OF GRA	RATION VITY
Z,m	H,m'	P,mb	P/P <sub>o</sub>	$\rho$ , kg/m <sup>3</sup>	P/P <sub>o</sub>	g,m/sec <sup>2</sup>	g/go
-5,000 -4,996.1 -4,000 -3,997.5 -3,000 -2,998.6 -2,000	-5,003.9 -5,000 -4,002.5 -4,000 -3,001.4 -3,000 -2,000.6	1.7776 <sup>+3</sup> 1.7769 1.5960 1.5956 1.4297 1.4295 1.2778	1.75438 1.75365 1.57515 1.57469 1.41104 1.41082 1.26112 1.26103	1.9312 1.9305 1.7698 1.7694 1.6189 1.6187 1.4782	1.57644 1.57591 1.44472 1.44437 1.32157 1.32140 1.20667 1.20660	9.82210 9.82209 9.81901 9.81900 9.81592 9.81591 9.81283 9.81282	1.001575 1.001574 1.001260 1.001259 1.000945 1.000944 1.000630
-1,999.4 -1,000 - 998.8	-2,000 -1,000.2 -1,000	`1.2777 1.1393 1.1393	1.12441	1.3470	1.09960 1.09958	9.80774 9.80774	1.000315
0 1,000 1,000.2 2,000 2,000.6 3,000 3,001.4 4,000 4,002.5	0 999.8 1,000 1,999.4 2,000 2,998.6 3,000 3,997.5 4,000	1.01325 <sup>+3</sup> 8.9876 <sup>+2</sup> 8.9875 7.9501 7.9495 7.0121 7.0108 6.1660 6.1640	1.00000 8.87008-1 8.86994 7.84615 7.84556 6.92039 6.91917 6.08537 6.08339	1.2250 1.1117 1.1117 1.0066 1.0065 9.0926-1 9.0913 8.1935 8.1935	1.00000 9.07475 9.07464 8.21671 8.21622 7.42243 7.42137 6.68847 6.68671	9.80665 9.80356 9.80356 9.80048 9.80048 9.79740 9.79740 9.79432 9.79431	1.00000 .9996854 .9996854 .9993710 .9993708 .9990568 .9990563 .9987427
5,000 5,003.9 6,000 6,005.7 7,000 7,007.7 8,000 8,010.1 9,000 9,012.8	4,996.1 5,000 5,994.3 6,000 6,992.3 7,000 7,989.9 8,000 8,987.3 9,000	5.4048+2 5.4020 4.7217 4.7181 4.1105 4.1060 3.5651 3.5599 3.0800 3.0742	5.33413 <sup>-1</sup> 5.33133 4.65998 4.65635 4.05676 4.05233 3.51851 3.51339 3.03977 3.03401	7.3643 <sup>-1</sup> 7.3612 6.6011 6.5969 5.9002 5.8950 5.2578 5.2516 4.6706	6.01161 6.00906 5.38859 5.38519 4.81643 4.81216 4.29206 4.28701 3.81270 3.80685	9.79124 9.79123 9.78816 9.78505 9.78506 9.78506 9.78201 9.7894 9.77894	.9984287 .9984275 .9981149 .9981131 .9978013 .9977968 .9974877 .9974846 .9971744
10,000 10,016 11,000 11,019 12,000 12,023 13,000 13,027 14,000 14,031	9,984.3 10,000 10,981 11,000 11,977 12,000 12,973 13,000 13,969 14,000	2.6500 <sup>+2</sup> 2.6436 2.2700 2.2632 1.9399 1.9330 1.6579 1.6510 1.4170 1.4102	2.61532 2.60903 2.24030 2.23358 1.91455 1.90774 1.63626 1.62943 1.39849	4.1351 <sup>-1</sup> 4.1270 3.6480 3.6391 3.1193 3.1082 2.1659 2.6548 2.2785 2.2675	3.37554 3.36896 2.97792 2.97069 2.54637 2.53731 2.17624 2.16716 1.86001 1.85100	9.77582 9.77280 9.77274 9.76973 9.76966 9.76666 9.76658 9.76360	

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AL/I	ITUDE	PRE	SSURE	DE	NSITY	ACCELE OF GR	RATION :
Z,m	H,m'	P,mb	P/P <sub>o</sub>	$\rho$ , kg/m <sup>3</sup>	P/Po	g,m/sec <sup>2</sup>	g/go
15,000 15,035 16,000 16,040 17,000 17,046 18,000 18,051 19,000	15,000 15,960 16,000 16,955 17,000	1.2112 <sup>+2</sup> 1.20 <sup>1</sup> ; 1.0353 1.0287 8.8496+1 8.7866 7.5652 7.5048 6.4674 6.4099	1.19533 <sup>-1</sup> 1.18869 1.02173 1.01528 8.73388-2 8.67167 7.46623 7.40662 6.38285 6.32611	1.9367 1.6647 1.6542 1.4230 1.4129 1.2165 1.2067 1.0399	1.58097 1.35891 1.35033 1.16162 1.15334 9.93016-2 9.85088 8.48925	9.76042 9.75747 9.75735 9.75441 9.75427 9.75135 9.75119 9.74829	•9952973 •9952862 •9949849 •9949723 •9946728 •9946585 •9943607 •9943448 •9940448
20,000 20,063 21,000 21,070 22,000 22,076 23,000 23,084 24,000 24,091	19,937 20,000 20,931 21,000 21,924 22,000 22,917 23,000 23,910 24,000	5.5293 <sup>+1</sup> 5.4748 4.7275 4.6761 4.0420 3.9940 3.4562 3.4113 2.9554 2.9137	5.45694-2 5.40323 4.66564 4.61498 3.98918 3.94173 3.41101 3.36670 2.91677 2.87555	8.8909-2 8.8034 7.6016 7.5191 6.4995 6.4222 5.5575 5.4853 4.7522 4.6851	8.41379 7.25779-2 7.18634 6.20534 6.13797 5.30565 5.24255 4.53667 4.47774 3.87934 3.82451	9.74811 9.74523 9.74504 9.74218 9.74196 9.73912 9.73889 9.73607 9.73581 9.73302 9.73274	.9940311 .9937371 .9937174 .9934255 .9934038 .9931140 .9930902 .9928027 .9928027 .9924916 .9924633
25,000 25,099 26,000 26,107 27,000 27,115 28,000 28,124 29,000 29,133	24,902 25,000 25,894 26,000 26,886 27,000 27,877 28,000 28,868 29,000	2.5273 <sup>+1</sup> 2.4886 2.1632 2.1278 1.8555 1.8233 1.5949 1.5655 1.3737 1.3469	2.49428-2 2.45606 2.13493 2.10001 1.83126 1.79943 1.57407 1.54504 1.35573 1.32930	4.0639 <sup>-2</sup> 4.0016 3.4359 3.3748 2.9077 2.8528 2.4663 2.4169 2.0966 2.0521	3.31742-2 3.26658 2.80476 2.75490 2.37361 2.32877 2.01332 1.97296 1.71147 1.67520	9.72997 9.72967 9.72692 9.72659 9.72387 9.72352 9.72083 9.72045 9.71778	.9921805 .9921498 .9918697 .9918365 .9915589 .9915232 .9912484 .9912099 .9909379 .9908967
30,000 30,142 31,000 31,152 32,000 32,162 33,000 33,172 34,000 34,183	29,859 30,000 30,850 31,000 31,840 32,000 32,830 33,000 33,819 34,000	1.1855 <sup>+1</sup> 1.1611 1.0251 1.0028 8.8801 <sup>+0</sup> 8.6776 7.7068 7.5224 6.7006 6.5327	1.17002 <sup>-2</sup> 1.14592 1.01167 9.89735 <sup>-3</sup> 8.76402 8.56423 7.60604 7.42412 6.61300 6.44726	1.7861 <sup>-2</sup> 1.7461 1.5248 1.4889 1.3044 1.2721 1.1160 1.0890 9.6019 <sup>-3</sup> 9.3404	1.45803-2 1.42540 1.24472 1.21538 1.06478 1.03840 9.12666-3 8.88944 7.83821 7.62473	9.71431 9.71170 9.71124 9.70866 9.70816 9.70562 9.70510 9.70258	.9906276 .9905835 .9903175 .9902704 .9900075 .9899573 .9896977 .9896443 .9893879

		<u> </u>		n		r	
ALT	ITUDE	PRE	SURE	DEI	SITY	ACCELE OF GR	
Z,m	H,m'	P,mb	P/P <sub>o</sub>	p,kg/m <sup>3</sup>	P/Po	g,m/sec <sup>2</sup>	g/go
35,000 35,194 36,000 36,205 37,000 57,217 38,000 38,229 39,000 59,241	34,898 35,000 35,797 36,000 36,786 37,000 37,774 38,000 38,762 39,000	5.8359+0 5.6829 5.0914 4.9519 4.4493 4.3221 3.8944 3.7785 5.4142 3.3084	5.75960 <sup>-3</sup> 5.60855 5.02486 4.88717 4.39115 4.26562 3.84344 3.72908 3.36952 3.26514	8.2619 <sup>-3</sup> 8.0265 7.1221 6.9101 6.1507 5.9597 5.3209 5.1489 4.6112 4.4560	6.74437-3 6.55217 5.81390 5.64082 5.02089 4.86497 4.34354 4.20313 3.76419	9.69955 9.69896 9.69551 9.69589 9.69348 9.69282 9.69045 9.68975 9.68742 9.68569	.9890784 .9890184 .9637690 .9887056 .9884597 .9883927 .9881506 .9830800 .9878416
40,000 40,253 41,000 41,266 42,000 42,279 43,000 43,293 44,000 44,307	39,750 40,000 40,737 41,000 41,724 42,000 42,711 43,000 43,698 44,000	2.9977 <sup>+0</sup> 2.9013 2.6361 2.5481 2.3215 2.2411 2.0474 1.9739 1.8082 1.7411	2.95851 <sup>-3</sup> 2.86333 2.60159 2.51474 2.29110 2.21176 2.02060 1.94812 1.78454 1.71828	4.0027 <sup>-3</sup> 3.8629 3.4803 3.3541 3.0310 2.9169 2.6438 2.5408 2.3096 2.2165	3.26751 <sup>-3</sup> 3.15332 2.84105 2.73803 2.4 <sup>-1</sup> 422 2.38115 2.15815 2.07408 1.88534 1.80933	9.68439 9.68362 9.68136 9.68056 9.67834 9.67749 9.67531 9.67443 9.67229	.9875328 .9874546 .9872241 .9871420 .9869155 .9868294 .9866072 .9865169 .9862044
45,000 45,321 46,000 46,335 47,000 47,350 48,365 49,000 49,381	44,684 45,000 45,670 46,655 47,000 47,640 48,000 48,625 49,000	1.5991 <sup>+0</sup> 1.5378 1.4161 1.3600 1.2558 1.2044 1.1147 1.0673 9.8961 <sup>-1</sup> 9.4578	1.57820 <sup>-3</sup> 1.51765 1.39763 1.34224 1.25936 1.18866 1.10014 1.05333 9.76671-4 9.33411	2.0206 <sup>-3</sup> 1.9364 1.7704 1.6942 1.5535 1.4845 1.3739 1.3155 1.2197 1.1657	1.64946 <sup>-3</sup> 1.58073 1.44523 1.383C <sup>4</sup> 1.26812 1.21179 1.12155 1.07383 9.95675 <sup>-4</sup> 9.5157 <sup>4</sup>	9.66927 9.66830 9.66625 9.66523 9.66323 9.66217 9.66021 9.65911 9.65719 9.65605	.9859908 .9858920 .9856828 .9855796 .9853750 .9852673 .9850673 .9849550 .9847598
50,000 50,396 51,000 51,412 52,000 52,429 53,000 53,446 54,000 54,463	49,610 50,000 50,594 51,000 51,578 52,000 52,562 53,000 53,545 54,000	8.7058-1 8.3810 7.8003 7.4269 6.9256 6.5813 6.1493 5.8320 5.4588 5.1637	8.67088-4 8.27142 7.69829 7.32973 6.83507 6.49524 6.06886 5.75576 5.38738 5 09615	1.0829 <sup>-3</sup> 1.0330 9.6140 <sup>-4</sup> 9.1537 8.5360 8.1116 7.5791 7.1881 6.7790 6.4534	8.83961 <sup>-14</sup> 8.43237 7.84809 7.47235 6.96807 6.62162 6.18694 5.86775 5.53383 5.26800	9.65418 9.65299 9.65117 9.64992 9.64815 9.64686 9.64515 9.64380 9.64214 9.64074	.9844524 .9843306 .9841452 .9840185 .9838381 .9837064 .9835311 .9833944 .98352243

		1	<del></del>			ACCETE	RATION
ALT	ITUDE	PRE	SSURE	DE	NSITY		AVITY
Z,m	H,m'	P,mb	P/Po	ρ,kg/m <sup>3</sup>		g,m/sec <sup>2</sup>	g/go
55,000 55,430 56,000 56,498 57,516 58,000 58,534 59,000 59,553	55,000 55,511 56,000 56,493 57,000	4.8388-1 4.5641 4.2822 4.0270 3.7833 3.5467 3.3367 3.1179 2.9375	4.77557-4 4.50445 4.22624 3.97438 3.73384 3.50036 3.29306 3.07713 2.89912 2.69987	6.0924-4 5.7850 5.4674 5.1777 4.8991 4.6268 4.3832 4.1276 3.9154 3.6761	4.97335-4 4.72241 4.46310 4.22666 3.99926 3.77692 3.57809 3.36945 3.19620 3.00083	9.63769 9.63612 9.63463 9.63312 9.63157 9.63012 9.62851 9.62711	.9829176 .9827705 .9826111 .9824586 .9823047 .9821467 .9819985 .9818350 .9816924
60,000 60,572 61,000 61,591 62,000 62,611 63,000 64,651 64,000 65,672 66,000 66,692	59,439 60,000 60,420 61,401 62,000 62,382 63,000 63,362 64,000 64,342 65,000 65,322 66,000	2.581¼-1 2.3955 2.26¼1 2.093¼ 1.9820 1.8255 1.7315 1.588¼ 1.5096 1.3790 1.3132-1 1.19¼5 1.1399 1.0322	2.54761-4 2.36417 2.23453 2.06598 1.95606 1.80159 1.70885 1.56764 1.48982 1.36099 1.29606-4 1.17885 1.12503 1.01866	3.4918-4 3.2681 3.1089 2.9002 2.7631 2.5689 2.4514 2.2711 2.1709 2.0038 1.9189-4 1.7643 1.6928 1.5502	2.85042 <sup>-14</sup> 2.66784 2.53783 2.36750 2.25558 2.09705 2.00114 1.85394 1.77218 1.63573 1.56640 <sup>-14</sup> 1.44026 1.38185 1.26546	9.62545 9.62411 9.62240 9.62111 9.61934 9.61812 9.61523 9.61323 9.61213 9.61018 9.60913 9.60913 9.60614 9.60407	.9815232 .9813864 .9810806 .9808999 .9807749 .9805884 .9804694 .9802768 .9801640 .9799653 .9798588 .9796539 .9795537
67,000 67,714 68,000 68,735 69,000 69,757	66,301 67,000 67,280 68,000 68,259 69,000	9.8726-2 8.8969 8.5301 7.6491 7.3523 6.5590	9.74349 <sup>-5</sup> 8.78052 8.41856 7.54912 7.25615 6.47323	1.4903 1.3591 1.3093 1.1888 1.1479 1.0374	1.20546 1.21658 1.16544 1.06883 9.70447-5 9.37010 8.46874	9.60315 9.60102 9.60016 9.59796 9.59717	.9793425 .9792488 .9790312 .9789439 .9787199 .9786395 .9784087
70,000 70,779 71,000 71,802 72,000 72,825 73,000 73,848 74,000 74,872	69,238 70,000 70,216 71,000 71,194 72,000 72,171 73,000 73,148 74,000	6.3212-2 5.6088 5.4206 4.7826 4.6357 4.0662 3.9535 3.4464 3.3619 2.9118	6.23854 <sup>-5</sup> 5.53547 5.34974 4.72010 4.57513 4.01299 3.90181 3.40132 3.31794 2.87373	1.0040-4 9.0313-5 8.7624 7.8424 7.6286 6.7922 6.6252 5.8666 5.7391 5.0529	8.19618 <sup>-5</sup> 7.37244 7.15289 6.40188 6.22739 5.54461 5.40830 4.78903 4.68489 4.12479	9.59186 9.59120 9.58881 9.58822 9.58576 9.58524 9.58271 9.58225	.9783347 .9780975 .9780304 .9777864 .9777261 .9774753 .9774220 .9771642 .9771181

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#### METRIC TABLE II CONTINUED

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ALTI	TUDE	PRE	SSURE	DE	nsity		ERATION RAVITY
Z,m	H,m'	P,mb	P/Po	ρ,kg/m <sup>3</sup>	PIPo	g,m/sec <sup>2</sup>	g/go
75,000 75,895 76,000 76,920 77,000 77,944 78,000 78,969 79,000	74,125 75,000 75,102 76,000 76,078 77,000 77,0.5 78,000 78,030	2.8503 <sup>-2</sup> 2.452 2.409 2.061 2.034 1.733 1.717 1.457	2,4200 2.3775 2.0344 2.0069 1.7103 1.6942 1.4378 1.4303	4.9582 <sup>-5</sup> 4.339 4.263 3.648 3.599 3.067 3.058 2.578 2.565	4.04747 3.5423 3.4801 2.9780 2.9377 2.5035 2.4799 2.1046 2.0936	5 9.57928 9.57661 9.57630 9.57356 9.57332 9.57051 9.57035 9.56746 9.56737	.9768142 .9765423 .9765106 .9762314 .9762070 .9759206 .9759036 .9756098
79,994 80,000 81,000 81,020 82,000 82,045 83,000 83,072 84,000 84,098	79,000 79,006 79,981 80,000 80,956 81,000 81,930 82,000 82,904 83,000	1.225 1.224 <sup>-2</sup> 1.033 1.030 8.723 <sup>-3</sup> 8.656 7.365 7.277 6.220 6.117	1.2087 1.2075 <sup>-5</sup> 1.0195 1.0162 8.6085 <sup>-6</sup> 8.5425 7.2690 7.1815 6.1383 6.0373	2.167 2:165 <sup>-5</sup> 1.828 1.822 1.544 1.532 1.303 1.288 1.101 1.083	1.7695 1.7676 <sup>-5</sup> 1.4924 1.4874 1.2601 1.2504 1.0640 1.0512 8.9851-6 8.8373	9.56143 9.56137 9.55846 9.55832 9.55549 9.55528	.9752990 .9752973 .9749943 .9749883 .9746915 .9746777 .9743888 .9743671 .9740862 .9740566
85,000 85,125 86,000 86,152 87,000 87,179 88,000 88,207 89,000 89,235	83,878 84,000 84,652 85,000 85,825 86,000 86,798 87,000 87,771 88,000	5.252-3 5.143 4.436 4.323 3.746 3.635 3.164 3.055 2.673 2.569	5.1837-6 5.0754 4.3778 4.2668 3.6974 3.5870 3.1229 3.0155 2.6378 2.5351	9.295-6 9.101 7.850 7.651 6.630 6.432 5.600 5.407 4.730 4.546	7.5878-6 7.4293 6.4081 6.2456 5.4121 5.2506 4.5712 4.4140 3.8611 3.7108	9.54956 9.54919 9.54659 9.54614 9.54363 9.54310 9.54067 9.54006 9.53771 9.53701	.9737838 .9737461 .9734816 .9734356 .9731795 .9731253 .9728774 .9728149 .9725046
90,000 90,264 91,000 91,293 92,000 92,322 93,000 93,351 94,000 94,381	88,744 89,000 89,716 90,000 90,688 91,000 91,659 92,000 92,630 93,000	2.258-3 2.159 1.907 1.815 1.612 1.528 1.367 1.367 1.162 1.162	2.2282-6 2.1312 1.8823 1.7916 1.5913 1.5085 1.3490 1.3490 1.1468 1.0789	3.995-6 3.822 3.375 3.213 2.819 2.658 2.350 2.206 1.965 1.837	3.2615-6 3.1196 2.7552 2.6225 2.3012 2.1695 1.9806 1.6057 1.4992	9.53475 9.53397 9.53179 9.53093 9.52884 9.52789 9.52588 9.52485 9.52293 9.52180	.9722739 .9721944 .9719724 .9718842 .9716709 .9715740 .9713697 .9712639 .9710685 .9709539
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		PRESS	VIRE.	DENS	SITY	ACCELER OF GRA	
Z,m	H,m'	P,mb	P/P <sub>O</sub>	$\rho$ ,kg/m <sup>3</sup>	P/P <sub>o</sub>	g,m/sec <sup>2</sup>	g/go
95,000 95,411 96,000 96,441 97,000 97,472 98,000 98,503 99,000	93,601 94,000 94,572 95,000 95,542 96,000 96,512 97,000 97,482 98,000	9.905 <sup>-4</sup> 9.264 8.466 7.905 7.254 6.749 6.231 5.777 5.365 4.957	9.7759 <sup>-7</sup> 9.1622 8.3552 7.8021 7.1590 6.6612 6.1492 5.7015 5.2914 4.8920	1.647 <sup>-6</sup> 1.534 1.386 1.285 1.169 1.079 9.882 <sup>-7</sup> 9.092 8.379 7.680	1.3449-6 1.2521 1.1311 1.0488 9.5392-7 8.8106 8.0671 7.4220 6.8399 6.2691	9.51998 9.51876 9.51703 9.51573 9.51408 9.51269 9.51113 9.50965 9.50818 9.50661	.9707675 .9706439 .9704666 .9703339 .9700240 .9698654 .9697142 .9695649 .9694044
100,000 100,566 101,000 101,598 102,000 102,631 103,000 103,663 104,000 104,696	98,451 99,000 99,420 100,000 100,389 101,000 101,358 102,000 102,326 103,000	4.629-4 4.263 4.004 3.675 3.471 3.175 3.015 2.749 2.624 2.385	4.5689-7 4.2073 3.9516 3.6268 3.4253 3.1333 2.9753 2.7129 2.5896 2.3538	7.123-7 6.504 6.069 5.522 5.184 4.699 4.439 4.009 3.809 3.428	5.8142 <sup>-7</sup> 5.3091 4.9545 4.5074 4.2321 3.8363 3.6234 3.2728 3.1092 2.7986	9.50524 9.50357 9.50230 9.50053 9.49935 9.49750 9.49641 9.49347 9.49347	.9692646 .9690946 .9689644 .9687849 .9686644 .9684753 .9683645 .9681657 .9680648
105,000 105,730 106,000 106,764 107,000 107,798 108,000 108,832 109,000	107,162	1.576 1.535 1.378 1.349	2.2585-7 2.0463 1.9735 1.7826 1.7277 1.5559 1.5153 1.3605 1.3314 1.1918	3.276-7 2.938 2.825 2.523 2.438 2.172 2.110 1.873 1.829 1.619	2.6739-7 2.3984 2.3044 2.0600 1.9901 1.7731 1.7222 1.5292 1.4932 1.3216	7 9.49053 9.48839 9.48760 9.48536 9.48232 9.48173 9.47929 9.47880 9.47626	.9677652 .9675466 .9674657 .9672372 .9671664 .9669278 .9668672 .9666184 .9665682 .9663091
110,000 110,902 111,000 111,937 112,000 112,973 113,000 114,000	109,000 109,095 110,000 110,061 111,000 111,026	1.060 1.047 9.316-5 9.244 8.203 8.176 2 7.243	1.0459 1.0331	1.402 1.383	1.1444 1.1289 9.9280 9.8432 8.6292 8.5975	9.47322 9.47293	.9659705 .9656906 .9656718 .9653815 .9653733

				Density		ACCELERATION OF GRAVITY	
	TUDE	PRES	P/Po	$\rho$ , kg/m <sup>3</sup>	P/P <sub>o</sub>	g,m/sec <sup>2</sup>	g/go
Z,m 115,000 115,045 116,000 116,082 117,119 118,000 118,156 119,000 119,194	112,957 113,000 113,921 114,000 114,885 115,000 116,813 117,000	6.426-5 6.392 5.710 5.655 5.081 5.011 4.528 4.447 4.040 3.952	6.3422-8 6.3083 5.6354 5.5815 5.5815 5.0145 4.9459 4.4684 4.3893 3.9873 3.9008	8.076 <sup>-8</sup> 8.029 7.090 7.015 6.234 6.140 5.490 5.383 4.842 4.726	6.5928-8 6.5540 5.7875 5.7265 5.0887 5.0120 4.4812 4.3938 3.9524 3.8579	9.46123 9.46110 9.45830 9.45807 9.45538 9.45503 9.45246 9.45201 9.44954 9.44898	.9647768 .9647633 .9644787 .9644543 .9641807 .9641454 .9638829 .9638365 .9635276
120,000 120,232 121,000 121,270 122,000 122,309 123,000 123,348 124,000 124,387	117,777 118,000 118,740 119,000 119,703 120,000 120,665 121,000 121,627	3.610 <sup>-5</sup> 3.518 3.230 3.135 2.894 2.798 2.595 2.500 2.531	3.5628-8 3.4716 3.1876 3.0939 2.8557 2.7610 2.5615 2.4671 2.3005 2.2074	4.277-8 4.156 3.783 3.660 3.351 3.229 2.973 2.852 2.640 2.523	3.4911-8 3.3927 3.0881 2.9881 2.7355 2.6357 2.4265 2.3282 2.1554 2.0595	9.44663 9.44595 9.44371 9.44292 9.44079 9.43788 9.43687 9.43584	.9632877 .9632188 .9629904 .9629100 .9626931 .9626013 .9623960 .9623987 .9620990 .9619840
125,000 125,427 126,000 126,467 127,000 127,507 128,000 128,548 129,000	122,589 123,000 123,551 124,000 124,512 125,000 125,473 126,000	2.096-5 2.004 1.887 1.797 1.701 1.614 1.534 1.451 1.386	2.0685-8 1.9775 1.8622 1.7737 1.6783 1.5928 1.5143 1.4320 1.3681 1.2903	2.348-8 2.235 2.092 1.982 1.865 1.761 1.665 1.566 1.476	1.9171-8 1.8243 1.7073 1.6181 1.5225 1.4372 1.3593 1.2781 1.2049 1.1170 •	9.43206 9.43081 9.42915 9.42624 9.42476 9.42333 9.42174 9.42043 9.41872	.9618022 .9616755 .9615055 .9613670 .9612089 .9610585 .9609125 .9609125 .9606162 .9604417
130,000 130,630 131,670 131,670 132,000 132,71 135,000 137,92	127,395 128,000 128,355 2 129,000 129,315 4 130,000 0 132,193	1.182 1.141 1.071 1.039 1.039 1.039 1.039 1.039	1.1662 1.1259 1.0571 1.0255	1.201 1.147 1.058 1.017	9.8014" 9.3657 8.6327 8.3007	3 9.41752 9 9.41569 9.41462 9.41267 9.41172 9.40965 9.40302 9.39454	.9603200 .9601334 .9600240 .9598251 .9597281 .9595169 .9588413

ALTI	UDE .	PRES	SURE	DENS	STTY	ACCELER OF GRA	
Z,m	H,m'	P,mb	P/P <sub>o</sub>	$\rho$ ,kg/m <sup>3</sup>	PIPo	g,m/sec <sup>2</sup>	g/go
143,153 145,000 148,385 150,000 153,625 155,000 158,874 160,000 164,131 165,000 169,397 170,000	135,983 140,000 141,766 145,000 146,542 150,000 151,311 155,000 156,072 160,000 160,826 165,000 165,572 170,000	5.336-6 4.238 3.729 2.985 2.698 2.173 2.008 1.624 1.531-6 1.531-6 1.191 9.692-7 9.431 7.689	5.2662 <sup>-9</sup> 4.1831 3.6807 2.9464 2.6628 2.1442 1.9820 1.6032 1.5110 <sup>-9</sup> 1.2264 1.1756 9.5657 <sup>-10</sup> 9.3080 7.5881	4.296 <sup>-9</sup> 3.190 2.704 2.028 1.779 1.345 1.215 9.234 <sup>-10</sup> 8.554 <sup>-10</sup> 6.531 6.183 4.737 4.573 3.511	3.5071 <sup>-9</sup> 2.6042 2.2072 1.6555 1.4525 1.0977 9.9162 <sup>-10</sup> 7.5380 6.9824 <sup>-10</sup> 5.3312 5.0476 3.8668 3.7326 2.8663	9.38855 9.37945 9.37945 9.36437 9.35972 9.34930 9.34535 9.33424 9.33101 9.31671 9.30244 9.30244 9.28914	.9573660 .9564375 .9558941 .9548996 .9544256 .9533630 .9529605 .9518276 .9518276 .9518935 .9502935 .9502935 .9487606 .9485852 .9472289
175,000	170,311	7.582 6.189	7.4832 6.1085	3.1119 2.653	2.8152 2.1655	9.28621 9.27413	.9471335 .9456985
185,000 185,245 190,000 190,545 195,000	175,043 179,768 180,000 184,486 185,000 189,196 190,000	6.178-7 5.082 5.035 4.208 4.124 3.506 3.400	6.0974 <sup>-10</sup> 5.0159 4.9689 4.1533 4.0704 3.4602 3.3559	2.647 <sup>-10</sup> 2.107 2.083 1.689 1.650 1.364 1.316	2.1609 <sup>-10</sup> 1.7196 1.7008 1.3790 1.3469 1.1138 1.0747	9.27400 9.25983 9.25913 9.24569 9.24415 9.23159 9.22918	.9456852 .9442401 .9441693 .9427984 .9426413 .9413599 .9411146
201,171 205,000 206,497 210,000	193,899 195,000 198,595 200,000 203,284 205,000 207,966 210,000	2.938-7 2.821 2.475 2.354 2.096 1.974 1.783 1.665	2.8995-10 2.7840 2.4427 2.3229 2.0685 1.9486 1.7600 1.6430	1.110 <sup>-10</sup> 1.058 9.079 <sup>-11</sup> 8.560 7.474 6.970 6.188 5.709	9.0572 <sup>-11</sup> 8.6369 7.4117 6.9880 6.1014 5.6899 5.0511 4.6605	9.21751 9.21422 9.20347 9.19927 9.18946 9.18433 9.17548 9.16941	.9359247 .9355891 .9384929 .9380648 .9370643 .9365418 .9356389 .9350200
220,000 222,526 225,000 227,887 230,000 233,256 235,000 238,634	212,641 215,000 217,308 220,000 221,969 225,000 226,622 230,000	1.524-7 1.410 1.308 1.200 1.128 1.026 9.759-8 8.805	1.5044-10 1.3920 1.2914 1.1846 1.1131 1.0126 9.6315-11 8.6900	4.703 4.308 3.894 3.620 3.241	4.2039-11 3.8389 3.5164 3.1789 2.9554 2.6457 2.4953 2.2124	9.16154 9.15450 9.14762 9.13960 9.13374 9.12472 9.11989 9.10984	.9342168 .9334995 .9327979 .9319802 .9313823 .9304621 .9299699 .9289453

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### METRIC TABLE II CONTINUED

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Δ1.77	TUDE	PRE	SURE	DENSITY		ACCELERATION OF GRAVITY	
· Z,m	H,m'	P,mb	P/P <sub>o</sub>	ρ,kg/m <sup>3</sup>	P/Po	g,m/sec <sup>2</sup>	g/go
240,000 244,021 245,000 249,417 250,000 254,821 255,000	231,268 235,000 235,908 240,000 240,540 245,000 245,165	8.475-8 7.586 7.387 6.560 6.459 5.692 5.666	8.3646 <sup>-11</sup> 7.4869 7.2900 6.4741 6.3746 5.6179 5.5920	2.592 <sup>-11</sup> 2.277 2.207 1.921 1.886 1.627 1.618	2.1158 <sup>-11</sup> 1.8535 1.801 <sup>4</sup> 1.5679 1.5398 1.3282 1.3210	9.10607 9.09498 9.09228 9.08013 9.07852 9.06529 9.06480	.9285607 .9274297 .9271547 .9259154 .9257519 .9244022 .9243522
260,000 260,235 265,000 265,657 270,000 271,088 275,000 276,528	249, 784 250, 000 254, 395 255, 000 258, 999 260, 000 263, 597 265, 500	4.986 <sup>-8</sup> 4.956 4.400 4.329 3.893 3.792 3.454 3.332	4.9204 <sup>-11</sup> 4.8913 4.3421 4.2722 3.8423 3.7428 3.4092 3.2886	1.793-11 1.384 1.204 1.181 1.043 1.012 9.071-12 8.697	1.1374-11 1.1295 9.8261-12 9.6414 8.5166 8.2591 7.4048 7.0992	9.05046	.9229558 .9228904 .9215625 .9213797 .9201724 .9198703 .9187854 .9183622
280,000 281,977 285,000 287,435 290,000 292,902 295,000 298,377	2.0 2.7 273.0 273.0 277,567 280,000 281,917 285,000	3.073-8 2.936 2.740 2.594 2.449 2.297 2.194 2.040	3.0327 <sup>-11</sup> 2.8975 2.7044 2.5598 2.4173 2.2672 2.1655 2.0130	7.910 <sup>-12</sup> 7.499 6.918 6.487 6.067 5.629 5.334 4.898	6.4574 <sup>-12</sup> 6.1220 5.6473 5.2958 4.9525 4.5948 4.3546 3.9980	8.99664 8.99128 8.98709 8.97651 8.96958 8.96176 8.95611 8.94702	.9174015 .9168552 .9160207 .9153496 .9146431 .9138451 .9132686 .9123419
300,000 303,862 305,000 309,356 310,000 314,859	286,480 290,000 291,036 295,000 295,585 300,000	1.970 <sup>-8</sup> 1.815 1.772 1.619 1.598 1.447	1.9442 <sup>-11</sup> 1.7914 1.7492 1.5979 1.5769 1.4284	4.273 4.156 3.738 3.681 3.279	3 4883 3.3923 3.0517 3.0048 2.6765	8.93229 8.92924 8.91757 8.91585 8.90287	.9118972 .9108399 .9105288 .9093391 .9091636 .9078396
320,000 325,893 330,000 336,963 340,000 348,069	304,663 310,000 513,714 320,000 322,738 330,000	1.306 <sup>-8</sup> 1.16 <sup>4</sup> 1.075 9.431 <sup>-9</sup> 8.91 <sup>4</sup> 7.698	1.1485	2.541 2.316	2.3736 <sup>-12</sup> 2.0739 1.8909 1.6217 1.5183 1.2788	2 8.88916 8.87349 8.86259 8.84417 8.83615 8.81489	.9064422 .9048444 .9037331 .9018540 .9010361 .8988686

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# METRIC TABLE II CONȚINUED

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	•		CCIEF	DENS	TTY	ACCELERA OF GRA	ATION VITI
ALTI	TUDE	PRE	SSURE		1	, 2	g/go
Z,m	H,m'	P,mb	P/P <sub>o</sub>	ρ,kg/m <sup>3</sup>	P/Po	g,m/sec <sup>2</sup>	<u>ку ко</u>
350,000 359,213 360,000 370,000 370,394 380,000 381,612	331,735 340,000 340,705 349,648 350,000 358,565 360,000	7.437 <sup>-9</sup> 6.326 6.241 5.266 5.232 4.467 4.352	7.3393 <sup>-12</sup> 6.2432 6.1589 5.1973 5.1631 4.4088 4.2954	1.245 1.226 1.005 9.971-13 8.289 8.040	1.2282-12 1.0165 1.0005 9.2026-1 8.1396 6.7665 6.5634 5.6141	2 8.80982 8.78566 8.78560 3 8.75751 8.75648 8.73153 8.72735 8.72735	.89835,12 .8958882 .8956782 .8930172 .8929127 .8903680 .8899421 .8877305
390,000	367,456 370,000	3.808 3.641	3.7584 3.5934	6.877 6.526	5.3270	8.69827	.8869765
392,867 400,000 404,160 410,000 415,491 420,000 426,860 430,000 438,267 440,000 449,713 450,000 461,197 470,000 472,721 480,000 484,283 490,000 495,884	376, 320 380,000 385,158 390,000 400,000 402,756 410,000 411,516 420,000 428,959 450,000 437,642 440,000 446,300 454,932	3.262-9 3.062 2.806 2.586 2.424 2.197 2.102 1.874 1.830 1.605 1.599-9 1.402 1.380 1.233 1.191 1.088 1.032 9.625-1	3.2190 <sup>-14</sup> 3.0220 2.7691 2.5517 2.3922 2.1687 2.0747 1.8496 1.8062 1.5841 1.3621 1.3621 1.2168 1.1756 1.0735 1.0184	2 5.737 <sup>-13</sup> 5.329 4.811 4.373 4.054 5.615 3.432 3.000 2.918 2.503 2.136 2.036 1.839 1.766 1.588 1.493	4.5501 3.9273 3.5696 3.3096 2.9508 2.7848 2.4495 2.3824 2.0434	3 8.67991 8.66923 8.65428 8.64025 8.62876 8.61131 8.60335 8.56242 8.57805 8.55358 13 8.55286 8.52779 8.52479 8.50282 8.49605 6.47797 6.46735 8.45322 8.45322	.8851048 .8840159 .8824907 .8810602 .8798882 .8781094 .8772971 .8751636 .8747175 .8722228 .8721492 .8695923 .8692869 .8670466 .8663559 .8645120 .8634299 .8619885 .8605088
500,000 507,525 510,000 519,205 520,000 530,000 530,925 540,000 542,68	463,540 5 470,000 172,122 5 480,000 180,679 189,212 5 490,000 197,719	7.821 7.600 6.841 6.780 6.064 6.002 7.436	10 8.4293-1 7.7184 7.5003 6.7515 6.6912 5.9842 5.9234 5.3647 5.2117	13 1.197 <sup>-1</sup> 1.080 1.044 9.231 <sup>-1</sup> 9.135 8.016 7.921 7.054 6.819	8.8124 8.5219	14 8.42858 8.41011 8.40405 8.38156 8.37963 8.35531 8.35306 8.35110 8.32461	.8495354

#### METRIC TABLE III

Velocity of Sound, Particle Speed, Molecular-Scale Temperature Gradient, and Scale Height as Functions of Germetric and Geopotential Altitude

TILIA	Tule	MOL-SCALE TEMP.GRAD	SCALE	HEIGHT,	PARTIGLE RATIO SPIED		
Z,n	H <sub>p</sub> m <sup>4</sup>	Lي•Q/m'	H <sub>g</sub> , km	H <sub>s</sub> /H <sub>so</sub>	√,m/sec	V/V₀ c₅/c₅o	Cg,m/sec
-5,000 -4,996.1 -4,000 -3,997.5 -3,000 -2,998.6 -2,000 -1,999.4 -1,000 - 999.8	-5,003,9 -5,000 -4,002,5 -4,000 -3,001,4 -3,000 -2,000,6 -2,000 -1,000,2 -1,000		9.3717 9.3709 9.1843 9.1839 8.9969 8.9967 8.8095 8.6220 8.6220	1.11112 1.11104 1.08891 1.08886 1.06670 1.06666 1.04447 1.04446 1.02224	484.15 484.15 479.21 479.20 474.22 474.22 469.19 469.18 464.09 464.09	1.05493 1.05489 1.04417 1.04414 1.03330 1.03328 1.02232 1.02231 1.01122 1.01122	358.98 358.97 355.32 355.31 351.62 351.62 347.89 347.88 344.11 344.11
0 1,000 1,000.2 2,000 2,000.6 3,000 3,001.4 4,000 4,002.5	999.8 1,000 1,999.4 2,000 2,998.6 3,000 3,997.5 4,000	-o.op65	8.2468 8.2468 8.0591 8.0590 7.8713 7.8711 7.6835 7.6831	•977754 •977751 •955501 •955487 •933241 •933210 •910975	453.74 453.74 448.47 448.47 443.15 443.11 437.76	•988659 •988657 •977190 •977183 •965589 •965572 •953850	336.43 336.43 332.53 332.53 328.58 328.58 324.59
5,000 5,003.9 6,000 6,005.7 7,000 7,007.7 8,000 8,010.1 9,000 9,012.8	4,996.1 5,000 5,994.3 6,000 6,992.3 7,000 7,989.9 8,000 8,987.3 9,000		7.4957 7.1919 7.5077 7.3067 7.1198 7.1183 6.9317 6.9298 6.7436 6.7412	.888700 .898613 .866419 .866293 .844131 .843959 .821836 .821611 .799534	432.29 426.79 426.76 421.20 421.15 415.53 415.47 409.79	.941968 .941921 .929939 .929870 .917756 .917661 .905412 .892902	20.53 516.45 516.43 512.30 512.27 508.10 508.06 503.85
10,000 10,016 11,000 11,019 12,000 12,023 13,000 13,027 14,000 14,031	9,984.3 10,000 10,981 11,000 11,977 12,000 12,973 13,000 13,969 14,000	0.0000	6.5554 6.5525 6.3672 6.3636 6.3656 6.3656 6.3676 6.3696 6.3696	•777225 •776873 •754908 •754483 •754715 •754720 •754958 •755189	403.88 398.07 397.95 397.95 397.95 397.95 397.95 397.95	.880219 .800018 .867554 .867107 .867107 .867107 .867107 .867107 .867107	299.46 295.15 295.07 295.07 295.07 295.07 295.07

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ALTI	TUDE	MOL-SCALE TEMP.GRAD.	SCALE	height	PARTICLE SPEED	→RATIO ←	SOUND SPEED
Z,m	H,m'	L <sub>l,i</sub> , °C/m'	H <sub>s</sub> , km	H <sub>B</sub> /H <sub>SO</sub>	⊽,m/sec	V/Vo c <sub>s</sub> /c <sub>so</sub>	C <sub>s</sub> ,n√sec
15,000 15,035 16,000 16,040 17,000 17,046 18,000 18,051 19,000	14,965 15,000 15,960 16,955 17,000 17,949 18,000 18,943 .19,000		6.3716 6.3717 6.3736 6.3737 6.3757 6.3776 6.3777 6.3796 6.3797	•755426 •755435 •755664 •755673 •755915 •756140 •756152 •756377 •756389	397.95 397.95 397.95 397.95 397.95 397.95 397.95 397.95 397.95	.867107 .867107 .867107 .867107 .867107 .867107 .867107 .867107	295.07 295.07 295.07 295.07 295.07 295.07 295.07 295.07 295.07
20,000 20,063 21,000 21,070 22,000 22,076 23,000 23,084 24,000 24,091	19,937 20,000 20,931 21,000 21,924 22,000 22,917 23,000 23,910 24,000	0.0000	6.3816 6.3817 6.3836 6.3837 6.3856 6.3857 6.3876 6.3878 6.3896 6.3898	•756615 •756626 •756852 •756864 •757089 •757101 •757326 •757350 •757563 •757587	397.95 397.95 397.95 397.95 397.95 397.95 397.95 397.95 397.95	.867107 .867107 .867107 .867107 .867107 .867107 .867107 .867107 .867107	295.07 295.07 295.07 295.07 295.07 295.07 295.07 295.07 295.07
25,000 25,099 26,000 26,107 27,000 27,115 28,000 28,124 29,000 29,133	24,902 25,000 25,894 26,000 26,886 27,000 27,877 28,000 28,868 29,000	+0.0030	6.3916 6.3918 6.4728 6.4823 6.5626 6.5730 6.6525 6.6636 6.7424 6.7543	.757800 .757824 .767427 .768554 .778074 .779307 .788733 .790049 .799392 .800803	397.95 397.95 400.41 400.70 403.11 403.42 405.80 406.13 408.58 408.82	.867107 .867107 .8721:58 .873089 .878355 .879031 .884206 .884933 .890026	295.07 295.07 296.89 297.11 298.90 299.13 300.89 301.14 302.87 303.13
30,000 30,142 31,000 31,152 32,000 32,162 33,000 33,172 34,000 34,183	29,859 30,000 30,850 31,000 31,840 32,000 32,830 33,000 33,819 34,000		6.8323 6.8451 6.9223 6.9360 7.0123 7.0269 7.1023 7.1178 7.1923 7.2088	.610050 .811568 .820721 .822345 .851392 .835123 .842062 .843900 .852733 .854689	411.12 411.50 413.75 414.15 416.37 416.79 418.97 419.41 421.55 422.02	.895802 .896621 .901539 .902407 .907238 .908158 .912900 .913871 .918525 .919550	304.83 305.11 306.79 307.08 308.13 309.04 310.65 310.98 312.57 312.92

TULA,	'rude	MOL SCALE	SCALE I	गहर (स्था	PARTICLE SPEED	~>RAUTO≪	SCOID SPEED
Z,m	H,m'	L <sub>台</sub> ,*C/m'	H <sub>H</sub> , km	e <sub>s</sub> /H <sub>so</sub>	V,m/sec	v/√o cu/cao	C <sub>H</sub> , 11/450
35,000 35,194 36,000 36,205 37,000 37,217 38,000 38,229 39,000 39,241	34,808 35,000 35,797 36,000 36,786 37,000 37,774 33,000 38,762 39,000		7.2824 7.2999 7.3725 7.3910 7.4627 7.4822 7.5528 7.5734 7.6430 7.6647	.865415 .865490 .874098 .876291 .884792 .887104 .895474 .697917 .9061.69	424.11 424.61 426.66 427.18 429.20 429.74 431.71 432.29 434.22 434.82	.924114 .925193 .929668 .930803 .935187 .936378 .940672 .941921 .946124 .947433	214.47 214.34 216.26 216.74 318.24 318.64 320.10 320.53 321.96 322.40
40,000 40,253 41,000 41,266 42,000 42,279 43,000 43,293 44,000 44,307	39,750 40,000 40,737 41,000 41,724 42,000 42,711 43,000 43,698 44,000	+0.0030	7.7332 7.7561 7.8235 7.8475 7.9138 7.9389 8.0040 8.0305 8.0943 6.1220	.916863 .919578 .927569 .930414 .938275 .941251 .948969 .952111 .959676	436.70 437.33 439.17 439.83 441.63 442.32 444.08 444.79 446.50 447.25	.951543 .952910 .956929 .958357 .962283 .963773 .967606 .969159 .972899	323.80 324.27 325.64 326.12 327.46 327.96 329.27 329.80 331.62
45,000 45,321 46,000 46,335 47,000 47,350 48,000 48,365 49,000	44,684 45,000 45,670 46,655 47,000 47,640 48,000 48,625 49,000	<u> </u>	8.1847 8.2137 8.2751 8.3054 8.3655 8.3971 8.3998 8.3998 8.4015 8.4025	.970394 .973832 .981111 .984704 .991829 .995576 .995896 .995896	448.92 449.69 451.32 452.12 453.71 454.54 454.54 454.54 454.54	.978161 .979843 .983393 .985141 .988596 .590411 .990411 .990411	332.86 333.43 354.64 335.24 337.03 337.03 337.03 337.03 337.03
50,000 50,396 51,000 51,412 52,020 52,429 53,446 54,000 54,463	49,610 50,000 50,594 51,000 51,578 52,562 53,562 53,545 54,000	0.0000	8.4041 8.4051 8.4067 8.4078 8.4093 8.4105 8.4100 8.4131 8.7513 8.2997	.996406 .996525 .996714 .996845 .997165 .997165 .997473 .997473	454.54 454.54 454.54 454.54 454.54 454.54 452.83	.590411 .590411 .990411 .990411 .990411 .990411 .936679 .983554	337.03 357.03 357.03 337.03 337.03 337.03 337.03 337.03 337.70

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ALTZ	TUDE	MOL-SCALE	SCALE F	EIGHT	PARTICLE SPEED	-> PATIO &	SPEED
Z,m	H,m'	L <sub>M</sub> ,°C/m'	H <sub>S</sub> , km	H <sub>s</sub> /H <sub>so</sub>	⊽;m/sec	⊽/V <sub>o</sub> c <sub>c</sub> /c <sub>so</sub>	C <sub>s</sub> ,m/sec
55,000 55,480 56,000 56,498 57,000 57,516 58,000 58,534 59,000 59,553 60,000 60,572	54,528 55,000 55,511 56,000 56,493 57,000 57,476 58,000 58,457 59,000		8.2397 8.1862 8.1281 8.0726 8.0165 7.9589 7.9048 7.8452 7.7931 7.7314 7.6814 7.6175	.976918 .970567 .963687 .957100 .950451 .943623 .937121 .930139 .923967 .916645 .910719 .903143	435.32 433.83 432.03	.979913 .976650 .975103 .969696 .966247 .962692 .959344 .955637 .952393 .948530 .945393 .941368 .938343	533.46 532.55 531.14 529.98 528.81 527.60 526.46 525.20 521.71 520.34 319.31
61,000 61,591 62,000 60,611 63,000 63,631 64,000 61,651	60,420 61,000 61,401 62,000 62,382 63,000 63,362 64,000		7.5696 7.5035 7.4578 7.3895 7.3459 7.2754 7.2341 7.1612	.897467 .839632 .884210 .876113 .870949 .860585 .857685 .849048	428.72 427.39 425.38 424.10 422.02 420.69 418.63	.934152 .931242 .926880 .924088 .919550 .916881 .912161	317.88 316.89 315.41 314.46 312.92 512.01 310.40
65,000 65,672 66,000 66,692 67,714 68,000 68,735 69,000 69,757			7.1821 7.0470 7.0102 6.9327 6.8982 6.8183 6.7862 6.7108 6.6741 6.5893	.844415 .835503 .831142 .821949 .817865 .808386 .604583 .795644 .791697 .781235	415.21 414.10 411.76 410.72 408.29 407.31 404.78 403.87	.909620 .904712 .902502 .897001 .894906 .889627 .887499 .879997	307.87 307.05 305.31 504.54 302.73 302.01 300.13 299.46 297.51
70,000 70,779 71,000 71,802 72,000 72,825 73,000 73,848 74,000 74,872	70,000 70,21: 71,000 71,19 72,000 72,17: 75,000 73,14	0 6 0 14 0 1 0 8	6.5620 6.4746 6.1499 6.3600 6.3377 6.2452 6.2255 6.1304 6.1133	.77800° .767646 .76471 .754049 .75141 .74044 .73811 .726826 .72480	5 397.68 5 396.90 9 394.08 5 393.37 5 390.44 5 389.82 8 386.77 6 386.23		294.87 294.29 292.20 291.68 289.50 289.04 286.78 286.38

ALTI	TUDE	MOL-SCALE TEMP.GRAD.	SCALE	IEIGHT	PARTICLE SPEED	→ RATIO «	SOUND SIEED
Z,nı	H,m'	C/m'	II <sub>s</sub> , kan	ll <sub>e</sub> /il <sub>so</sub>	V,1n/BGC	V/Vo Cs/Cso	C <sub>s</sub> ,m/see
75,000 75,895 76,000 76,920 77,000 77,944 78,000 78,969 79,000 79,994	7 <sup>1</sup> 4,125 75,000 75,102 76,000 76,078 77,000 77,055 78,000 .78,030 79,000	-0.0039 중	6.0010 5.900 5.901 5.902 5.903 5.904 5.904 5.906 5.906 5.908	.71150 .69957 .69960 .69980 .69981 .70002 .70003 .70024 .70025 .70046	382.60 379.3 379.3 379.3 379.3 379.3 379.3 379.3 379.3	.833666 .82654 .82654 .82654 .82654 .82654 .82654 .82654 .82654	283.69 281.26 281.26 281.26 281.26 281.26 281.26 281.26 281.26
80,000 81,000 81,020 82,000 82,045 83,000 83,072 84,000 84,098	79,006 79,981 80,000 80,956 81,000 81,930 82,000 82,904 83,000	0.0000	5.908 5.910 5.910 5.912 5.914 5.914 5.915 5.916	.70047 .70068 .70069 .70090 .70091 .70112 .70113 .70134	379.3 379.3 379.3 379.3 379.3 379.3 379.3	.82654 .82654 .82654 .82654 .82654 .82654 .82654 .82654	281.26 281.26 281.26 281.26 281.26 281.26 281.26 281.26
85,000 85,125 86,000 86,152 87,000 87,179 88,000 88,207 89,000 89,235	85,878 84,000 84,852 85,000 85,825 86,000 86,798 87,000 87,771 88,000		5.917 5.917 5.919 5.919 5.921 5.923 5.923 5.925 5.925	.70155 .70158 .70177 .70181 .70199 .70203 .70221 .70225 .70243	379.3 379.3 379.3 379.3 379.3 379.3 379.3 379.3	.82654 .82654 .82654 .82654 .82654 .82654 .82654 .82654 .82654	281.26 281.26 281.26 281.26 281.26 281.26 281.26 281.26 281.26
90,000 90,264 91,000 91,293 92,000 92,322 93,000 93,351 94,000 94,381	88,744 89,000 89,716 90,000 90,688 91,000 91,659 92,000 92,630 93,000	+0.0035	5.926 5.927 5.928 5.929 5.984 6.036 6.107 6.144 6.211 6.251	.70264 .70270 .70286 .70293 .70944 .71565 .72404 .72839 .73641 .74113	379.3 379.3 379.3 379.3 381.6 382.7 384.9 386.0 388.1 389.3	.82654 .82654 .82654 .82654 .83157 .83385 .83863 .84110 .84564 .84829	281.26 281.26 281.26 281.26

ALTIT	UDE	MOL-SCALE TEMP.GRAD	SCALE	HFIGHT	PARTICL	E SPEED
Z,m	H,m'	L <sub>M</sub> , C/m'	H <sub>s</sub> ,kn	H <sub>s</sub> /H <sub>so</sub>	V,m/sec	₹/₹ <sub>0</sub>
2,m  95,000 95,411 96,000 96,441 97,000 97,472 98,000 98,503 99,000 99,534  100,000 101,598 102,000 103,663 104,000 103,663 104,000 105,730 106,764 107,000 106,764 107,798 108,000 107,798 108,000 109,867 110,000 110,902	H,m'  93,601 94,000 94,572 95,000 95,542 96,000 96,512 97,000 97,482 98,000  98,451 99,000 100,389 101,000 101,358 102,000 103,394 104,000 104,261 105,000 105,229 106,000 105,229 106,000 107,162 108,000 108,129 109,000	<del> </del>	H <sub>B</sub> , lnt  6.316 6.359 6.420 6.466 6.524 6.574 6.629 6.682 6.733 6.789 6.897 7.047 7.1152 7.225 7.329 7.361 7.466 7.574 7.676 7.763 7.780 7.871 7.885 7.980	H <sub>B</sub> /H <sub>BO</sub> ·74879 ·75388 ·76117 ·76664 ·77355 ·77940 ·78594 ·79218 ·79833 ·80496  ·81073 ·81775 ·82313 ·83055 ·84794 ·85617 ·86035 ·84794 ·85617 ·86035 ·84899  ·87276 ·88182 ·889466 ·89760 ·991003 ·92246 ·93323  ·93489 ·94610	₹,m/sec 391.3 394.4 395.8 397.6 403.2 405.4 406.6 409.7 408.8 409.7 412.9 412	85259 85259 85259 85259 85259 85259 85259 86249 86249 86330 86950 87646 87976 87976 88646 89964 899645 912662 912662 91269
111,000 111,937 112,000 112,973 113,000	109,095 110,000 110,061 111,000 111,026		7.990 8.088 8.095 8.197 8.2% 8.305	•94733 •95898 •95977 •97187 •97221 •98466	439.0 441.7 441.8 444.5 444.6 447.4	.95660 .96233 .96272 .96862 .96879 .97482
114,000 114,009	111,992		8.306	•98477	447.4	.97487

AI/II	TUDE	MOL-SCALE TEMP.GRAD.	SCALE	HEIGHT	PARTICI.	e speed
Z,m	H,m¹	L <sub>M</sub> , °C/m'	H <sub>s</sub> ,kan	H <sub>s</sub> /II <sub>so</sub>	V̄,m/sec	₹/Ÿ <sub>o</sub>
115,000 115,045 116,000 116,082 117,000 117,119 118,000 118,156 119,000	112,957 113,000 113,921 114,000 114,885 115,000 115,850 116,000		8.410 8.415 8.515 8.524 8.620 6.633 8.725 8.742 8.831	.99712 .99768 1.0096 1.0106 1.0220 1.0235 1.0345 1.0364	450.1 450.3 452.9 453.1 455.6 455.9 458.3 460.0	.98081 .98108 .98677 .98725 .99268 .99338 .99856 .99948
119,194 120,000 120,232 121,000 121,270 122,000 122,309 123,000 123,348 124,000 124,387	117,000 117,777 118,000 118,740 119,000 119,703 120,000 120,665 121,000 121,627 122,000	+0.0035	8.851 8.936 8.960 9.041 9.069 9.146 9.179 9.251 9.288 9.357 9.398	1.0494 1.0594 1.0623 1.0719 1.0753 1.0844 1.0882 1.0969 1.1012 1.1094 1.1142	461.5 463.6 464.2 466.3 467.0 468.9 469.7 471.5 472.4 474.1 475.1	1.0055 1.0102 1.0116 1.0160 1.0175 1.0217 1.0235 1.0274 1.0294 1.0353 1.0353
125,000 125,427 126,000 126,467 127,000 127,507 128,000 128,548 129,000 129,589 130,000 130,630 131,600 131,672 132,000 132,774 135,000 137,929	122,589 123,500 123,551 124,000 124,512 125,000 125,473 126,000 126,434 127,000 127,395 128,000 128,355 129,000 129,315 130,000 132,193 135,000	+0.0100	9.462 9.507 9.567 9.617 9.673 9.726 9.778 9.836 9.970 10.14 10.27 10.45 10.76 10.86 11.07 11.75 12.61	1.1218 1.1272 1.1343 1.1402 1.1468 1.1532 1.1593 1.1662 1.1820 1.2027 1.2171 1.2392 1.2522 1.2758 1.2673 1.3124 1.3926 1.4956	476.7 477.8 479.3 480.5 481.9 485.8 485.0 493.3 496.6 5003.2 5003.9 515.0 530.3 549.3	1.0367 1.0411 1.0449 1.0470 1.0527 1.0555 1.0585 1.0656 1.0748 1.0811 1.0908 1.0964 1.1066 1.1115 1.1222 1.1556 1.1970

ALTIT	e fore	MOL-SCALE	SCALE	HEIGHT	PARTICL	SPEED
	H,m'	TEMP.GRAD.	H <sub>s</sub> ,km	H <sub>s</sub> /H <sub>so</sub>	⊽,m/sec	⊽/v̄o
2,m 140,000 143,153 145,000 148,385 150,000 153,625 155,000 158,874	136,983 140,000 141,766 145,000 146,542 150,000 151,311 155,000	+0.0100	13.23 14.16 14.71 15.72 16.20 17.28 17.69 18.85	1.5684 1.6794 1.7445 1.8638 1.9208 2.0488 2.0974 2.2344	. 675.1	1.2254 1.2674 1.2913 1.3266 1.3540 1.3976 1.4138 1.4584
160,000 164,13i 165,000 169,397 170,000 174,671 175,000 179,954	156,072 160,000 160,826 165,000 165,572 170,000 170,311 175,000	5 <del>\frac{17}{44}</del>	20.42 20.68 21.99 22.17 23.57 25.16 25.16	2.4206 2.4514 2.6074 2.6288 2.7948 2.8065 2.9828	696.1 700.4 721.8 724.7 746.7 748.3 770.8	1.5167 1.5261 1.5728 1.5791 1.6271 1.6304 1.6795
185,000 185,245 190,000 190,545 195,000	179,760 180,000 184,480 185,000 189,19	6	26.05 26.10 26.94 27.04 27.83 27.99	3.0943 3.1945 3.206 3.3003 3.3186	784.4 796.5 797.8 808.9 811.0	1.7092 1.7355 1.7384 1.7626 1.7671
200,000 201,171 205,000 206,497 210,000 211,831 215,000 217,175	. 195,00 . 198,59 . 200,00 . 203,28 . 205,00 . 207,96	0 15 10 10 10 10 10 10 10 10 10 10 10 10 10	28.73 28.94 29.62 29.89 30.53 30.89 31.4: 31.8	3.430 2 3.511 3.543 1 3.617 4 3.656 1 3.724	7 824.0 8 833.1 5 836.7 9 845.0 7 849.3	1.7954 1.8154 1.8232 1.8412 1.8506 1.8667
920,000 222,521 225,000 227,80 230,00 233,25 235,00 238,63	212,6 6 215,0 0 217,3 7 220,0 0 221,9 6 225,0 0 226,6	41 00 08 00 69 00 22	32.3 32.7 33.2 33.7 34.1 34.6 35.0	6 3.884 1 3.937 3 3.998 1 4.04 9 4.11	43 873.9 70 879.5 86 886.0 87 890.7 833 897.8	1.9042 1.9164 1.9304 1.9407 1.9563 7 1.9647

ALTIT	TIME I	MOL-SCALE	SCALE	HEIGHT	PARTICLE	SPEED
CLTAIA .		TEMP.GRAD.	<u>  </u>	H <sub>s</sub> /H <sub>so</sub>	⊽,m/sec	$\overline{\overline{v}}/\overline{v}_{o}$
Z,n	H,m'	L <sub>M</sub> , °C/m'	H <sub>s</sub> ,km			
240,000	231,268		35.91	4.2576	912.5	1.9883 2.0071
244,021	235,000	ļ	36.64	11.3438	921.2 923.2	2.0117
245,000	235,908		36.81.	4.3647	932.6	2.0320
249,417	21,0,000	1	37.61	4.4595	933.8	2.0347
250,000	240,540	1	37.72	4.4721	943.9	2.0566
254,621	245,000	1	38.59	4.5757	944.3	2.0575
255,000	245,165		38.63	4.5796	947.7	
	alia 78h	1	39.53	4.6872	954.6	2.0799
260,000	249,784	l l	39.58	4.6923	955.0	2.0810
260,235	250,000 254,395		40.44	4.7950	964.8	2.1021
265,000	255,000		40.56	4.8092	966.1	2.1050 2.1240
265,657	258 999		41.35	4.9030	974.8	2.1288
270,000	260,000	1	41.55	4.9265	977.0	2.1200
271,088	263,597	ł	42.27	5.0111	984.8	2.1523
275,000 276,528	265,000	ļ	42.54	5.0442	987.8	2.1727
		1	1. z 1 R	5.1194	994.6	2.1671
280,000	268,187	+0.0058	43.18 43.54	5.1622		2.1755
281,977	270,000	i	44.09	5.2278	1,004	2.1883
285,000	272,771		44.54	5.2806	1,009	2.1986
287,435	275,000	1	45.01	5.3364	1,014	2.2093
290,000	277,347	l	45.54	5.3995		2.2168
292,902	280,000		45.93	5.4451	1,023	2.2300
295,000	281,917		46.55	5.5187	1,030	2.2439
298,377	285,000		•		- 037	0.0505
300,000	286,480	,	46.84		1,033	2.2505 2.2662
303,862			47.56		1,040	2.2708
305,000			47.76			2.2883
309,356			48.57	5.7582	1,050	2.2908
310,000			48.69			2.3102
314,859			49.58	5.8786	, 1,000	<u> </u>
		.	50.53	5.9912	2 1,070	2.3304
320,000	304,66	ζ	51.62	6.120	5 1,080	2.3533
<del>525,893</del>	310,000 313,71	ĭ. I	52.38	6.210		2.3691
350,000			53.68	· 6.3644		2.3957
336,963			54.24	6.430		2.4072
340,000 348,069		ŏ l	55.74	6.609	0 1,119	2.4373
240,00	, ,,,,,,,,,					

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ALTITUDE	MOL-SCALE TEMP.GRAD.	SCALE HEIGHT	PARTICLE SPEED
Z,m H,m'	IM, C/m'	H <sub>s</sub> ,km H <sub>s</sub> /H <sub>so</sub>	V,m/sec V/V <sub>O</sub>
350,000 331,735 359,213 340,000 360,000 349,648 370,394 350,000 380,000 358,565 381,612 360,000 390,000 367,456 392,867 370,000 400,000 376,320 404,160 380,000 410,000 385,158 415,491 390,000 420,000 393,970 426,860 400,000 430,000 402,756 438,267 410,000 440,000 411,516 449,713 420,000 450,000 428,959 461,197 430,000 470,000 437,642 472,721 440,000 484,283 450,000 484,283 450,000 495,884 460,000 507,525 470,000 507,525 470,000 510,000 472,122 519,205 480,000	<del> </del>	56.10 6.6517 57.82 6.8557 57.97 6.8731 59.84 7.0952 59.92 7.1040 61.72 7.3179 62.03 7.3538 63.61 7.5412 64.15 7.6053 65.49 7.7652 66.28 7.8585 67.39 7.9897 68.43 8.1133 69.29 8.2149 70.59 8.3698 71.19 8.4408 72.77 8.6280 73.10 8.6672 74.96 8.8878 75.02 8.8943 76.94 9.1221 77.17 9.1494 78.87 9.3504 79.39 9.4127 80.80 9.5794 81.63 9.6777 82.73 9.8090 83.88 9.9444 84.68 10.0393 86.14 10.2129 86.62 10.2701 88.42 10.4832	1,122 2.4445 1,137 2.4783 1,139 2.4812 1,155 2.5172 1,156 2.5186 1,171 2.5582 1,187 2.5582 1,187 2.5587 1,192 2.5973 1,201 2.6178 1,210 2.6357 1,219 2.6553 1,227 2.6736 1,234 2.6885 1,244 2.7110 1,249 2.7212 1,261 2.7479 1,264 2.7534 1,278 2.7843 1,278 2.7852 1,294 2.8165 1,294 2.8165 1,307 2.8473 1,311 2.8556 1,322 2.8812 1,307 2.8473 1,311 2.8556 1,322 2.8812 1,327 2.9978 1,343 2.9253 1,348 2.9595 1,362 2.9667 1,374 2.9933
520,000 480,679 530,000 489,212 530,925 490,000 540,000 497,719 542,686 500,000		88.58 10.5016 90.53 10.7338 90.71 10.7553 92.50 10.9665 93.02 11.0292	1,375 2.9956 1,388 3.0241 1,389 3.0267 1,401 3.0523 1,404 3.0598

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#### METRIC TABLE IV

VISCOSITY, KINEMATIC VISCOSITY, AND SPECIFIC WEIGHT AS FUNCTIONS OF GEGMETRIC AND GEOPOTENTIAL ALTITUDE

ALTI	TUDE	visco	SITY		VISCOSITY	SPECIFI	C WEIGHT
Z,m	H,m'	$\mu$ , $\frac{kg}{m \text{ sec}}$	μ/μ <sub>0</sub>	$ \eta, \frac{m^2}{\sec} $	η/η <sub>ο</sub>	m, kg m <sup>2</sup> sec <sup>2</sup>	ω/ω <sub>o</sub>
-5,000 -4,995.1 -4,000 -3,997.5 -3,000 -2,998.6 -2,000 -1,999.4 -1,000 999.8	-5,003.9 -5,000 -4,002.5 -4,000 -3,001.4 -3,000 -2,000.6 -2,000 -1,000.2 -1,000	1.9423 <sup>-5</sup> 1.9422 1.9123 1.9123 1.8821 1.8820 1.8515 1.8515 1.8206 1.8206	1.08542 1.08536 1.06863 1.06864 1.05177 1.05175 1.03469 1.03468 1.01744	1.0058 <sup>-5</sup> 1.0060 1.0805 1.0808 1.1625 1.1627 1.2526 1.2526 1.3516	6.88539 6.88717 7.39715 7.39864 7.95852 7.95935 8.57478 8.57517 9.25274 9.25289	1.8968 <sup>+1</sup> 1.8961 1.7378 1.7373 1.5891 1.5889 1.4505 1.4504 1.3214	1.57892 1.57839 1.44654 1.44619 1.32282 1.32265 1.20743 1.20735 1.09994 1.09993
0 1,000 1,000.2 2,000 2,000.6 3,000 3,001.4 4,000 4,002.5	0 , 999.8 1,000 1,999.4 2,000 2,998.6 3,000 3,997.5 4,000	1.7894 <sup>-5</sup> 1.7579 1.7579 1.7260 1.7260 1.6938 1.6938 1.6612 1.6612	1.00000 .982380 .982377 .964571 .964560 .946567 .946542 .928364	1.4607 <sup>-5</sup> 1.5813 1.5813 1.7148 1.7149 1.8629 1.8631 2.0275 2.0279	1.00000 <sup>+0</sup> 1.08254 1.08255 1.17391 1.17397 1.27528 1.27543 1.38801 1.38830	1.2013 <sup>+1</sup> 1.0898 1.0898 9.8648 <sup>+0</sup> 9.8642 8.9083 8.9071 8.0249 8.0228	1.00000 9.07189-1 9.07179 8.21154 8.21105 7.41543 7.41437 6.68006 6.67830
5,000 5,003.9 6,000 6,005.7 7,000 7,007.7 8,000 8,010.1 9,000 9,012.8	4,996.1 5,000 5,994.3 6,000 6,992.3 7,000 7,989.9 8,000 8,987.3 9,000	1.6285 <sup>-5</sup> 1.6282 1.5950 1.5948 1.5613 1.5610 1.5272 1.5268 1.4927 1.4922	.909955 .909882 .891335 .891229 .872499 .872352 .853439 .853246 .834151	2.2111-5 2.2118 2.4162 2.4175 2.6461 2.6480 2.9045 2.9073 3.1958 3.1998	1.51366 <sup>40</sup> 1.51418 1.65411 1.65496 1.81150 1.81281 1.93841 1.99031 2.18782 2.19053	7.2106 <sup>+0</sup> 7.2075 6.4613 6.4572 5.7734 5.7683 5.1432 5.1372 4.5674 4.5603	6.00216 <sup>-1</sup> 5.99961 5.37843 5.37503 4.80584 4.80157 4.28128 4.27625 3.80193 3.79608
10,000 10,016 11,000 11,019 12,000 12,025 13,000 13,027 14,000 14,051	9,984.3 10,000 10,981 11,000 11,977 12,000 12,973 13,000 13,969 14,000	1.4577 <sup>-5</sup> 1.4572 1.4223 1.4217 1.4217 1.4217 1.4217 1.4217 1.4217	.814627 .814317 .794861 .794482 .794482 .794482 .794482 .794482	3.9066 4.5576 4.5739 5.3327 5.3551 6.2394	2.41332 <sup>+0</sup> 2.41711 2.66918 2.67440 3.12005 3.13119 3.65070 3.66601 4.27139 4.29217	4.0424 <sup>+0</sup> 4.0344 3.5651 3.5564 3.0475 3.0366 2.6037 2.5928 2.2247 2.2139	3.36494 <sup>-1</sup> 3.35831 2.96764 2.96042 2.53678 2.52774 2.16737 2.15830 1.85184 1.84286

ALTI	TUDE	VISCOS	YTE	KINEMATIC	VISCOSITY	SPECIFIC	WEIGHT
Z,m	H,m'	μ, kg m sec	μ/μ <sub>o</sub>	η, <u>m²</u> sec	դ/ւյ <sub>օ</sub>	ω, m <sup>2</sup> sec <sup>2</sup>	ω/ω <sub>o</sub>
15,000 15,035 16,000 16,040 17,000 17,046 18,000 18,051 19,000 19,057	14,965 15,000 15,960 16,000 16,955 17,000 17,949 18,000 18,943 19,000	1.4217 <sup>-5</sup> 1.4217 1.4217 1.4217 1.4217 1.4217 1.4217 1.4217 1.4217 1.4217 1.4217	.794482 .794482 .794482 .794482 .794482 .794482 .794482 .794482 .794482	7.2999 <sup>-5</sup> 7.3406 8.5401 8.5944 9.9907 1.0062-4 1.1687 1.1781 1.3671 1.3793	4.99737 <sup>+0</sup> 5.02527 5.84644 5.88359 6.83946 6.88852 8.00069 8.06509 9.35868 9.44261	1.9009 <sup>+0</sup> 1.8903 1.6243 1.6140 1.3881 1.3781 1.1862 1.1767 1.0138	1.58232 <sup>-1</sup> 1.57352 1.35209 1.34354 1.15543 1.14718 9.87416 <sup>-2</sup> 9.79517 8.43869 8.36357
20,000 20,063 21,000 21,070 22,000 22,076 23,000 23,084 24,000 24,091	19,937 20,000 20,931 21,000 21,924 22,000 22,917 23,000 25,910 24,000	1.4217 <sup>-5</sup> 1.4217 1.4217 1.4217 1.4217 1.4217 1.4217 1.4217 1.4217 1.4217	.794482 .794482 .794482 .794482 .794482 .794482 .794482 .794482 .794482	1.5990 <sup>-4</sup> 1.6149 1.6702 1.8907 2.1874 2.2137 2.5581 2.5918 2.9916 3.0345	1.09466 <sup>+1</sup> 1.10554 1.28032 1.29437 1.49743 1.51545 1.75124 1.77429 2.04798 2.07734	8.6644-1 8.5789 7.4056 7.4056 6.3500 6.2545 5.4108 5.3404 4.6254 4.5599	7.21234 <sup>-2</sup> 7.14119 6.16454 6.09748 5.26912 5.20633 4.50402 4.44540 3.85021 3.79569
25,000 25,099 26,000 26,107 27,000 27,115 28,000 28,124 29,000 29,133	24,902 25,000 25,894 26,000 26,886 27,000 27,877 28,000 28,868 29,000	1.4217 <sup>-5</sup> 1.4217 1.4364 1.4381 1.4526 1.4544 1.4687 1.4707 1.4847 1.4868	.794482 .794482 .802698 .803668 .811760 .812801 .820768 .821880 .829720	3.4983 <sup>-4</sup> 3.5527 4.1805 4.2613 4.9956 5.0984 5.9550 6.0850 7.0817 7.2453	2.39488 <sup>+1</sup> 2.43215 2.66191 2.91724 3.41994 3.49026 4.07667 4.16571 4.84801 4.96004	3.9541 <sup>-1</sup> 3.8934 3.3420 3.2825 2.8274 2.7739 2.3975 2.3493 2.0374 1.9941	3.29148 <sup>-2</sup> 3.24094 2.78196 2.73241 2.35357 2.30903 1.99570 1.95562 1.69596
50,000 30,142 31,000 31,152 32,000 32,162 33,000 33,172 34,000 34,183	29,859 30,000 30,850 31,000 31,840 32,000 32,830 33,000 33,819 34,000	1.5634	.838619 .839880 .847464 .848803 .856257 .857676 .864998 .866498 .873688	9.9454 1.0202-3 1.1747 1.2065 1.3844 1.4239 1.6282	5.75171 <sup>+1</sup> 5.89224 6.80848 6.98383 8.04162 8.25957 9.47770 9.74750 1.11465+2 1.14794	1.7352 <sup>-1</sup> 1.6962 1.4808 1.4459 1.2664 1.2349 1.0851 1.0569 9.3163 <sup>-2</sup> 9.0621	1.44436 <sup>-2</sup> 1.41198 1.23267 1.20355 1.05414 1.02797 9.03263 <sup>-3</sup> 8.79738 7.75503 7.54338

						<del></del>	
ALI	TUDE	VISC	OSITY	TY KINEMATIC VISCOSITY		SPECIF.	IC WEIGHT
Z,m	H,m'	μ, kg m sec	μ/μ <sub>ο</sub>	$\eta, \frac{m^2}{\sec}$	η/η <sub>ο</sub>	$\omega, \frac{kg}{m^2 sec^2}$	ω/ω <sub>0</sub>
35,000 35,194 36,000 36,20° 37,017 38,000 38,229 39,000 39,241	35,000	1.5818 1.5942 1.5974 1.6095 1.6128 1.6247 1.6282 1.6398 1.6434	.882327 .883996 .890917 .892672 .899457 .901,500 .907948 .909882 .916392 .918417	1.9708 2.2384 2.3117 2.6168 2.7062 3.0535 3.1622	1.30824 <sup>+2</sup> 1.34916 1.53239 1.58252 1.79143 1.85263 2.09034 2.16477 2.43450 2.52483	8.0137 <sup>-2</sup> 7.7 <sup>4</sup> 49 6.9060 6.6999 5.9621 5.7766 5.1562 4.9892 4.4670 4.3164	6.67071 <sup>-3</sup> 6.48022 5.81390 5.57711 4.96294 4.80850 4.29207 4.15303 3.71842 3.59305
40,000 40,253 41,000 41,266 42,000 42,279 43,000 43,293 44,000 44,307	39,750 40,000 40,737 41,000 41,724 42,000 42,711 43,000 43,698 44,000	1.6548 <sup>-5</sup> 1.6586 1.6698 1.6737 1.6846 1.6888 1.6994 1.7037 1.7141 1.7186	.924788 .926907 .933137 .935352 .941440 .943751 .949698 .952107 .957910 .960420	4.1343 <sup>-3</sup> 4.2938 4.7978 4.9901 5.5581 5.7895 6.4280 6.7055 7.4218 7.7538	2.83025 <sup>+2</sup> 2.93946 3.28448 3.41615 3.80500 3.96342 4.40051 4.59050 5.08083 5.30816	3.8764 <sup>-2</sup> 3.7407 3.3694 3.2470 2.9355 2.8229 2.5579 2.4581 2.2339 2.1436	3.22677 <sup>-3</sup> 3.11376 2.80475 2.70282 2.44185 2.34979 2.12925 2.04611 1.85951 1.78437
45,000 45,321 46,000 46,335 47,350 48,000 48,365 49,381	44,684 45,000 45,670 46,000 45,655 47,000 47,640 48,625 49,000	1.7287 <sup>-5</sup> 1.733 <sup>4</sup> 1.7435 1.7481 1.7577 1.7628 1.7628 1.7628 1.7628	.966078 .968689 .974202 .976916 .982282 .985101 .985101 .985101 .985101	8:5554 <sup>-3</sup> 8:9516 9:8466 1:0318 <sup>-2</sup> 1:1315 1:1875 1:2830 1:3400 1:4452 1:5122	5.85692 <sup>+2</sup> 6.12811 6.74082 7.06356 7.74600 8.12930 8.78339 9.17372 9.89320 1.03523 <sup>+3</sup>	1.9538 <sup>-2</sup> 1.8722 1.7113 1.6375 1.5011 1.4343 1.3272 1.2706 1.1779 1.1256	1.62635 <sup>-3</sup> 1.55843 1.42454 1.36310 1.24957 1.19394 1.10480 1.05767 9.80501
50,000 50,396 51,000 51,412 52,000 52,429 53,000 53,446 54,000 54,463	49,610 50,000 50,594 51,000 51,578 52,000 52,562 53,000 53,545 54,000		.985101 .985101 .985101 .985101 .985101 .985101 .985101 .985101 .979305 .974452	1.6279 <sup>-2</sup> 1.7065 1.8335 1.9257 2.0651 2.1731 2.3258 2.4524 2.5850 2.7020	1.76967	7.8251 7.3101 6.9320 6.5364	8.70218 <sup>-14</sup> 8.30024 7.72366 7.35293 6.85545 6.51373 6.08505 5.77031 5.148100

	<del></del>			3 mv	13	NEMATIC		CCCSLTY		SPECIFIC	WEIGHT
ALTI	rude		Viscos	111					_	kg	,
			kg_	/	١,	m <sup>2</sup>		η/η <sub>ο</sub>	ω,	m <sup>2</sup> sec <sup>2</sup>	യ/യ <sub>ഠ</sub>
Z,m	H,m'	μ,	т вес	$\mu/\mu_{o}$		n sec		· ··	<u> </u>	III Beic	<u></u>
	<u></u>	<u>!</u>				2	٠,	.94798+3	ς,	.8726 <sup>-3</sup>	4.88839-4
55,000	54,528	1.7	7336 <sup>-5</sup>	.968799	2	.8455-2	7	.04076	5	.5754	4.64105
55,480	55,000	1.	(247	.963733		.9810 .1362	2	14700	5	.2684	4.38549
56,000	55,511	1.	7147	.958227		.1502 5.293 <sup>1</sup> 1	2	25459	4	.9885	4.15252
56,498	56,000	1.	7052	952941		5.4611	2	.36941	4	.7194	3.92849
57,000	56.493	1.0	6956	947588		5.6435	2	49429	4	.4563	3.70949
57,516	57.000	1.	6858	94207		5.8248	2	2.61839	4	.2211	3.51368
58,000	57,476	1.	6765	.93688		.0367	7	2.76346		.9743	3.30824
58,534	58,000	1.	6662	.93113		1.2325	:	2.89752	3	5.76 <u>9</u> 3	3.13769 2.94538
59,000	58.457	1.	6572	.92010	•	+.4790		3.06622	3	5.5384	
59,553	59,000	1.	6465	,5/2011		-	_	т.		(06-3	2.79736-4
	_		6578-5	.91525	9 1	4.6904-2	2	3.21097 <sup>+</sup>	•	3.3606 <sup>-3</sup>	2.61772
60,000	59,439	) <u> </u>	6266	.90902		4.9772		3.40734		3.1447	2.48982
60,572	60,000	) I	6183	90434	1	5.2053		3.56345	- 3	2.9911	2.32281
61,000	60,420		6066	.89754	L .	5.5397		3.79241		2.7898 2.6576	2.21222
61,591	61,000		.5986	.89335	51	5.7855		3.96064		2.4703	2.05634
62,000			5865	.88660	)3	6.1758		4.22786		2.3571	1.96206
62,611			.5788	.88228	37	6.4403		4.40893		2.3833	1.81737
65,000			.5662	.8752	71	6.8964		4.72115		2.0867	1.73703
63,631 64,000	63,36	-	.5589	.8711	48	7.1805		4.91569 5.28116		1.9257	1.60296
64,65			.5458	.8638	57	7.7144		5.20110			
04,07.	<u>.</u>					8.0193	-2	5.48985	۶-	1.8439	3 1.53485 <sup>-4</sup>
65,00	0 64,34	2 1	5388 <sup>-5</sup>	.8599	22	8.6448		5.91811		1.6950	1.41096
65,67	2 65,00	XO I	5252	.8525 .8486	10	8.9709		6.14136		1.6261	1.35360
66,00	0 65,39	22 ]	.5186	.8407	140 175	9.7052		6.64401		1.4888	1.23932
66,69	2 66,00	00 ]	L 5045	.8372	60	1.0053	-1	6.88216		1.4312	1.19134 1.08618
67,00	0 66,3		1.4982	.8291		1.0916		7.47317		1.3049	1.00010
67,71	4 67,0		1.4836 1.4777	.8258		1.1286	•	7.72637		1.2570	1.04632 9.49796
68,00	0 67,2		1.4626	.817	347	1.2303	j	8.42238		1.1410	9.16995
68,7	35 68,0		1.4571	.8142	286	1.2694	1	8.69026		1.1016	-4 8.28589
69,00	00 68,2 57 69,0		1.4414	.805	499	1.3891	l	9.51143	,	9.9541	•
69,7	51 09,0						1	0 20201	+3	9.6330	-4 8.01867 <sup>-5</sup>
70,0	00 69,2	38	1.4363	·5 .802	671	1.4305	2 -	9.79321 1.07639	+4	8.6527	
70,7		000	1.4200	•795	560	1.5723		1.1058	7	8.4042	6.99574
71,0	00 70,2	216	1.4154	.790	972	1.615		1.2207	3	7.5200	6.25967
71,8		000	1.3985	.781	.528	1.783		1.2512		7.3145	6.08868
72,0	00 71.	194	1.3943	•779	TQQ	1.827		1.3876	Ś	6.5109	5.41972
72,8	325 72,0	900	1.3768	.769	3402			1.4187	8	6.3504	5.28619
73,0	ر 72 000	171	1.3731	.767	22 BV			1.5810	7	5.6218	4.67967
73,8	348 73,	000	1.3549		TOO			1.6123	ż	5.4993	5 4.57769
74,0	000 73.	148	1.3517		1861	2.637	8	1.8058	2	4.840	4.02931
74,8	372 74,	000	1.3329	• 1-4.	,501	>1					

$\begin{array}{c ccccccccccccccccccccccccccccccccccc$	ALTITUDE	ij	VISCO	SITY	KINEMATIC	VISCOSITY	SPECIFI	CWEIGHT
75,895 75,000 1.311 .73244 3.020 2.0677 4.156 2.4596 76,000 75,102 1.311 .73244 3.074 2.1047 4.083 3.3984 77,000 76,078 1.311 .73244 3.593 2.4595 3.493 2.9072 2.076,078 1.311 .73244 3.593 2.4595 3.493 2.9072 2.0700 76,078 1.311 .73244 3.642 2.4533 3.445 2.8678 77,944 77,000 1.311 .73244 4.274 2.9257 2.935 2.4432 2.9257 2.935 2.4432 2.9257 2.935 2.4432 2.9257 2.935 2.4432 2.9257 2.935 2.4432 2.9257 2.935 2.4432 2.9257 2.935 2.4432 2.9257 2.935 2.4432 2.9257 2.935 2.4432 2.9257 2.935 2.4432 2.9257 2.935 2.4432 2.9257 2.935 2.4432 2.9257 2.935 2.4432 2.9257 2.935 2.4432 2.9257 2.935 2.4432 2.9257 2.935 2.4432 2.9257 2.935 2.4432 2.935 2.454 2.9257 2.935 2.4432 2.9257 2.935 2.4432 2.9257 2.935 2.4432 2.925 2.907 2.4201 2.459 2.900 1.311 .73244 5.084 3.4801 2.467 2.0533 79,000 79,000 1.311 .73244 5.110 3.4985 2.454 2.071 4.1438 4.1297 2.073 1.7256 2.000 2.000 1.311 .73244 7.195 4.9242 1.742 1.4502 2.025 2.000 2.000 1.311 .73244 7.195 4.9242 1.742 1.4502 2.025 2.000 2.000 1.311 .73244 7.195 4.9242 1.742 1.4502 2.025 2.000 2.311 .73244 1.006 6.3337 1.245 1.0367 2.282 2.000 1.311 .73244 1.006 6.3337 1.245 1.0367 2.3244 1.006 6.3337 1.245 1.034 2.038 2.325 2.32		.	. kg			η/η <sub>ο</sub>	ω, kg m <sup>2</sup> sec <sup>2</sup>	<sup>ພຸພ</sup> ວ
92,000 91,000 92,322 91,000 93,000 91,659 93,351 92,000 94,000 92,630 1.749 1.4557	75,000 74,1 75,895 75,0 76,000 75,1 76,920 76,0 77,944 77,0 78,969 78,0 79,994 79,0 80,000 79,0 81,000 79,0 81,020 80,0 82,000 81,0 82,000 81,0 82,000 82,0 84,000 82,0 84,000 82,0 84,000 82,0 84,000 85,0 85,125 84,0 85,125 84,0 86,152 85,0 86,152 85,0 87,179 86,0 87,179 86,0 88,207 87,0 88,207 87,0 89,235 88,0 90,000 88,9 91,000 89,9 91,000 90,9 92,322 91,9 93,551 92,0	25 0.02 0.00 0.00 0.00 0.00 0.00 0.00 0.	1.3301 <sup>-5</sup> 1.311	• 743309 • 73244 • 73244	2.6826 <sup>-1</sup> 3.020 3.074 3.593 3.642 4.274 4.314 5.084 5.110 6.047 6.053 <sup>-1</sup> 7.169 7.193 8.491 8.556 1.006 1.018 1.191 1.410 1.440 1.670 1.713 1.977 2.038 2.341 2.183 3.280 40 3.483	1.83648 <sup>+4</sup> 2.0677 2.1047 2.14595 2.4595 2.4935 2.9257 2.9557 2.9557 3.4985 4.1397 4.14587 4.9242 5.8126 5.8577 6.9676 8.1518 8.2881 9.8588 9.8588 1.1727 1.35533 1.3950 1.6594 1.89738 2.2457 2.3479 2.6584	4.1083 5.2.2.2.2.2.2.2.2.2.2.2.2.2.2.2.2.2.2.2	3.3984 2.9072 2.8678 2.4432 2.4201 2.0533 2.0425 1.7256 1.7256 1.4551 1.4552 1.2282 1.2282 1.0243 6.0797 1.0243 6.0797 1.0243 6.0797 1.0243 6.0797 1.0329 1.

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ALTITUDE	SPECIFIC	WEICHT	ALTI	TUDE	SPECIFIC	WEIGHT
Z,m H,m'	ω, kg m <sup>2</sup> eec <sup>2</sup>	ω/ω <sub>o</sub>	Z,m	H,m'	ω, <u>kg</u> m <sup>2</sup> sec <sup>2</sup>	ω/ω <sub>o</sub>
95,000 93,601 95,411 94,000 96,000 94,572 96,441 95,000 97,000 95,542 97,472 96,000 98,000 96,512 98,503 97,000 99,000 97,482 99,534 98,000 100,000 98,451 100,566 99,000 101,000 99,420 101,598 100,000 102,000 100,389 102,631 101,000 103,603 101,558 103,663 102,000 104,000 102,326 104,696 103,000	1.568-5 1.460 1.319 1.223 1.112 1.027-6 9.399 8.646 7.967 7.301 6.770-6 6.181 5.767 5.246 4.925 4.463 4.215 3.807 3.616 3.254 5.109-6	1.3056-6 1.2153 1.0177 9.2546-7 8.5465 7.8240 7.1972 6.6317 6.0773 5.6355-7 5.1450 4.8007 4.3667 4.0995 3.7154 3.5088 3.1686 3.0099 2.7086	115,000 115,045 116,000 116,082 117,000 117,119 115,000 118,156 119,000 120,232 121,000 120,232 121,000 121,270 122,000 123,348 124,000 124,387	112,957 113,000 113,921 114,000 114,885 115,000 115,850 116,000 116,813 117,000 118,740 119,000 119,703 120,000 120,665 121,000 121,627 122,000	7.641 <sup>-7</sup> 7.596 6.706 6.635 5.894 5.805 5.189 5.088 4.575 4.466 4.040 <sup>-7</sup> 3.926 3.573 3.457 3.164 3.605 2.692 2.491 2.380 2.215 <sup>-7</sup>	6.3606-8 6.3231 5.5819 5.5230 4.9064 4.8323 4.3194 4.2349 3.8085 3.7172 3.3629-8 3.2679 2.9738 2.8773 2.63371 2.53753 2.2404 2.0737 1.9812
105,000 103,29h 105,730 104,000 106,000 104,261 106,764 105,000 107,000 105,229 107,798 106,000 108,000 106,196 108,832 107,000 109,000 107,162 109,867 108,000 110,000 108,129 110,902 109,000 111,000 109,095 111,937 110,000 112,973 111,000 113,000 111,961 114,000 111,992 11h,000 111,992 11h,000 111,992	5.109-0 2.788 2.678 2.394 2.312 2.060 2.000 1.776 1.734 1.534 1.528 1.328 1.328 1.328 1.328 1.328 1.328 1.328 1.328 1.328 1.328 1.328 1.328 1.328 1.328	2.5877-1 2.3206 2.2294 1.9925 1.9248 1.7145 1.6651 1.4782 1.4435 1.2771 1.2534-7 1.1055 1.0905 9.5053 8.3305 8.3305 8.3305 8.32998 7.2597	125,000 125,427 126,000 126,467 127,507 128,000 128,548 129,000 130,630 131,672 132,000 135,000 137,929	122,589 123,000 123,551 124,000 124,512 125,000 125,473 126,000 126,434 127,000 127,395 128,000 128,355 129,000 129,315 130,000 132,193 135,000	2.215 2.108 1.972 1.869 1.758 1.659 1.475 1.391 1.289 1.223 <sup>-7</sup> 1.131 1.080 9.570 8.795 6.783 4.965	1.8439 1.7544 1.6416 1.5556 1.4634 1.3812 1.2280 1.1575 1.0728 1.0183-8 9.4107-9 8.2859 7.9664 7.3214 5.6462 4.1333

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ALTITUD	)E	SPECIF	IC WEIGHT	ALI	TTUDE	SPECIF	IC WEIGHT
Z,m	H,m'	ω, kg m <sup>2</sup> sec <sup>2</sup>	ω/ω <sub>ο</sub>	2,5	K,m'	ω, kg m <sup>2</sup> sec <sup>2</sup>	ω/ω <sub>ο</sub>
143,153 146 145,000 141 148,385 145 150,000 146 153,625 150 155,000 156 164,131 160 169,397 165 170,000 165 174,671 170 175,000 170 179,954 175 180,000 175 180,000 175 185,000 179 185,245 180	,000 ,311 ,000 ,043 ,768 ,000 ,486 ,000 ,196 ,000 ,827 ,000	1.951 1.929 1.562 1.525 1.260 1.215	3.3576-9 2.4908 2.1098 1.5808 1.5808 1.3863 1.0465 9.4498-10 7.1749 6.6437-10 5.0662 4.7954 3.6687 3.5407 2.7150 2.6664 2.01;79 2.0435-10 1.6257 1.6058 1.3001 1.2696 1.0114 8.5131-11 8.1151 6.9558	210,000 211,831 215,000 217,175 220,000 222,526 225,000 227,887 230,000 233,256 235,000 244,021 245,000 244,021 245,000 254,821 255,000 260,000 260,235 265,657 270,000 271,088 275,000	205,000 207,966 210,000 212,641 215,000	6.869 <sup>-10</sup> 6.402 5.678 5.235 4.718 <sup>-10</sup> 4.305 3.941 3.559 3.307 2.957 2.788 2.469 2.360 <sup>-10</sup> 2.006 1.744 1.713 1.475 1.475 1.467 1.261 <sup>-10</sup> 1.252 1.088 1.067 9.415 <sup>-11</sup> 9.127	

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ALTI	TUDE	SPECIFI	WEIGHT	ALTI:	LUDE	SPECIFIC	WEIGHT
Z, m	H,m'	ω, kg m sec	ω/ω <sub>o</sub>	Z,m	H,m'	w, kg m <sup>2</sup> sec <sup>2</sup>	ω/ω <sub>0</sub>
280,000 281,977 285,000 287,435 290,000 292,902 295,000 298,377 300,000 303,862 305,000 309,356 310,000 314,859 320,000 325,893 340,000 348,069 359,213 360,000 370,394 380,000 381,612 390,000 392,867	331,735 340,000 340,705 349,648 350,000 358,565 360,000	7.117-11 6.743 6.215 5.823 5.442 5.044 4.778 4.382 4.205-11 3.334 3.282 2.919 2.585 2.254 2.053 1.757 1.644 1.326-11 1.326-11 1.094 1.077-12 8.600 1.772-12 8.731 7.238 7.017 5.987	5.9240 <sup>-12</sup> 5.6130 5.1730 4.8475 4.5298 4.1989 3.9769 3.6475 3.5006 <sup>-12</sup> 3.5006 <sup>-12</sup> 3.1773 3.0888 2.7750 2.4298 2.1515 1.8766 1.7081 1.4625 1.3680 1.1495 1.1034 <sup>-12</sup> 9.1067 <sup>-13</sup> 8.9613 7.3251 7.2680 6.0247 5.8410 4.9838 4.7249	450,000 460,000 461,197 470,000 472,721 480,000 484,283 490,000 495,884 500,000	490,000 497,719	2.504 2.141 2.131 <sup>-12</sup> 1.622 1.789 1.563 1.564 1.264 1.163 1.069 1.009 <sup>-12</sup> 9.080 <sup>-13</sup> 8.773 7.737 7.654 6.698 6.616 5.877	4.1455-13 3.8456 3.4658 3.1450 2.9121 2.5911 2.1437 2.0839 1.7823 1.7742-13 1.5166 1.4887 1.3013 1.2490 1.1207 1.0523 9.6836-14 8.3964 7.5575 7.5031 6.3716 5.5751 5.5074 4.8919 4.7253

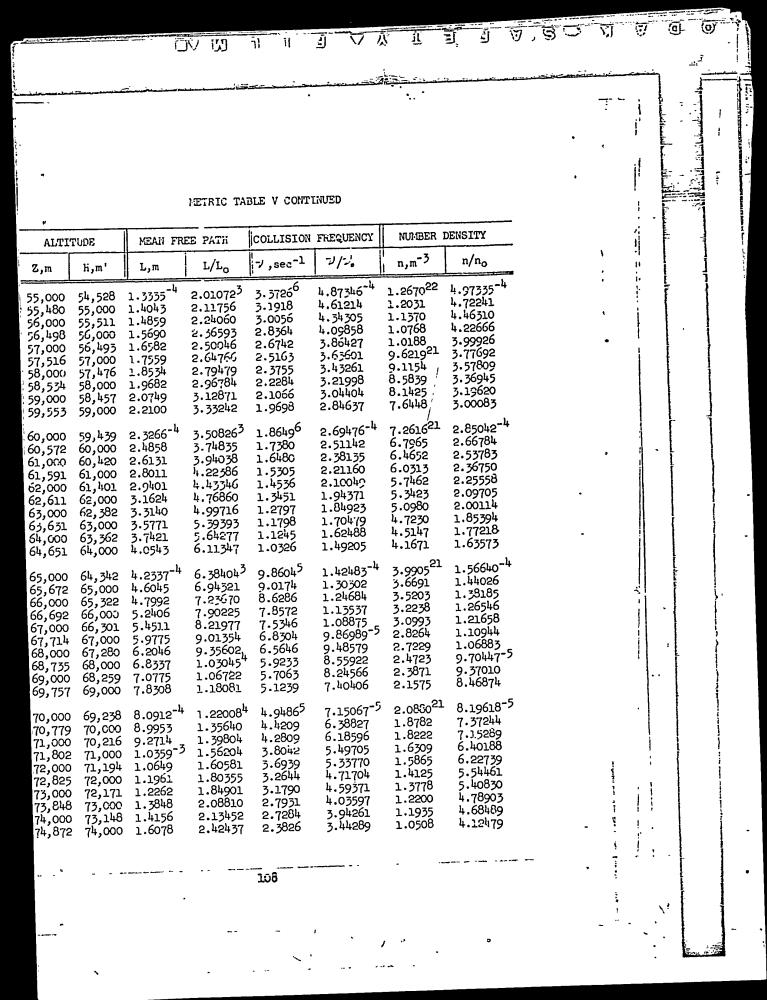
METRIC TABLE V

MEAN FREE PATH, COLLISION FREQUENCY AND NUMBER DENSITY AS FUNCTIONS OF GEOVETRIC AND GEOPOTENTIAL ALTITUDE

ALTI	TUDE	MEAN F	REE PATH	COLLISIO	FREQUENCY	NUMBER	DENSITY
Z,nı	H,m'	L,m	L/L <sub>o</sub>	マノ,sec <sup>-1</sup>	ブノン。	n,m <sup>-3</sup>	n/n <sub>o</sub>
-4,000 -3,997.5 -3,000 -2,998.6 -2,000 -1,999.4 -1,000	-5,003.9 -5,000 -4,002.5 -4,000 -3,001.4 -3,000 -2,000.6 -2,000 -1,000.2 -1,000	4.2068-8 4.2082 4.5903 4.5914 5.0181 5.0187 5.4959 5.4962 6.0310 6.0311	6.34342 <sup>-1</sup> 6.34554 6.92177 6.92545 7.56678 7.56678 8.28729 8.28776 9.09419 9.09435	1.1509 <sup>10</sup> 1.1505 1.0440 1.0457 9.45039 9.4490 8.5370 8.5364 7.6951 7.6949	1.66303 1.66241 1.50853 1.50812 1.36557 1.36538 1.23359 1.23359 1.211194 1.11192	4.0161 <sup>25</sup> 4.0147 3.6805 3.6796 3.3668 3.3663 3.0740 3.0739 2.8013 2.8012	1.57644 1.57591 1.44472 1.44437 1.32157 1.32140 1.20667 1.20660 1.09960 1.09958
0 1,000 1,000.2 2,000 2,000.6 3,000 3,001.4 4,000 4,002.5	0 999.8 1,000 1,999.4 2,000 2,998.6 3,000 3,997.5 4,000	6.6317 <sup>-8</sup> 7.3079 7.3080 8.0710 8.0715 8.9347 8.9360 9.9151 9.9178	1.00000 <sup>0</sup> 1.10196 1.10197 1.21703 1.21695 1.34727 1.34746 1.49511	6.92049 6.2089 6.2088 5.5566 5.5565 4.9599 4.9591 4.4151 4.4138	1.00000 8.97183-1 8.97171 8.02929 8.02910 7.16701 7.16587 6.37980 6.37792	2.5476 <sup>25</sup> 2.5118 2.5118 2.0933 2.0931 1.8909 1.8906 1.7059 1.7035	1.00000 9.07475-1 9.07464 8.21671 8.21622 7.42243 7.42137 6.68847 6.68671
5,000 5,003.9 6,000 6,005.7 7,000 7,007.7 8,000 8,010.7 9,000 9,012.8	4,996.1 5,000 5,994.3 6,000 6,992.3 7,000 7,989.9 8,000 8,987.3 9,000	1.1032 <sup>-7</sup> 1.1036 1.2307 1.2315 1.3769 1.3781 1.5451 1.5469 1.7394 1.7421	1.66345 <sup>0</sup> 1.66415 1.85577 1.85694 2.07623 2.07807 2.32988 2.33263 2.62282 2.62684	3.9189 <sup>9</sup> 3.9170 5.4679 5.4654 3.0590 3.0560 2.6895 2.6858 2.3560 2.3519	5.66275 <sup>-1</sup> 5.66006 5.01106 5.00755 4.42031 4.41594 3.88608 3.88097 3.40437 3.39853	1.5315 <sup>25</sup> 1.5308 1.5728 1.3719 1.2270 1.2259 1.0934 1.0921 9.7130 <sup>24</sup> 9.6981	6.01161 <sup>-1</sup> 6.00906 5.38859 5.38519 4.81643 4.81216 4.29206 4.28701 3.81270 3.80685
10,000 10,016 11,000 11,019 12,000 12,023 13,000 13,027 14,000 14,031	9,984.3 10,000 10,981 11,000 11,977 12,000 12,973 13,000 13,979 14,000	1.9646 <sup>-7</sup> 1.9685 2.2269 2.2524 2.6044 2.6137 3.0473 3.0601 3.5653 3.5828	2.96249 <sup>0</sup> 2.96827 3.35805 3.36622 3.92715 3.94118 4.51434 5.37614 5.40247	2.0562 <sup>9</sup> 2.0517 1.7875 1.7626 1.5280 1.5226 1.3059 1.3005 1.1162 1.1107	2.97121 <sup>-1</sup> 2.964/4 2.58291 2.57590 2.20798 2.20012 1.88703 1.87916 1.61288 1.60502	8.5994 <sup>24</sup> 8.5526 7.5864 7.5680 6.4870 6.4639 5.5441 5.5210 4.7385 4.7155	3.37554-1 3.36896 2.97792 2.97069 2.54637 2.53731 2.17624 2.16716 1.86001 1.85100

ALTI	TUDE	MEAN FR	EE PATH	COLLISION	FREQUENCY	NUMBER	DENSITY
Z,m	H,m'	L,m	L, L <sub>o</sub>	ア, sec-1	יר/יר.	n,m-3	n/n <sub>o</sub>
15,000 15,035 16,000 16,040 17,000 17,046 18,000 18,051 19,000 19,057	14,965 15,000 15,960 16,000 16,955 17,000 17,949 18,000 18,945 19,000	4.1714 <sup>-7</sup> 4.1947 4.8802 4.9112 5.7091 5.7500 6.6784 6.7521 7.8119 7.8820	6.29011 <sup>0</sup> 6.32521 7.35882 7.40557 8.60870 8.67046 1.00703 <sup>1</sup> 1.01514 1.17796 1.18852	9.55998 9.4870 8.1545 8.1030 6.9705 6.9209 5.9588 5.9112 5.0942 5.0489	1.57352 <sup>-1</sup> 1.57087 1.17852 1.17088 1.00724 1.00007 8.61051 <sup>-2</sup> 8.54176 7.36108 7.29565	4.0501 <sup>24</sup> 4.0276 5.4619 5.4609 2.9593 2.9582 2.5298 2.5298 2.5096 2.1627 2.1435	1.58980 <sup>-1</sup> 1.58097 1.35891 1.35033 1.16162 1.15334 9.93016 <sup>-2</sup> 9.85088 8.48925 8.41379
20,000 20,063 21,000 21,070 22,000 22,076 23,000 23,084 24,000 24,091	19,937 20,000 20,931 21,000 21,924 22,000 22,917 23,000 23,910 24,000	9.1374-7 9.2282 1.0687-6 1.0804 1.2499 1.2650 1.4618 1.4810 1.7095 1.7340	1.37783 <sup>1</sup> 1.39153 1.61151 1.62920 1.88478 1.90747 2.20426 2.23327 2.57776 2.61471	4.35528 4.3125 3.7237 3.6832 3.1838 3.1459 2.7223 2.6870 2.3279 2.2950	-6.29328 <sup>-2</sup> 6.23133 5.38070 5.32227 4.60056 4.54585 3.93378 3.88268 3.36380 3.31626	1.8490 <sup>24</sup> 1.8308 1.5808 1.5637 1.3516 1.3356 1.1557 1.1407 9.8828 <sup>23</sup> 9.7431	7.25779 <sup>-2</sup> 7.18634 6.20534 6.13797 5.30565 5.24255 4.53667 4.47774 3.87934 3.82451
25,000 25,099 26,000 26,107 27,000 27,115 28,000 28,124 29,000 29,153	24,902 25,000 25,894 26,000 26,886 27,000 27,877 28,000 28,868 29,000	1.9991-6 2.0302 2.3645 2.4073 2.7939 2.8477 3.2939 3.3613 3.8749 3.9588	3.01439 <sup>1</sup> 3.06131 3.56537 3.62990 4.21299 4.29412 4.96693 5.06852 5.84294 5.96943	1.9907 <sup>8</sup> 1.9602 1.6934 1.6645 1.4428 1.4166 1.2320 1.2083 1.0542	2.87655 <sup>-2</sup> 2.83247 2.44703 2.40527 2.08687 2.04706 1.78019 1.74594 1.52325 1.49226	8.4513 <sup>23</sup> 8.3218 7.1453 7.0182 6.0469 5.9327 5.1290 5.0262 4.3601 4.2679	3.31742 <sup>-2</sup> 3.26658 2.80476 2.75490 2.37361 2.32877 2.01332 1.97296 1.71147 1.67520
30,000 30,142 31,000 31,152 32,000 32,162 33,000 33,172 34,000 34,183	29,859 30,000 30,850 31,000 31,810 32,000 32,830 33,000 33,819 34,000	4.5484-6 4.6525 5.3279 5.4565 6.2282 6.3865 7.2663 7.4602 8.4608 8.6976	6.85855 <sup>1</sup> 7.01556 8.03395 8.22785 9.39160 9.63017 1.09569 <sup>2</sup> 1.12493 1.27580 1.31152	9.03887 8.8446 7.7658 7.5901 6.6852 6.5262 5.7659 5.6220 4.9824 4.8521	1.30611 <sup>-2</sup> 1.27804 1.12216 1.09677 9.66010 <sup>-3</sup> 9.43033 8.33173 8.32381 7 19959 7.01132	3.7144 <sup>23</sup> 3.6313 3.1710 3.0963 2.7126 2.6454 2.3251 2.2646 1.9968 1.9424	1.45803 <sup>-2</sup> 1.42540 1.24472 1.21538 1.06478 1.03840 9.12666-3 8.88944 7.83821 7.62473

ALTT	TIME	MEAN ER	EE PATH	LOGITIE	FREQUENCY	NUMBER	DENCIAN
<del></del>		I PEAU FR	EE PAIN	<del>'</del>	<del> </del>		DENSITY
Z,m	H,m'	L,m	L, L <sub>Q</sub>	' ₩,scc <sup>-1</sup>	:/:	n,m-3	n/n <sub>o</sub>
35,000	34,808	9.8570-6	1.4827.22	4.31327		1.718223	6.74457-3
35,194	35,000	1.0121-5	1.52621	4.1952	6.00203	1.6692	6.55217
36,000	35,797	1.1407	1.72002	3.7405	5.40500	1.4811	5.81390
36,205 37,000	36,000 36,786	1.1757 1.3208	1.77279	3.6336 3.005	5.25049	1.4370	5.64082
37,217	36,786 37,000	1.5652	1.99168 2.05551	3.2495	4.69547 4.55545	1 2791	5.02089
58,000	37,777	1.5268	2.50227	3.1526 2.8276	4:08585	1.2594	4.86497
38,229	38,000	1.5778	2.50227	2.7398	3.95902	1.1065 1.0708	4.34354
39,000	<i>7</i> 8, 762	1.7618	2.65661	2.4646	3.56139	9.589522	4.20313 3.76419
39,241	39,000	1.8231	2.74911	2.3850	5.44631	9.2668	3.63753
					•	•	_
40,000	39,750	2.0296 <sup>-5</sup>	3.06044 <sup>2</sup>	2.1517 <sup>7</sup>	3.10917 <sup>-3</sup>	8.3241 <sup>22</sup>	3.26751 <sup>-3</sup>
40,253	40,000	2.1031	3.17126	2.0795	3.00483	8.0332	3.15332
41,000	40,737	2.3342	3.51982	1.3814	2.71869	7.2377	2.84105
41,266	41,000	2.4221	3.65228	1.8159	2.62401	6.9753	2.73803
42,000	41,724	2.6803	4.04168	1.6477	2.38090	6.3032	2.47422
42,279	42,000	2.7851	4.19964	1.5882	2.29489	6.0661	2.38115
43,000	42,711	3.0729	4.63359	1.4451	2.08824	5.4980	2.15815
43,293	43,000	3.1974	4.82141	1.3911	2.01011	5.2838	2.07408
1:4,000	43,698	3.5175	5.30407	1.2694	1.83425	4.8030	1.88534
44,307	44,000	3.6653	5.52691	1.2202	1.76322	4.6094	1.80933
45,000	44,684	4.0205-5	6.06258 <sup>2</sup>	1.11667	1.61344-3	4.202122	1.64946 <sup>-3</sup>
45,321	45,000	4.1954	6.32619	1.0719 9.8355 <sup>6</sup>	1.54887	4.0270	1.58073
46,000	45,670	4.5887	6.91932	9.83556	1.42123	3.6818	1.44523
46, 335	46,000	4.7950	7.23046	9.4290	1.36249	3.5234	1.38304
47,000	46,655	5.2296	7.88572	8.6758	1.25365	3.2306	1.26812
47,350	47,000	5.4727	8.25225	8. <i>3</i> 057	1.20017	3.0871	1.21179
48,000	47,640	5.9130	8.91623	7.6872	1.11080	2.8572	1.12155
48,365	48,000	6.1758	9.31247	7.3601	1.06353	2.7356	1.07383
49,000	48,625	6.6605	1.004343	6.8244	9.86127 <sup>-4</sup>	2.5365	9.95675
49,381	49,000	6.9692	1.05089	6.5221	9.42449	2.4242	9.51574
50,000	49,610	7.5023 <sup>-5</sup>	1.13127 <sup>3</sup>	6.05876	8.75484-4	2.2519 <sup>22</sup>	8.83961 <sup>-1</sup>
50,396	50,000	7.8646	1.18591	5.7796	8.35151	2.1482	8.43237
51,000	50,594	8.4501	1.27420	5.3791	7.77283	1.9993	7.84809
51,412	51,000	8.8750	1.33827	5.1216	7.40070	1.9036	7.47235
52,000	51,578	9.5173 1	1.43512	4.7759	6.90125	1.7752	6.96807
52,429	52,000	1.0015-4	1.51020	4.5385	6.55813	1.6869	6.62162
53,000	52,562	1.0719	1.61631	4.2406	6.12761	1.5761	6.18694
53,446	53,000	1.1302	1.70423	4.0218	5.81148	1.4948	5.86775
54,000	53,545	1.1984	1.80707	3.7786	5.46012	1.4098	5.53383
54,463	54,000	1.2589	1.89826	3.5857	5.18136	1.3420	5.26800



ALTI	TUDE	MEAN FR	EE PATH	COLLISION	FREQUENCY	NUMBER	DENSITY
Z,m	H,m'	L,m :	L/L <sub>o</sub>	sec <sup>-1</sup> ,	ن <i>-ا</i> لا	n,m <sup>-3</sup>	n/n <sub>o</sub>
75,000 75,895 76,000 76,920 77,000 77,944 78,000 78,969	74,125 75,000 75,102 76,000 76,078 77,000 77,055 78,000	1.6385 <sup>-3</sup> 1.872 1.906 2.227 2.257 2.649 2.674 3.151	2.47068 <sup>14</sup> 2.8250 2.8755 5.5580 3.4040 5.9944 4.0524 4.7514	2.5551 <sup>5</sup> 2.026 1.991 1.703 1.680 1.452 1.419 1.204	5.57424-5 2.9279 2.8764 2.4614 2.4281 2.0692 2.0498 1.7396	1.0311 <sup>21</sup> 9.024 20 8.866 7.587 7.484 6.378 6.318 5.362	4.04747 <sup>-5</sup> 3.5423 3.4801 2.9780 2.9377 2.5035 2.4799 2.1046
79,000 79,994	78,030 79,000	3.168 3.748	4.7764 5.6519	1.198	1.7304 1.4624	5.354 4.507	2.0936 1.7693
80,000 81,000 81,020 82,000 82,045 33,000 83,072 84,000 84,098	79,006 79,981 80,000 80,956 81,000 81,930 82,000 82,904 83,000	3.752 <sup>-3</sup> 4.444 4.459 5.263 5.304 6.233 6.309 7.381 7.504	5.6575 <sup>4</sup> 6.7007 6.7250 7.9359 7.9972 9.3983 9.5128 1.11205 1.1316	1.0115 8.536* 6.508 7.208 7.152 6.086 6.013 5.139 5.055	1.4610 <sup>-5</sup> 1.2335 1.2294 1.0415 1.0335 8.7945 <sup>-6</sup> 8.6887 7.4265 7.5044	4.503 <sup>20</sup> 3.802 3.789 3.210 3.186 2.711 2.678 2.289 2.251	1.7676 <sup>-5</sup> 1.4924 1.4874 1.2601 1.2504 1.0640 1.0512 8.9851 <sup>-6</sup> 8.8373
85,000 65,125 86,000 86,152 87,000 87,179 88,000 88,207 89,000 89,255	83,878 84,000 84,852 85,000 85,825 86,000 86,798 87,000 87,771 88,000	8.740 <sup>-3</sup> 3.926 1.035 <sup>-2</sup> 1.062 1.225 1.263 1.451 1.502 1.718 1.787	1.3179 <sup>5</sup> 1.3460 1.5605 1.6011 1.8477 1.9046 2.1876 2.2655 2.5899 2.6949	4.540 4.250 5.665 3.572 3.096 3.003 2.615 2.525 2.209 2.123	6.2716-6 6.1406 5.2965 5.1623 4.4733 4.3398 3.7783 3.6484 3.1914 3.0671	1.933 <sup>20</sup> 1.893 1.632 1.591 1.379 1.328 1.165 1.124 9.836 <sup>19</sup> 9.453	7.5878-6 7.4293 6.4081 6.2456 5.4121 5.2506 4.5712 4.4140 3.8611 3.7108
90,000 90,264 91,000 91,293 92,000 92,322 93,000 93,351 94,000 94,381	88,744 89,000 89,716 90,000 90,688 91,000 91,659 92,000 92,630 93,000	2.033 <sup>-2</sup> 2.126 2.407 2.529 2.849 3.007 3.369 3.572 3.978 4.237	3.0660 <sup>5</sup> 3.2056 3.6295 3.8131 4.2956 4.5340 5.0795 5.3859 5.9991 6.3891	1.866 <sup>4</sup> 1.78 <sup>1</sup> 1.576 1.500 1.3 <sup>1</sup> 0 1.275 1.143 1.080 9.755 <sup>3</sup> 9.188	2.6958-6 2.5784 2.2773 2.1676 1.9359 1.8391 1.6510 1.5617 1.4096 1.3277	8. 309 <sup>19</sup> 7.947 7.019 6.681 5.931 5.619 5.015 4.730 4.247 3.987	3.261;-6 3.1196 2.7552 2.6225 2.3280 2.2055 1.9687 1.8567 1.6669 1.5652

				,			
ALTI	TUDE	mean fi	REE PATH	COLLISION	FREQUENCY	NUMBER	DENSITY
Z,m	H,m'	L,m	ı,/L <sub>o</sub>	γ,sec <sup>-l</sup>	v/v <sub>o</sub>	n,m <sup>-3</sup>	n/n <sub>o</sub>
95,000 95,411 96,000 96,441 97,000 97,472 98,000 98,503 99,000 99,534	93,601 94,000 94,572 95,000 95,542 96,000 96,512 97,000 97,482 98,000	4.692 <sup>-2</sup> 5.018 5.524 5.934 6.493 7.003 7.617 8.249 8.920 9.698	7.0748 <sup>5</sup> 7.5674 8.3298 8.9474 9.79056 1.0560 1.1486 1.2439 1.3451 1.4624	8.340 <sup>+3</sup> 7.823 7.141 6.671 6.123 5.698 5.260 4.876 4.526 4.180	1.2051-6 1.1304 1.0318 9.6396-7 8.8483 8.2341 7.6010 7.0461 6.5407 6.0404	3.601 <sup>19</sup> 3.366 3.058 2.847 2.602 2.413 2.218 2.048 1.894 1.742	1.4135-6 1.3215 1.2005 1.1176 1.0214 9.4699-7 8.7060 8.0393 7.4344 6.8380
100,000 100,566 101,000 101,598 102,000 102,631 103,000 103,663 104,000 104,696	98,451 99,000 99,420 100,000 100,389 101,000 101,358 102,000 102,326 103,000	1.043 <sup>-1</sup> 1.138 1.216 1.333 1.416 1.558 1.646 1.817 1.910 2.115	1.5722 <sup>6</sup> 1.7160 1.8342 2.0095 2.1357 2.3487 2.4822 2.7397 2.8795 3.1897	4.902 <sup>+3</sup> 3.590 3.370 3.089 2.915 2.663 2.526 2.300 2.193 1.990	5.6382 <sup>-7</sup> 5.1878 4.8690 4.4637 4.2123 3.8479 5.6506 3.3231 3.1693 2.8752	1.620 <sup>19</sup> 1.485 1.389 1.268 1.193 1.085 1.026 9.299 8.847 7.987	6.3605 <sup>-7</sup> 5.8276 5.4520 4.9763 4.6822 4.2577 4.0286 3.6500 3.4728 3.1351
105,000 105,730 106,000 106,764 107,000 107,798 108,000 108,832 109,000 109,867	103,294 104,000 104,261 105,000 105,229 106,000 106,196 107,000 107,162 108,000	2.211 <sup>-1</sup> 2.458 2.556 2.851 2.949 3.300 3.396 3.814 3.904 4.400	3.3342 <sup>6</sup> 3.7065 3.8537 4.2990 4.4461 4.9768 5.1205 5.7512 5.8869 6.6341	1.908 <sup>+3</sup> 1.725 1.662 1.497 1.450 1.303 1.268 1.135 1.110 9.906 <sup>+2</sup>	2.7564 <sup>-7</sup> 2.4921 2.4014 2.1639 2.0956 1.8822 1.8319 1.6400 1.6040 1.4314	7.641 <sup>18</sup> 6.873 6.611 5.926 5.730 5.119 4.975 4.430 4.327 3.840	2.9992 <sup>-7</sup> 2.6979 2.5949 2.3261 2.2492 2.0093 1.9529 1.7388 1.6987 1.5074
110,000 110,902 111,000 111,937 112,000 112,973 113,000 114,000	108,129 109,000 109,095 110,000 110,061 111,000 111,026 111,992 112,000	4.481 <sup>-1</sup> 5.066 5.134 5.824 5.873 6.683 6.707 7.648 7.656	6.7565 <sup>6</sup> 7.6391 7.7416 8.7814 8.8556 1.0077 1.0114 1.1532 1.1545	9.735 <sup>+2</sup> 8.661 8.551 7.584 7.523 6.652 6.629 5.850 5.844	1.4067 <sup>-7</sup> 3.251 <sup>4</sup> 1.2357 1.0959 1.0871 9.6120 <sup>-8</sup> 9.5790 8.4531 8.4439	3.771 5.335 3.291 2.901 2.877 2.528 2.519 2.209 2.207	1.4801 <sup>-7</sup> 1.3090 1.2917 1.1388 1.1292 9.9234 9.8876 8.6715 8.6616

ALT	ITUDE	MEAN F	REE PATH	COLLISION	FREQUENCY	2	DENSITY
Z,m	H,m'	L,m	L/L <sub>o</sub>	γ,sec <sup>-l</sup>	v/v <sub>o</sub>	n,m <sup>-3</sup>	r/n <sub>o</sub>
115,000 115,045 116,000 116,082 117,000 117,119 118,000 118,156 119,000 119,194	112,957 113,000 113,921 114,000 114,885 115,000 115,850 116,000 116,813 117,000	8.707 <sup>-1</sup> 8.758 9.898 1.000 1.123 1.140 1.273 1.298 1.441 1.476	1.3129 <sup>7</sup> 1.3206 1.4925 1.5081 1.6941 1.7196 1.9201 1.9577 2.1731 2.2255	5.17c <sup>+2</sup> 5.141 4.576 4.530 4.055 3.998 3.599 3.533 3.199 3.127	7.4705-8 7.4293 6.6117 6.5463 5.8598 5.7769 5.2007 5.1053 4.6220 4.5182	1.940 <sup>18</sup> 1.929 1.707 1.689 1.50 <sup>1</sup> 1.481 1.327 1.301 1.172 1.145	7.6167-8 7.5725 6.7003 6.6308 5.9030 5.8154 5.2081 5.1080 4.6017 4.4933
120,000 120,232 121,000 121,270 122,000 122,309 123,000 123,348 124,000 124,387	117,777 118,000 118,740 119,000 119,703 120,000 120,665 121,000 121,627 122,000	1.6290 1.675 1.838 1.899 2.072 2.150 2.332 2.430 2.622 2.743	2.4561 <sup>7</sup> 2.5263 2.7721 2.8636 3.1245 3.2413 3.5171 3.6639 3.9539 4.1359	2.846 <sup>+2</sup> 2.771 2.536 2.459 2.263 2.185 2.022 1.944 1.808 1.732	4.1131-8 4.0041 3.6651 3.5534 3.2700 3.1577 2.9212 2.8096 2.6129 2.5032	1.037 <sup>18</sup> 1.008 <sub>1</sub> 7 9.190 <sup>17</sup> 8.896 8.153 7.860 7.243 6.953 6.443 6.160	4.0715 <sup>-8</sup> 3.9584 3.6074 3.4921 3.2005 3.0852 2.8432 2.7293 2.5291 2.4178
125,000 125,427 126,000 126,467 127,000 127,507 128,000 128,548 129,000 129,569	122,589 123,000 123,551 124,000 124,512 125,000 125,473 126,000 126,454 127,000	2.944° 3.092 3.302 3.481 3.698 3.915 4.137 4.597 4.662 5.025	4.43947 4.6627 4.9784 5.2497 5.5759 5.9032 6.2376 6.6298 7.0291 7.5770	1.619 <sup>+2</sup> 1.545 1.452 1.380 1 303 1.254 1.171 1.105 1.049 9.816 <sup>+1</sup>	2.3398-8 2.2329 2.0978 1.9943 1.8830 1.7834 1.6921 1.5966 1.5160 1.4185	5.738 <sup>1</sup> 7 5.464 5.117 4.853 4.569 4.316 4.084 3.843 3.624 3.624	2.2525-8 2.1447 2.0087 1.9049 1.7934 1.6940 1.6032 1.5083 1.4227 1.3198
130,000 130,630 131,000 131,672 132,000 132,714 135,000 137,929	127,395 128,000 128,355 129,000 129,315 130,000 132,193 135,000	5.291 <sup>0</sup> 5.720 5.984 6.488 6.745 7.332 9.481 1.291 <sup>+1</sup>	7.9784 <sup>7</sup> 8.6255 9.0233 9.7827 <sub>8</sub> 1.0171 1.1056 1.4296 1.9462	9.378 <sup>+1</sup> 8.752 8.409 7.828 7.563 7.024 5.594 4.256	1.3551-8 1.2646 1.2151 1.1312 1.0929 1.0149 8.0832-9 6.1502	3.193 <sup>17</sup> 2.953 2.823 2.604 2.505 2.304 1.782 1.309	1.2534-8 1.1593 1.1082 1.0222 9.8323-9 9.0446 6.9951 5.1361

ALTITU	DE:	MEAN FR	EE PATH	COLLISION	FREQUENCY	NUMBER	DENSITY
Z,m /	H,m'	L,m	L/L <sub>o</sub>	y ,5ec <sup>-1</sup>	v/v <sub>o</sub>	n,m-3	n/n <sub>o</sub>
140,000 1; 143,153 1; 145,000 1; 148,385 1; 150,000 1; 153,625 1; 155,000 1; 158,874 1; 160,000 1; 164,131 1; 165,000 1;	36,983 40,000 41,766 45,000 46,542 50,000 51,311 55,000 56,072 60,000 60,826	1.585 <sup>+1</sup> 2.130 2.510 3.340 3.803 5.022 5.555 7.295 7.872 <sup>+1</sup> 1.029 <sup>+2</sup> 1.087	2.390 <sup>+8</sup> 3.212 3.785 5.036 5.734 7.573 8.377 1.100 <sup>+9</sup> 1.187 <sup>+9</sup> 1.552 1.639	3.548 <sup>+1</sup> 2.731 2.361 1.823 1.634 1.277 1.168 9.175 8.577 6.763 6.451	5.126 <sup>-9</sup> 3.946 3.412 2.634 2.361 1.846 1.688 1.326 1.239 <sup>-9</sup> 9.773 <sup>-10</sup> 9.322	1.066 <sup>17</sup> 7.932 <sup>16</sup> 6.731 5.059 4.443 3.364 3.041 2.316 2.146 <sup>16</sup> 1.642 1.555	4.1834 <sup>-9</sup> 3.1135 2.6421 1.9858 1.7439 1.3205 1.1937 9.0908-10 8.4250-10 6.4436 6.1029 4.6827
169,397 1 170,000 1 174,671 1 175,000 1	65,000 65,572 70,000 70,311 75,000	1.416 1.467 1.907 1.942 2.521	2.136 2.212 2.876 2.928 3.801	5.097 4.941 3.915 3.854 3.058	7.365 7.140 5.658 5.569 4.419	1.193 1.152 8.858 <sup>15</sup> 8.701 6.703	4.5212 3.4772 3.4155 2.6311
185,000 1 185,245 1 190,000 1 190,545 1 195,000 1	.75,043 .79,768 .80,000 .84,486 .85,000 .89,196 .90,000	2.525 <sup>+2</sup> 3.119 3.151 3.826 3.910 4.664 4.822	5.808 <sup>+9</sup> 4.703 4.751 5.769 5.896 7.033 7.271	3.053 <sup>0</sup> 2.505 2.490 2.082 2.040 1.73 <sup>4</sup> 1.682	4.411 <sup>-10</sup> 3.619 3.598 3.008 2.948 2.506 2.430	6.690 <sup>15</sup> 5.417 5.362 4.416 4.320 3.622 3.504	2.6259 <sup>-10</sup> 2.1264 2.1049 1.7334 1.6959 1.4218 1.3754
200,000 1 201,171 1 205,000 1 206,497 2 210,000 2 211,831 2 215,000 2	193,899 195,000 198,595 200,000 203,284 205,000 207,966 210,000	8.169 8.720 9.747 <sub>+3</sub>	8.524 <sup>+9</sup> 8.909 1.027+10 1.085 1.232	1.453 <sup>0</sup> 1.395 1.223 1.162 1.034 9.740 <sup>-1</sup> 8.789 8.199	2.099.10 2.015 1.767 1.680 1.495 1.407 1.270 1.185	2.989 <sup>15</sup> 2.859 2.480 2.347 2.068 1.937 1.733	1.1731-10 1.1224 9.7332-11 9.2134 8.1177 7.6051 6.8040 6.3103
220,000 2 222,526 2 225,000 2 227,887 2 230,000 2 233,256 2	212,641 215,000 217,308 220,000 221,969 225,000 226,622 230,000	1.157 <sup>+3</sup> 1.260 1.368 1.504 1.611 1.788 1.889	1.745 <sup>+10</sup> 1.900 2.063 2.268 2.429 2.695 2.848 3.189	7.501 <sup>-1</sup> 6.934 6.427 5.890 5.529 5.023 4.774 4.500	1.084 <sup>-10</sup> 1.002 9.288 <sup>-11</sup> 8.511 7.990 7.258 6.899 6.214	1.460 <sup>15</sup> 1.341 1.235 1.123 1.049 9.45214 8.945 7.988	4.8465 4.4086

ALTI	PIDE	MEAN F	REE PATH	COLLISION	N FREQUENCY	NUMBER I	DENSITY
Z,m	H,m'	L,m	L/L <sub>o</sub>	v,sec-1	v/v <sub>o</sub>	n,m-3	n/n <sub>o</sub>
240,000 244,021 245,000 249,417 250,000 254,821 255,000	231,268 235,000 235,908 240,000 240,540 245,000 245,165	2.206 <sup>+3</sup> 2.493 2.567 2.926 2.977 3.423 3.440	3.3263 <sup>+10</sup> 3.7586 3.8708 4.4126 4.4886 5.1612 5.1877	4.137 <sup>-1</sup> 3.696 3.597 3.187 3.137 2.758 2.745	5.9776 <sup>-11</sup> 5.3400 5.1971 4.6051 4.5330 3.9848 3.9650	7.659 <sup>14</sup> 6.778 6.581 5.773 5.676 4.936	3.006 <sup>-11</sup> 2.661 2.583 2.266 2.226 1.938 1.928
260,000 260,235 265,000 265,657 270,000 271,088 275,000 276,528	249,784 250,000 254,395 255,000 258,999 260,000 263,597 265,000	3.963 <sup>+3</sup> 3.990 4.552 4.635 5.213 5.367 5.952 6.195	5.9765 <sup>+10</sup> 6.0158 5.8641 6.9885 7.8603 8.0925 8.9757 9.3421	2.408 <sup>-1</sup> 2.394 2.119 2.084 1.870 1.820 1.654 1.594	3.4802 <sup>-11</sup> 3.4591 3.0625 3.0121 2.7022 2.6306 2.3906 2.3039	4.263 <sup>14</sup> 4.235 3.711 3.645 3.241 3.148 2.838 2.727	1.673 <sup>-11</sup> 1.662 1.457 1.431 1.272 1.236 1.114 1.070
280,000 281,977 285,000 287,435 290,000 292,902 295,000 298,377	268,187 270,000 272,771 275,000 277,347 280,000 281,917 285,000	6.779 <sup>+3</sup> 7.131 7.700 8.185 8.724 9.369 9.860 1.070 <sup>+4</sup>		1.467 <sup>-1</sup> 1.400 1.304 1.233 1.162 1.086 1.038 9.628	2.1202 <sup>-11</sup> 2.0232 1.8848 1.7814 1.6795 1.5692 1.4999	2.492 <sup>14</sup> 2.369 2.194 2.064 1.937 1.803 1.713	9.783 <sup>-12</sup> 9.300 8.613 8.103 7.602 7.079 6.726 6.200
300,000 303,862 305,000 309,356 310,000 314,859	286,480 290,000 291,036 295,000 295,585	1.112 <sup>+4</sup> 1.218 1.251 1.384 1.404		9.289 <sup>-2</sup> 8.538 8.331 7.589 7.486 6.760	1.3423 <sup>-11</sup> 1.2338 1.2038 1.0966 1.0817 9.7678 <sup>-12</sup>	1.520 <sup>14</sup> 1.387 1.351 1.221 1.203 1.077	5.444 5.301 4.792 4.722 4.228
320,000 325,893 330,000 336,963 340,000 348,069	304,663 310,000 313,714 320,000 322,738	1.759 <sup>+4</sup> 2.002 2.187 2.533 2.698	2.6527 <sup>+11</sup> 3.0182 3.2971 3.8194 4.0682 4.7959	6.079 <sup>-2</sup> 5.396 4.973 4.341 4.095 3.517	8.7848 <sup>-12</sup> 7.7972 7.1856 6.2724 5.9171 5.0822	9.603 <sup>13</sup> 8.441 7.727 6.670 6.262 5.312	3.770 <sup>-12</sup> 3.313 3.033 2.618 2.458 2.085

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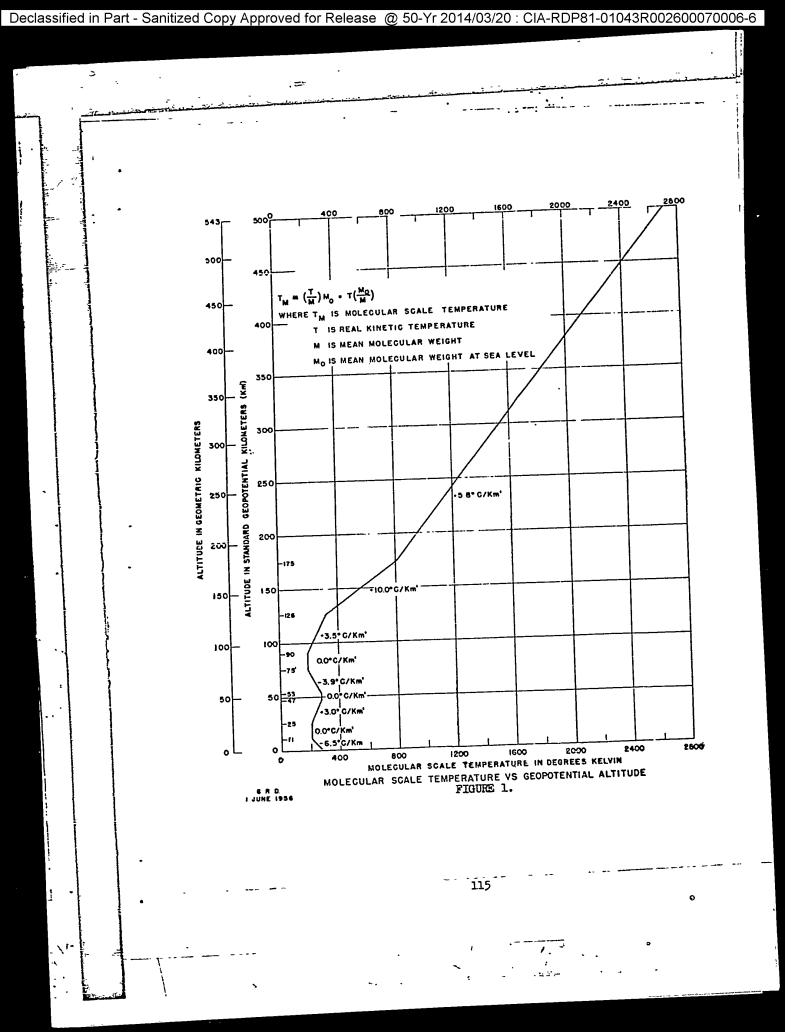
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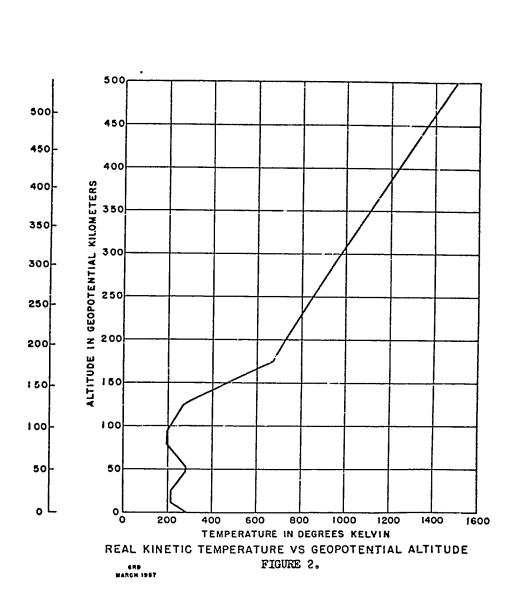
# METRIC TABLE V CONFINUED

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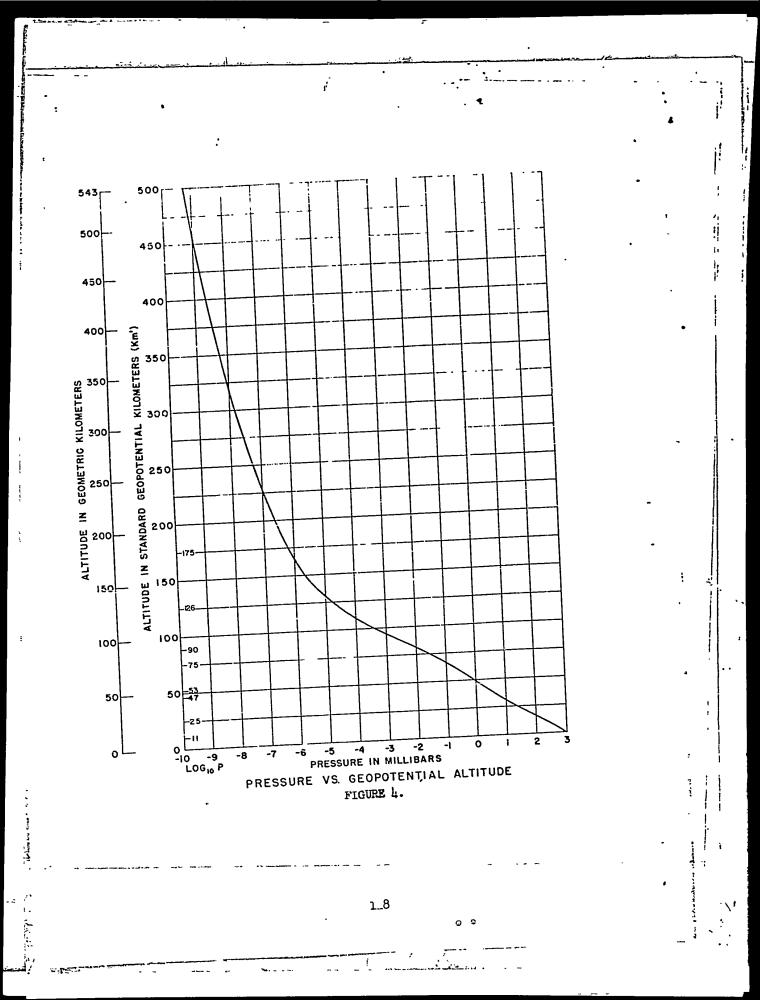
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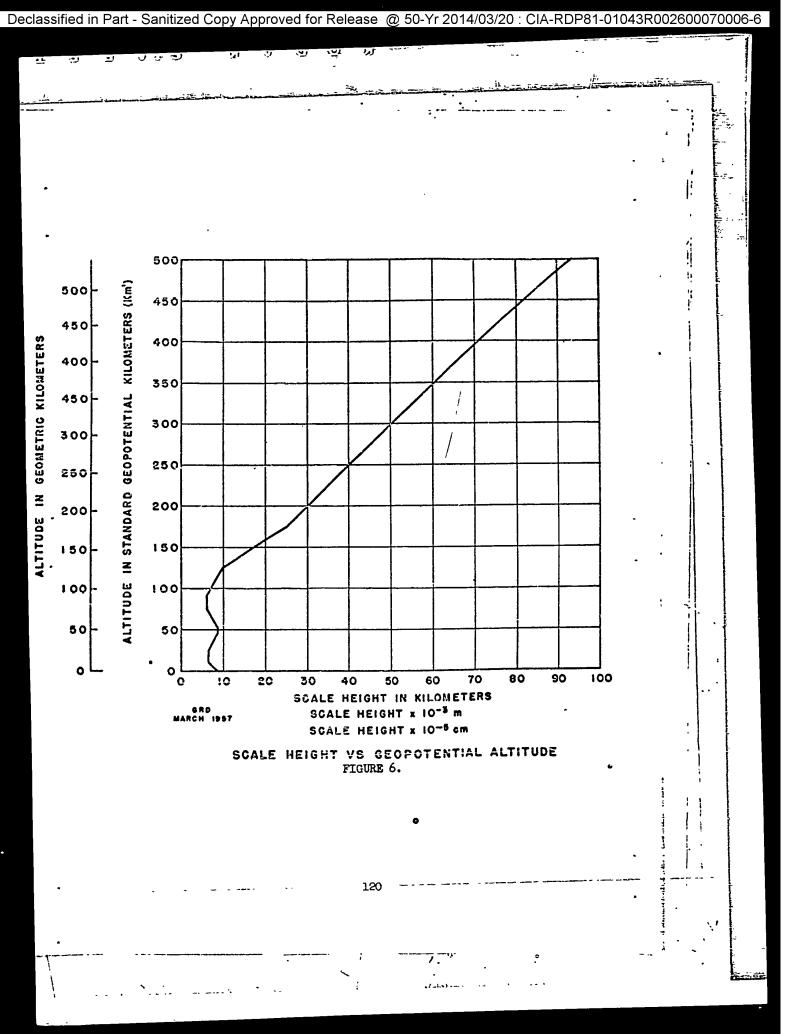
				COLLISION FREQUENCY		NUMBER DENSITY	
ALTITUDE		MEAN FREE PATH			v/v <sub>0</sub>	n,m-3	n/n <sub>o</sub>
Z,m	H,m'	L,m	L/L <sub>o</sub>	y,sec <sup>-l</sup>	V/V <sub>0</sub>	<u> </u>	
350,000 359,213 360,000 370,000	331,735 340,000 340,705 349,648	3.306 <sup>+4</sup> 3.964 4.026 4.872 4.908	4.9853 <sup>+11</sup> 5.9781 6.0701 7.3463 7.4009	3.393 <sup>-2</sup> 2.869 2.829 2.371 2.355	4.9035 <sup>-12</sup> 4.1456 4.0875 3.4264 3.4030	5.110 <sup>13</sup> 4.261 4.197 3.468 3.442 2.582	2.0059 <sup>-12</sup> 1.6728 1.6474 1.3612 1.3512 1.1312
370, 594 380,000 381,612 390,000 392,867	350,000 358,565 360,000 367,456 370,000	5.863 6.037 7.017 7.381	8.8404 9.1035 1.0581+12 1.1130	1.615	2.8874 2.8101 2.4453 2.3335	2.798 2.408 2.289 2,022 <sup>13</sup>	1.0985 9.4507-13 8.9846 7.9363 <sup>-13</sup>
400,000 404,160 410,000 415,491 420,000	576,320 380,000 385,158 390,000 393,970	8.356 <sup>+1</sup> 8.973 9.903 1.085 <sup>+5</sup> 1.168	1.2600 <sup>+12</sup> 1.3531 1.4933 1.6360 1.7616	1.438 <sup>-2</sup> 1.348 1.231 1.131 1.056 9.533 <sup>-3</sup>	2.0776 <sup>-12</sup> 1.9480 1.7782 1.6342 1.5262 1.3775	1.883 1.706 1.557 1.446 1.294	7.3906 6.6967 6.1125 5.6767 5.0812
426,860 430,000 438,267 440,000 449,713	400,000 402,756 410,000 411,516	1.305 1.372 1.562 1.605	1.9680 2.0691 2.3561 2.4202 2.8076	8.071 7.873 6.863	1.1663 1.1377 9.9168-13	1.231 1.081 1.053 9.074	
450,000 460,000 461,197 470,000 472,721 480,000 484,283 490,000	420,250 428,959 430,000 437,642 140,000 146,300 346,300 454,930	1.870 <sup>+5</sup> 2.171 2.209 2.510 2.610 2.894 3.072 2.3.324	2.8198 <sup>+1</sup> : 3.2731 3.3312 3.7856 3.9360 4.3634 4.6323 5.0129 5.4314	2 6.835 <sup>-3</sup> 5.955 5.859 5.205 5.021 4.570 4.318 4.014 3.727	8.4660 7.5215 7.2551 6.6032 6.2402 5.8006 5.3859	7.648 6.739 6.472 5.838 5.499 5.082 4.690	3.0552 3.0019 2.6416 2.5406 2.2918 2.1587 1.9948 1.8412
500,00 507,52 510,00 519,20 520,00 530,00 540,00 542,60	0 463,54 5 470,00 0 472,12 15 480,00 0 480,67 00 489,21 490,00 497,71	0 3.807 <sup>†</sup> 0 4.208 12 4.347 10 4.899 19 4.949 12 5.620 10 5.685 19 6.365	5 5.7411 <sup>+1</sup> 6.3453 6.5553 7.3876 7.4633 8.4737 8.5728 9.5976 9.9168	2 3.541 3.228 3.132 2.804 2.778 2.470 2.443 2.201 2.135	5.1165 <sup>-13</sup> 4.6640 4.5256 4.0518 4.0137 3.5688 3.5306 3.1803 3.0855	4.437 <sup>1</sup> 4.037 4.086 5.448 5.413 3.006 2.972 2.654 2.569	2 1.7418 <sup>-13</sup> 1.5760 1.5255 1.3536 1.3399 1.1801 1.1665 1.0419 1.0084

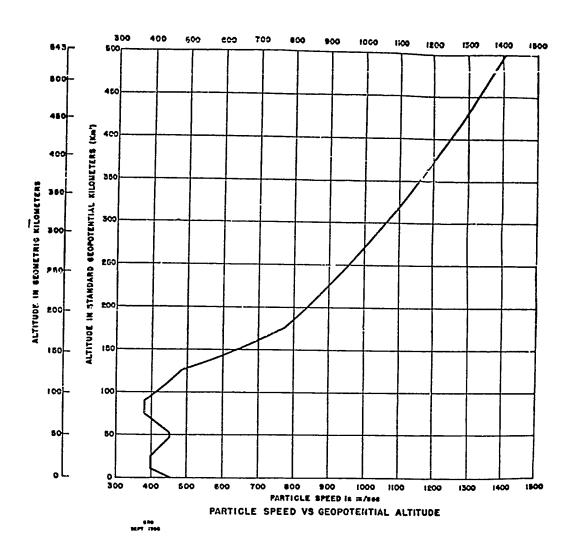


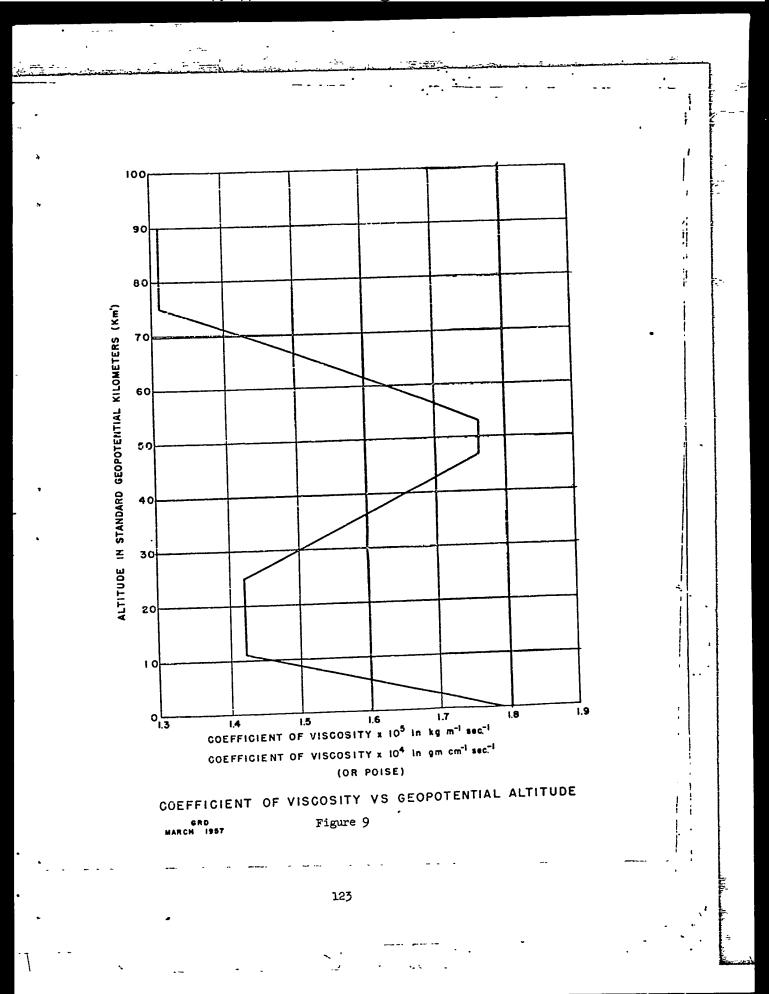


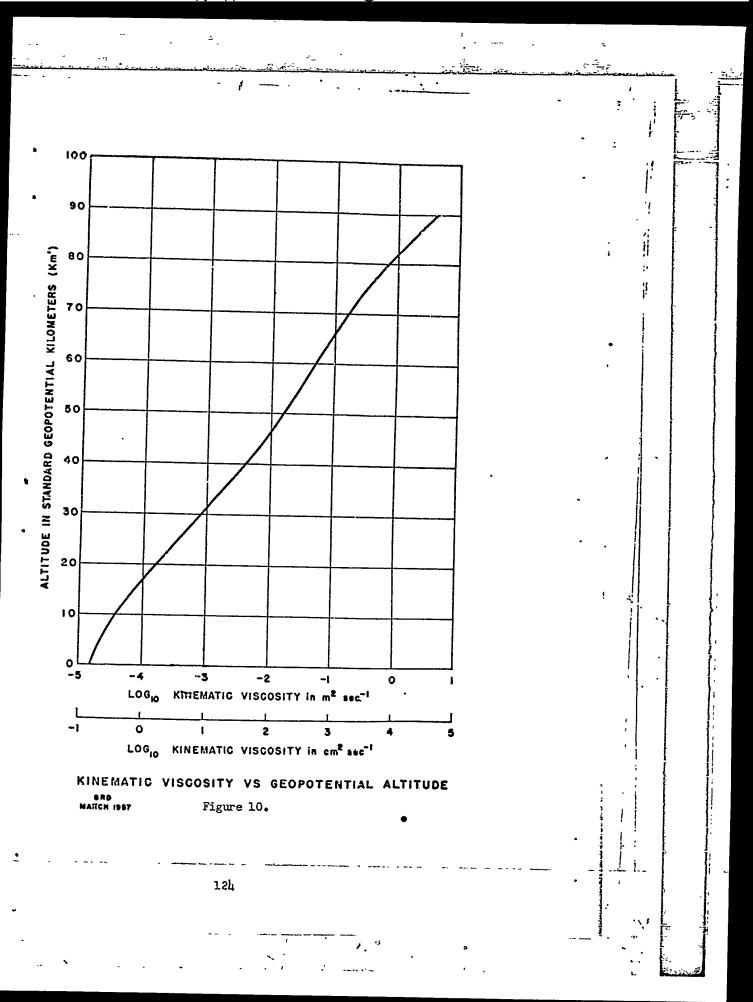
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ENGLISH TABLE I
TEMPERATURES, MOLECULAR WEIGHT, GRAVITATIONAL ACCELERATION
AS FUNCTIONS OF GEGMETRIC AND GEOPOTENTIAL ALTITUDE

ALTI	TUDE		TEMPE	RATURE		MOLECUL	AR WEIGHT	GRAVT. ACCEL.
Z,rt	H,ft'	t,°C	t, °F	T, °R	T/To	М	M/M <sub>o</sub>	g,ft sec-2
-15000 -14989 -12500 -12493 -10000 -9995.2 -7500 -7497.3 -5000 -4998.8 -2500 -2499.7  0 2500 2500.3 5000 5001.2 7500 7502.7 10000 10005 12500 12508 15000 12508 15000 17515 20000 20019 22500 22524 25000 25030 27500	-15011 -15000 -12508 -12500 -10005 -10000 -7502.7 -7500 -5001.2 -5000 -2500.3 -2500 0 2499.7 2500 4998.8 5000 7497.3 7500 9995.2 10000 12493 12500 14989 15000 17485 17500 19981 20000 22476 22500 24970 25000 27464	+, °C  44.668  7568  7768  7869  44.668  7869  44.668  7869	112.531 112.492 103.604 103.577 94.662 85.756 85.756 85.756 876.833 67.916 67.915 59.000 50.085 41.174 41.169 32.254 23.338 14.452 23.338 14.452 23.358 -3.408 -12.352 -21.239 -30.047 -38.940	572.218 563.227 563.227 563.227 563.227 563.527 515	1.10320 1.10313 1.08599 1.08594 1.06879 1.06879 1.05158 1.05157 1.03439 1.01719 1.00000 .982814 .982812 .965632 .965623 .948453 .948453 .931279 .931247 .914110 .914058 .896870 .879782 .879681 .862625 .862493 .845371 .845305 .828322 .828116 .811177 .810928	constant at 28.966 for altitudes to 299,516 feet.	1.00000 for altitudes to 299,516 feet 0.0000	32.2204 32.2204 32.2204 32.2127 32.2126 32.2049 32.1972 32.1972 32.1895 32.1818 32.1663 32.1663 32.1586 32.1586 32.1586 32.1586 32.1599 32.1509 32.1201 32.12124 32.12124 32.1201 32.1201 32.124 32.1047 32.1047 32.0970 32.0893
27536	27500	274.42	277		ą.	-		

#### ENGLISH TABLE I CONTINUED

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ALT	TUDE		TEMPER	RATURE		MOLECUI	LAR WELGHT	GRAVT. ACCEL.
Z,ft	H,ft'	t,°C	t,°F	T, R	T/T <sub>o</sub>	М	м/мо	g,ft sec-2
30000 30043 32500 32551 35000 35059 36152 37500 37568	29957 30000 32449 32500 34941 35000 36089 37433 37500	-44.351 -44.436 -49.289 -49.389 -54.226 -54.342 -56.500 -56.500	-47.831 -47.985 -56.720 -56.900 -65.607 -65.816 -69.700 -69.700	411.86 411.70 402.97 402.79 394.08 393.87 389.99 389.99 389.99	.794036 .793740 .776899 .776551 .759766 .759363 .751874 .751874	28.966	1.00000	32.0817 32.0816 32.0740 32.0739 32.0663 32.0662 32.0628 32.0587 32.0585
40000 40077 42500 42587 45000 45097 47500 47608 50000 52500 52500 52500 52500 55145 57500 57659 60000 60173 70000 70236 80308 82345 90000 90390	39923 40000 42414 42500 44903 45000 47392 47500 49880 50000 52368 52500 54855 55000 57342 57500 59828 60000 69766 70000 79694 80000 82021 89613 90000	-56.500 -56.500	-69.700 -69.700	589.99 589.99	.751874 .751874	constant at 28.966 for altitudes to 299,516 feet	1.00000 for altitudes to 299,516 feet	32.0510 32.0508 32.0433 32.0433 32:0357 32.0254 32.0203 32.0207 32.0203 31.9392 31.9592 31.9286 31.9277 31.9215 31.8982 31.8970 31.8677
100482 110000 110583	100000 109423 110000	-40.060 -31.444 -30.916	-40.108 -24.599 -23.649	419.58 435.09 436.04	.808926 .838827 .840658	28.966	1.00000	31.8663 31.8373 31.8356

#### ENGLISH TABLE I CONTINUED

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	TITUDE		<del></del>	PATURE	·	MOLECU	LAR WEIGHT	GRAVT. ACCEL.
7,ft	H,ft'	t,°C	t,°F	т,°R	T/T <sub>o</sub>	М	M/M <sub>O</sub>	g,ft sec-2
12000 12069 13000 13081	5 120000 0 129195	-21.772 -13.364	-7.190 7.944	452.50 467.63	.870212 .872390 .901567		1.00000	31.8070 31.8049 31.7767
140000 140946	139066 140000	-4.338 -3.484	24.192 25.729	483.88 485.42	.904123 .932893 .935855	į		31.7742 31.7464 31.7435
150000 151087 155346 160000	150000 154199 158782	5.660 9.500 9.500	40.425 42.188 49.100 49.100	501.88 508.79 508.79	.964188 .967587 .980913 .980913	#	•	31.7162 31.7129 31.7000 31.6860
161237 170000 171397 175346 180000	168626 170000 173885	9.500 9.500 9.500 9.500	49.100 49.100 49.100 49.100	508.79 508.79 508.79 508.79	.980913 .980913 .980913 .980913	299,516 feet	feet	31.6823 31.6559 31.6517 31.6398
181567 190000 191747	180000	4.061 2.230 -7.618 -9.657	39.310 36.015 18.288 14.618	499.00 495.70 477.98 474.31	.962040 .955686 .921510 .914434	ţ	299,516 1	31.6258 31.6211 31.5957 31.5905
200000 201937 210000	198100 200000 207907	-19.297 -21.544 -30.943	-2.735 -6.779 -23.697	456.95 452.91 435.99	.880978 .873182 .840565	altitudes	to	31.5657 31.5599 31.5358
212136 220000 222345 230000	210000 217704 220000 227491	-33.431 -42.589 -45.318 -54.223	-28.176 -44.659 -49.573 -65.602	431.51 415.03 410.11 394.09	.831929 .800151 .790677 .759775	28.966 for	r altítudes	31.5294 31.5059 31.4988
5,43001 5,45,43,4 5,4000 5,352,62	230000 237270 240000 246063	-57.206 -65.847 -69.093 -76.300	-70.970 -86.525 -92.367 -105.340	388.72 373.16 367.32 354.35	.749425 .719437 .708173 .683162	constant at 28.	1.0000 for	31.4760 31.4683 31.4461 31.4378 31.4193
250000 253033 260000 263282	247039 250000 256799 260000	-76.300 -76.300 -76.300 -76.300	-105.340 -105.340 -105.340	354.35 354.35 354.35	.683162 .683162 .683162	conste		31.4164 31.4073 31.3866
270000 273541 280000 283810	266549 270000 276291 280000	-76.300 -76.300 -76.300	-105.340 -105.340 -105.340 -105.340	354.35 354.35 354.35 354.35	.683162 .683162 .683162	. •		31.3768 31.3569 31.3464 31.3272
290000 294089 299516	286023 290000 295276	-76.300 -76.300	-105.340	354.35 354.35 354.35 354.35	.683162 .683162 .683162 .683162	28.966 1		31.3159 31.2976 31.2855 31.2694

#### ENGLISH TABLE I CONTINUED

ALTITU	me.		темог	ERATURE		MOLECITA	AR WEIGHT	GRAVT.
	H,ft'	t,°C	t, °F	T, R	T/T <sub>o</sub>	м	M/M <sub>O</sub>	g,ft sec-2
500000 2: 504378 3: 525000 3: 550000 3: 555974 3: 575000 3:	295746 20000 20013 25000 44223 50000 68376 75000	-76.30 -75.84 -67.84 -65.00 -52.52 -48.45 -34.90 -29.85	-105.3 -104.5 -90.12 -85.01 -62.53 -55.21 -30.81 -21.73	354.4 355.2 369.6 375.7 397.2 404.5 428.9 438.0	.68318 .68475 .71251 .72237 .76569 .77982 .82685 .84436	28.89 28.31 26.64 26.38 25.66 25.50 25.11 25.00	.99748 .97730 .91967 .91069 .88583 .88040 .86691	31.2680 31.2551 31.1943 31.1791 31.1207 31.1032 31.0475 31.0274
407822 40 421745 41 425000 41 433841 42 450000 44 459924 45 475000 46	92473 00000 13386 16512 25000 40495 50000 64422 75000	-16.25 -10.29 0.42 8.18 29.22 67.55 91.02 126.6 152.6	02.76 13.48 22.75 46.72 84.60 153.6 195.8 259.8 306.7	462.4 473.2 492.4 506.4 544.3 655.5 719.5 766.4	.89157 .91224 .94940 .97633 1.0494 1.1824 1.2638 1.3872 1.4776	24.76 24.68 24.54 24.52 24.45 24.25 24.20 24.15	.85481 .85188 .84736 .84642 .84403 .84025 .83823 .83553 .83377	30.9745 30.9517 30.9112 30.9018 30.8761 30.8293 30.8006 30.7571 30.7252
512282 50 550000 53 564897 55	88293 00000 35868 50000 74147	185.3 214.1 302.1 336.7 395.9	365.6 417.3 575.7 638.1 744.5	825.2 877.0 1035 1098 1204	1.5910 1.6908 1.9963 2.1164 2.3217	24.09 24.05 23.93 23.90 23.84	.83180 .63025 .82627 .82498 .82303	30.6851 30.6498 30.5419 30.4995 30.4270
617773 60 650000 63	83221 00000 30354 50000	402.0 413.6 434.9 448.9	755.7 776.5 814.9 840.0	1215 1236 1275 1300	2.3432 2.3833 2.4572 2.5058	23.59 23.17 22.48 22.09	.81458 .79991 .77621 .76252	30.3997 30.3494 30.2585 30.1998
724311 70 750000 72	77268 00000 23965 50000	468.5 485.1 502.8 522.1	875.4 905.3 937.0 971.7	1335 1365 1397 1431	2.5741 2.6316 2.6927 2.7597	21.59 21.22 20.85 20.50	.74534 .73242 .71998 .70766	30.1183 30.0505 29.9791 29.9016
831911 80 850000 81	70446 00000 16714 50000	537•3 559•5 572•1 597•3	999.2 1039 1062 1107	1459 1499 1521 1567	2.8126 2.8896 2.9333 3.0209	20.24 19.90 19.72 19.39	.69876 .68694 .68075 .66935	29.8408 29.7531 29.7035 29.6049
940590 90 950000 90	08611 00000	607.1 635.5 642.1 673.9	1125 1176 1188 1245	1584 1636 1647 1705	3.0546 3.1533 3.1762 3.2866	19.27 18.95 18.88 18.57	.66527 .65422 .65183 .64108	29.5671 29.4571 29.4317 29.3097

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#### ENGLISH TABLE I CONTINUED

ALTI	TUDE		TEMPA	PATURE		MOLECUL	AR WEIGHT	GRAVT.
Z,ft	H,ft'	t,°C	t,°F	T, R	T/T <sub>o</sub>	М	M/M <sub>C</sub>	g,ft sec-2
1000000 1050364 1100000 1161249 1200000 1273262 1500000 1586421 1400000	954245 1000000 1044889 1100000 1134710 1200000 1223721 1300000 1311932	677.2 712.5 747.3 790.2 817.3 868.4 887.1 947.0 956.4	1251 1315 1377 1454 1503 1595 1629 1737	1711 1774 1837 1914 1963 2055 2088 2196 2213	3.2980 3.4206 3.5415 3.6903 3.7844 3.9617 4.0263 4.2343 4.2669	18.54 16.24 17.97 17.68 17.51 17.23 17.14 16.86 16.82	.64004 .62955 .62034 .61028 .60454 .59481 .59158 .58213	29.2972 29.1626 29.0309 28.8696 28.7682 28.5781 28.5091 28.2880 28.2535
1500000 1500743 1600000 1616246 1700000 1732949 1780465 1850870	1399354 1400000 1485997 1500000 1571872 1600000 1640420 1700000	1025 1026 1094 1105 1162 1184 1216 1263	1878 1879 2001 2021 2123 2163 2221 2306	2337 2338 2461 2480 2583 2623 2681 2766	4.5061 4.5079 4.7438 4.7822 4.9797 5.0571 5.1683 5.0457	16.56 16.56 16.33 16.29 16.13 16.07 15.99 15.88	.57160 .57153 .56373 .56255 .55690 .55484 .55203 .54815	28.0013 27.9994 27.7525 27.7124 27.5069 27.4267 27.3117 27.1426

ENGLISH TABLE II

PRESSURE AND DENSITY AS FUNCTIONS OF GEOMETRIC AND CPOPOTENTIAL ALTITUDE

ALT	TUDE		PRES	SURE		DENS	SITY
Z,ft	H,ft'	P,mb	P,in Hg	P, lbf.	P/P <sub>o</sub>	ρ, lbfsec <sup>2</sup> ft <sup>4</sup>	Ρ/Ρ <sub>ο</sub>
-15000 -14989 -12500 -12493 -10000 -9995.2 -7500 -7497.3 -5000 -4998.8 -2500	-15011 -15000 -12508 -12500 -10005 -10000 -7502.7 -7500 -5001.2 -5000	1.3197 1.2103 1.2102	5.0140 <sup>1</sup> 5.0122 4.6163 4.6151 4.2446 4.2459 3.6976 3.8972 3.5740 3.5738 3.2726 3.2725	3.5462 <sup>3</sup> 3.5450 3.2649 3.2641 3.0020 3.0015 2.7566 2.7564 2.5276 2.3146 2.3145	1.67573 <sup>0</sup> 1.67514 1.54281 1.54242 1.41858 1.41835 1.30261 1.30249 1.19446 1.19441 1.09372	3.6105 <sup>-3</sup> 3.6094 3.3767 3.3761 3.1548 3.1544 2.9443 2.9441 2.7448 2.7447 2.5558	1.51897 1.51853 1.42064 1.42036 1.32728 1.32711 1.23871 1.23862 1.15475 1.15471 1.07524 1.07523
-2499.7 0 2500 2500.3 5000 5001.2 7500 7502.7	-2500 0 2499•7 2500 4998•8 5000 7497•3 7500	1.01325 <sup>3</sup> 9.2501 <sup>2</sup> 9.2500 8.4311 8.4307	2.9921 <sup>1</sup> 2.7315 2.7315 2.4897 2.4896 2.2656 2.2653	2.3143 2.1162 <sup>3</sup> 1.9319 1.9319 1.7609 1.7608 1.6023 1.6022	•	2.3769 <sup>-3</sup>	1.00000 9.28873 <sup>-1</sup> 9.28865 8.61699 8.61668 7.98321 7.98254
10000 10005 12500 12508 15000 15011 17500 17515	9995.2 10000 12493 12500 14989 15000 17485 17500	6.9694 <sup>2</sup> 6.9681 6.3200 6.3181 5.7207 5.7182 5.1683 5.1652	2.05807 <sup>1</sup> 2.0577 1.8663 1.8657 1.6895 1.6886 1.5262	1.4556 <sup>3</sup> 1.4553 1.3200 1.3196 1.1948 1.1943 1.0794 1.0788	6.87830 <sup>-1</sup> 6.87702 6.23738 6.23552 5.64584 5.64339 5.10069 5.09762	1.7556 <sup>-2</sup> 1.7553 1.6219 1.6215 1.4962 1.4956 1.3781 1.3774	7.38586 <sup>-1</sup> 7.38474 6.82345 6.82180 6.29453 6.29232 5.79768 5.79485
20000 20019 22500 22524 25000 25030 27500 27536	19981 20000 22476 22500 24970 25000 27464 27500	4.6600 <sup>2</sup> 4.6563 4.1931 4.1888 3.7650 3.7601 3.3730 3.3676	1.3761 <sup>1</sup> 1.3750 1.2382 1.2370 1.1118 1.1103 9.9605 <sup>0</sup> 9.91':4	9.7327 <sup>2</sup> 9.7249 8.7576 8.7485 7.8633 7.8531 7.0447 7.0333	4.59909 <sup>-1</sup> 4.59540 4.13831 4.13402 3.71574 3.71089 3.32890 3.32353	1.2673 <sup>-3</sup> 1.2664 1.1634 1.1624 1.0663 1.0651 9.7544 <sup>-4</sup> 9.7416	5.33151 <sup>-1</sup> 5.32805 4.89468 4.89057 4.48586 4.48112 4.10379 4.09843

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## ENGLISH TABLE II CONTINUED

ALT	ITUDE		PR	ESSURE		DENS	SITY
Z,ft	H,ft'	P,mb	P, in Hg	P, <u>lbf</u> .	P/P <sub>o</sub>	ρ, <u>lbfsec<sup>2</sup></u>	ρ/ο.
30000 30043 32500 32551 35000 35050 36152 37500 37568	7771	5.0146 <sup>2</sup> 5.0089 2.6882 2.6819 2.5909 2.3842 2.2632 2.1217 2.1148	8.9028 <sup>0</sup> 8.8854 7.9382 7.9196 7.0602 7.0406 6.6832 6.2653 6.2450	6.2966 <sup>2</sup> 6.2843 5.6144 5.6012 4.9934 4.9795 4.7268 4.4312 4.4169	2.97541 <sup>-1</sup> 2.96958 2.65303 2.64680 2.35960 2.35303 2.23359 2.09392 2.08716	8.8927 8.1169 8.1015 7.3820 7.3653 7.0611 6.6196 6.5982	3.74720 <sup>-1</sup> 3.74126 5.41490 3.40840 3.10569 3.09859 2.97069 2.78493 2.77594
40000 40077 42500 42587 45000 45097 47500 47608	59925 40000 42414 42500 44903 45000 47392 47500	1.8823 <sup>2</sup> 1.875 <sup>4</sup> 1.6700 1.6630 1.4816 1.4748 1.3146 1.3078	5.5584 <sup>0</sup> 5.5380 4.9314 4.9110 4.3753 4.3549 3.8820 3.8619	3.9312 <sup>2</sup> 3.9168 3.4878 3.4733 5.0945 3.0801 2.7456 2.7314	1.85767 <sup>-1</sup> 1.85085 1.64813 1.64130 1.46226 1.45547 1.29739 1.29068	5.8727 <sup>-4</sup> 5.8511 5.2103 5.1887 4.6227 4.6012 4.1015 4.0803	2.47072 <sup>-1</sup> 2.46165 2.19203 2.18294 1.94483 1.93579 1.72555
50000 50120 52500 52632 55000 55145 57500 57659	49880 50000 52368 52500 54855 55000 57342 57500	1.1664 <sup>2</sup> 1.1597 1.0349 1.0284 9.1834 9.1197 8.1489 8.0872	3.4444 <sup>0</sup> 3.4246 3.0562 3.0369 2.7119 2.6931 2.4064 2.3882	2.4361 <sup>2</sup> 2.4221 2.1615 2.1479 1.9180 1.90 <sup>1</sup> +7 1.7019 1.6890	1.15115 <sup>-1</sup> 1.14455 1.02141 1.01496 9.06327 <sup>-2</sup> 9.00048 8.04231 7.98144	3.6391 <sup>-14</sup> 3.6183 3.2290 3.2086 2.8652 2.8453 2.5412 2.5232	1.53104 <sup>-1</sup> 1.52226 1.35849 1.34991 1.20542 1.19707 1.06911
60000 60173 70000 70236 80000 80308 82345 90000 90390	59827 60000 69766 70000 79694 80000 82021 89613 90000	7.2311 <sup>1</sup> 7.1716 4.4850 4.4348 2.7831 2.7425 2.4886 1.7376 1.7067	2.1354 <sup>0</sup> 2.1178 1.3244 1.3096 8.2183-1 8.0985 7.3488 5.1312	1.5103 <sup>2</sup> 1.4978 9.3672 <sup>1</sup> 9.2623 5.8125 5.7278 5.1975 3.6291 3.5644	7.13658-2 7.07778 4.42637 4.37684 2.74666 2.70659 2.45605 1.71492 1.68434	2.2561-4 2.2375 1.3993 1.3837 8.6831-5 8.5564 7.7644 5.2531 5.1513	9.49172 <sup>-2</sup> 9.41352 5.88712 5.82124 3.65308 3.59980 3.26657 2.21004 2.16721
100000 100482 110000 110583	99523 100000 109423 110000	1.1053 <sup>1</sup> 1.0820 7.1565 <sup>0</sup> 6.9810	3.2640 <sup>-1</sup> 3.1951 2.1133 2.0615	2.3085 <sup>1</sup> 2.2598 1.4947 1.4580	1.09087 <sup>-2</sup> 1.06784 7.06294 <sup>-3</sup> 6.88969	3.2114 <sup>-5</sup> 3.1377 2.0014 1.9480	1.35107 <sup>-2</sup> 1.32007 8.42003 <sup>-3</sup> 8.19559

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# ENGLISH TABLE II CONTINUED

ALT	ITUDE		PR	ESSURE		DE	ENSITY
Z,ft	H,ft'	ř,mb	P,in Hg	P, lbf ft2	P/P <sub>o</sub>	ρ, <u>lbfsec</u> 2	
120000 120695 130000 130815 140000 140946	119313 120000 129195 130000 139066 140000	4.5779 3.1474 3.0476 2.1332 2.0575	1.3909 <sup>-1</sup> 1.3518 9.2943 <sup>-2</sup> 8.9995 6.2992 6.0759	9.5611 6.5735 6.3650 4.4552 4.2972	4.64848 <sup>-1</sup> 4.51799 3.10626 3.00773 2.10527 2.03062	1.2697 <sup>-5</sup> 1.2310 8.1894 <sup>-6</sup> 7.9072 5.3640 5.1574	5.34178 <sup>-3</sup> 5.17886 3.44540 3.32668 2.25672 2.16980
181567 190000	148929 150000 154199 158782 160000 168626 170000 173885 178460 180000 188285 190000	1.4650 <sup>0</sup> 1.4074 1.2044 1.0173 9.7267 <sup>-1</sup> 7.0788 6.7293 5.8320 4.9193 4.6418 3.3740 5.1537	4.3261-2 4.1561 5.5566 3.0041 2.8723 2.0904 1.9872 1.7222 1.4527 1.5707 9.9635-3	3.0597 <sup>0</sup> 2.9395 2.5155 2.1247 2.0315 1.4784 1.4054 1.2180 1.0274 9.6947-1 7.0468 6.5866	1.44582 <sup>-3</sup> 1.38502 1.18866 1.00401 9.59953 <sup>-4</sup> 6.98625 6.64128 5.75573 4.85495 4.58115 3.32989 3.11246	3.1122 2.8803 2.4320	1.49952 <sup>-3</sup> 1.43555 1.21179 1.02355 9.78631 <sup>-4</sup> 7.12218 6.77051 5.86773 5.04652 4.79357 3.61352 3.40370
201937 210000 212136 220000 222345 230000 232565 240000 242794 249001	246063	2.2752 <sup>-1</sup> 2.1047 1.5079 1.5775 9.7927 <sup>-2</sup> 8.8224 6.2217 5.5173 3.8580 3.3599 2.452	6.7186 <sup>-3</sup> 6.2153 4.4528 4.0676 2.8918 2.6053 1.8373 1.6293 1.1393 9.9218 7.241	4.7518-1 4.3953 3.1493 2.8769 2.0452 1.8426 1.2994 1.1523 8.0576-2 7.0173 5.121		6.0583-7 5.6545 4.2082 3.8841 2.8710 2.6175 1.9210 1.7270 1.2580 1.1130-8	2.54879 <sup>-4</sup> 2.37891 1.77043 1.63408 1.20785 1.10121 8.08180 <sup>-5</sup> 7.26582 5.29239 4.68242 3.5423
253033 2 260000 2 263282 2 270000 2 273541 2 280000 2 283810 2 290000 2 294089 2	250000 256799 260000 266549 270000 76291 80000 286023	6,912 4,956 4,073 2,962 2,400	1.463 1.203 8.746 <sup>-5</sup> 7.086	1.864-2 4.158 2.902 2.450 1.733 1.444 1.035 8.506-3 6.185 5.012 3.791	6.8219 4.8909 4.0195 2.9229 2.3683	4.772 4.028 2.849 2.374 1.702 1.399 1.017 8.240 <sup>-9</sup>	3.3641 <sup>-5</sup> 2.8764 2.8764 2.0075 1.6948 1.1986 9.9857 <sup>-6</sup> 7.1592 5.8837 4.2784 3.4667

# ENGLISH TABLE II CONTINUED

	A1	LTITUDE						·
		STITODE		P	RESSURE		4	SITY
:	Z,ft	H,ft	P,mb	P, in Hg	P, 1bf f+,2	P/P <sub>o</sub>	ρ, <u>lbfsec<sup>2</sup></u> ft <sup>4</sup>	P/P <sub>o</sub> .
	30000 30437 32500 33014 35000 35597 375000 38186	8 30000 0 32001 5 32500 0 34422 4 35000 0 36837 6 37500	00 1.418 3 5.317 10 4.225 1.826 1.439 6 6.987 0 5.453	4.188 1.570 1.248 5.393-6	3.698 <sup>-1</sup> 2.962 1.110 8.824 <sup>-1</sup> 3.814 3.005 1.459 1.139	1.4000	6.065 <sup>-9</sup> 4.749 1.610 1.250 4.957-10 3.810 1.718 1.308	2.5517 <sup>-6</sup> 1.9979 6.7730 <sup>-</sup> 7 5.2568 2.0853 1.6030 7.2292 <sup>-8</sup> 5.5011
1	400000 407822 421749 425000 433841 450000 459924 475000 486071	2 400000 5 413386 0 416512 425000 140495 450000	2.257 5 1.451 2 1.314 0 1.017 5 6.661 0 5.262 2 3.785	6.664 4.285 3.879	6.096-5 4.713 3.030 2.744 2.124 1.391 1.099 7.906-6 6.327	2.8807-8 2.2270 1.4320 1.2966 1.0036 6.5740-9 5.1933 3.7358 2.9900	6.565-11 4.943 3.038 2.672 1.919 1.111 8.187-12 5.348 4.010	2.7620 <sup>-8</sup> 2.0797 1.2781 1.1240 8.0726 <sup>-9</sup> 4.6718 3.4445 2.2501 1.6872
5555	500000 512282 550000 664897 590 <sup>1</sup> 401	500000	2.052 1.051 8.560 <sup>-7</sup>	6.892 <sup>-8</sup> 6.058 3.103 2.528 1.828	4.875 <sup>-6</sup> 4.285 2.195 1.788 1.293	2.3034-9 2.0247 1.0371 8.4481-10 6.1085	2.862 <sup>-12</sup> 2.363 1.020 7.827 <sup>-13</sup>	1.2043-9 9.9417-10 4.2928 3.2930 2.1655
6	00000 17773 50000 70910	630354 650000	3.138 2.518	1.629 <sup>-8</sup> 1.324 9.266-9 7.435	1.152-6 9.366-7 6.553 5.259	5.4443-10 4.4258 3.0967 2.4849	3.531 2.325	1.8927 <sup>-10</sup> 1.4854 9.7822 <sup>-11</sup> 7.5617
7: 7: 7:	00000 24311 50000 77977	677268 700000 723965 750000	1.879-7 1.489 1.175 9.187-8	5.550 <sup>-9</sup> 4.395 3.471 2.713	3.925 <sup>-7</sup> 3.109 2.455 1.919	1.8548-10 1.4690 1.1599 9.0665-11	9.718 <sup>-14</sup> 1 7.372	5.3707 <sup>-11</sup> 4.0885 3.1014 2.3249
83 85 88	00000 51911 50000 56115	770446 800000 816714 850000	7.624 <sup>-8</sup> 5.881 5.103 3.885		1.592-7 1.228 1.066 8.114-8	7.5340-11 5.8041 5.0364 3.8342	3.280 1 2.778 1	8692-11 3798 1688 .4955-12
94 95	0000 0590 0000 5339	862768 900000 908611 950000	3.511 <sup>-8</sup> 2.637 2.473 1.834	7.788-10 7.304	5.509 5.166	3.4646 <sup>-11</sup> 2.6029 2.4410 1.8098	1.284 5	.5456-12 .4002 .0094 .5301

# ENGLISH TABLE II CONTINUED

	TUDE		PRES	SURE		PENSITY	
Z,ft	H,ft'	P,mb	P,in Hg	P, lbf ft <sup>2</sup>	P/P <sub>o</sub>	ρ, lbfsec <sup>2</sup> ft <sup>4</sup>	
1000000 1050364 1100000 1161249 1200000 1273262 1300000 1386421	1044889 1100000 1134710 1200000 1223721 1300000	1.780 <sup>-8</sup> 1.302 9.732 <sup>-9</sup> 6.934 5.655 3.925 3.455 2.336 2.202	5.256 <sup>-10</sup> 3.845 2.874 2.048 1.670 1.159 1.020 6.897 6.502	2.720 2.033 1.448 1.181 8.197-9	1.7565 <sup>-11</sup> 1.2852 9.6045 <sup>-12</sup> 6.8436 5.5814 3.8733 3.4102 2.3050 2.1730	7.022	5.4089 <sup>-12</sup> 2.3653 1.6824 1.1317 8.9162 <sup>-13</sup> 5.8154 5.0107 3.1689 2.9576
1400000 1500000 1500743 1600000 1616246 1700000 1732949 1780465	1399354 140000 1485997 150000 1571872 1600000 1640420	1.454 <sup>-9</sup> 1.450 9.899 <sup>-10</sup> 9.324 6.922 6.185 5.281	4.293 <sup>-11</sup>	3.037 <sup>-9</sup> 3.028 2.068 1.947 1.446 1.292 1.103 8.802 <sup>-10</sup>	1.4349-12 1.4306 9.7699-13 9.2025 6.8319 6.1044 5.2116 4.1591	4.71	1.8201 <sup>-13</sup> 1.8138 1.1610 1.0825 7.6404 <sup>-14</sup> 6.6976 5.5666 4.2754

Livet in madentina and after Mark 1888 1888 1888 1888 and Monachite Chiefe.

#### ENGLISH TABLE III

SOUND SPEED, VISCOSITY, AND KINEMATIC VISCOSITY AS FUNCTIONS OF CECMETRIC AND GEOPOTENTIAL AUTITUDE

ALT	ITUDE	SOUNI	SPEED	VISCO	OSITY	KINEMATIO	VISCOSITY
Z,ft	H,ft'	C <sub>s</sub> , ft sec	C <sub>s</sub> /C <sub>so</sub>	μ, lbf sec ft <sup>2</sup>	μ/μ <sub>ο</sub>	$\eta, \frac{ft^2}{sec}$	η/η <sub>ο</sub>
-15000 -14989 -12500 -12493 -10000 -9995.2 -7500 -7497.3 -5000 -4998.8 -2500 -2499.7 0 2500 2500.3 5000 5001.2 7500 7502.7 10000 10005 12500 12508 15000 15011 17500 17515	-15011 -15000 -12508 -12500 -10005 -10000 -7502.7 -7500 -5001.2 -5000 -2500.3 -2500 4998.8 5000 7497.3 7500 9995.2 10000 12492 12500 14989 15000 17485 17500	1172.6 1172.6 1172.6 1163.4 1154.2 1154.2 1154.2 1144.9 1135.5 1135.5 1126.0 1116.4 1106.8 11097.1 1097.1 1087.3 1087.3 1077.4 1067.4 1067.4 1057.4 1057.3 1047.1 1056.9	1.05034 1.05030 1.04211 1.04209 1.05382 1.05381 1.02547 1.02546 1.01705 1.01704 1.00856 1.00000 .991370 .982666 .982661 .973886 .973876 .965028 .965011 .956091 .956064 .947071 .947032 .937967 .937913	4.0298-7 4.0296 3.9820 3.9820 3.9837 3.8853 3.8852 3.8363 3.7870 3.7870 3.7870 3.7872 3.6872 3.6872 3.6872 3.6366 3.5857 3.58	1.07828 1.07822 1.06548 1.06544 1.05256 1.05256 1.03959 1.03958 1.02650 1.02649 1.01330 1.01330 1.00000 .986589 .973073 .973066 .959443 .959443 .959428 .945673 .931786 .9317861 .903762 .903679	1.1162-4 1.1164 1.1792 1.1794 1.2469 1.2471 1.3196 1.3197 1.3977 1.4818 1.4818 1.5723-4 1.6700 1.6700 1.6700 1.7756 1.7756 1.7756 1.8897 1.8898 2.0132-4 2.0135 2.1472 2.1477 2.2928 2.2934 2.4510 2.4520	.709872 <sup>-1</sup> .710040 .7149997 .750122 .793123 .839253 .839305 .888933 .888958 .942397 .942403  1.00000 <sup>0</sup> 1.06214 1.06215 1.12925 1.12928 1.20183 1.20191 1.28042 <sup>0</sup> 1.28058 1.36564 1.36591 1.45819 1.45861 1.55983
20000 20019 22500 22524 25000 25030 27500 27536	20000 22476 22500 24970 25000 27464 27500	1036.9 1036.8 1026.6 1026.5 1016.1 1016.0 1005.5	.928776 .928705 .919495 .919405 .910122 .910009 .900654	3.3245 3.3241 3.2708 3.2703 3.2167 3.2161 3.1621 3.1613	.889539 .889429 .875190 .875050 .860712 .860537 .846102 .845889	2.6247 2.8114 2.8133 3.0169 3.0194 3.2418	1.66846 <sup>0</sup> 1.66933 1.78804 1.78926 1.91872 1.92036 2.06176

#### ENGLISH TABLE III CONTINUED

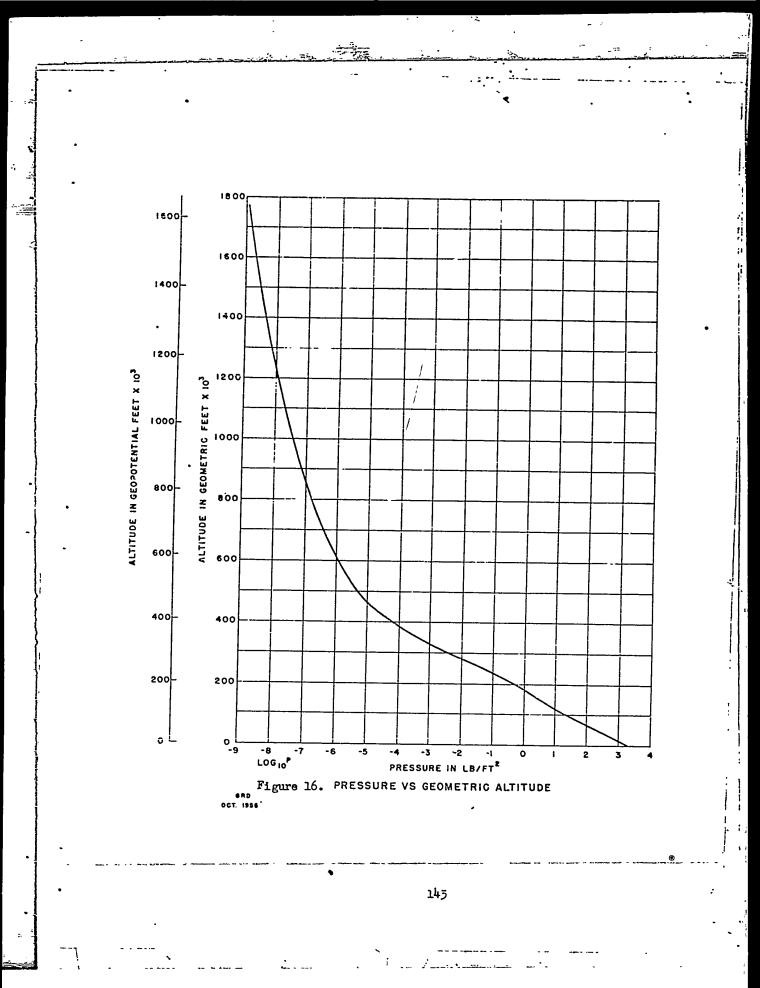
~		<del></del>	·	<del></del>		n	
ALT	TTUDE	SOLME	SPEED	VISCO	YTIE		VISCOSITY
Z,ft	H,ft	C <sub>s</sub> , ft sec	C <sub>s</sub> /C <sub>so</sub>	μ, lbf sec	μ/μ <sub>ο</sub>	$\eta, \frac{\text{ft}^2}{\text{sec}}$	η/ηο
30000 30043 32500 32551 35000 35059 36152 37500 37568	30000 32449 32500 34941 35000 36089 37433 37500	994.66 984.05 983.83 973.14 972.89 968.08 968.08 968.08	.891087 .890921 .881419 .881221 .871646 .871414 .867107 .867107	3.1061 3.0514 3.0503 2.9953 2.9940 2.9692 2.9692 2.9692	.831358 .831102 .816477 .816173 .801455 .801100 .794486 .794486	3.4884-4 3.4929 3.7593 3.7651 4.0576 4.0649 4.2051 4.4855 4.5001 5.0560-4	2.21861 <sup>0</sup> 2.22145 2.39093 2.39459 2.58060 2.58529 2.67441 2.85280 2.86204
40077 42500 42587 45000 45097 47500 47608	40000 42414 42500 44903 45000 47392 47500	968.08 968.08 968.08 968.08 968.08 968.08 968.08	.867107 .867107 .867107 .867107 .867107 .867107	2.9692 2.9692 2.9692 2.9692 2.9692 2.9692	. 794486 . 794486 . 794486 . 794486 . 794486 . 794486 . 794486	5.0746 5.6988 5.7225 6.4252 6.4532 7.2394 7.2771	3.22746 3.62443 3.63953 4.08513 4.10420 4.60426 4.62821
50000 50120 52500 52632 55000 55145 57500 57659	49880 50000 52368 52500 54855 55000 57342 57500	963.08 968.08 968.08 968.08 968.08 968.08 968.08	.867107 .867107 .867107 .867107 .867107 .867107 .867107	2.9692 <sup>-7</sup> 2.9692 2.9692 2.9692 2.9692 2.9692 2.9692 2.9692	· 794486 · 794486 · 794486 · 794486 · 794486 · 794486 · 794486 · 794486	8.1592 <sup>-14</sup> 8.2062 9.1955 9.25 <i>3</i> 9 1.0363 <sup>-3</sup> 1.0435 1.1684 1.1768	5.18921 <sup>0</sup> 5.21912 5.84830 5.88548 6.59093 6.63691 7.43128 7.48428
60000 60173 70000 70236 80000 80308 82345 90000 90390	59828 60000 69766 70000 79694 60000 82021 89613 90000	968.08 968.08 968.08 983.46	.867107 .867107 :867107 .867107 .867107 .867107 .867107 .087107 .080889 .881586	2.9692 <sup>-7</sup> 2.9692 2.9692 2.9692 2.9692 3.0484 3.0524	.794486 .794486 .794486 .794486 .794486 .815663	1.3270 2.1219 2.1459 3.4196 3.4702 3.8242 5.8030	8.37031 <sup>0</sup> 8.43984 1.34953 <sup>+1</sup> 1.36481 2.17484 2.20703 2.43217 3.69071
100000 100482 110000 110583	99523 100000 109423 110000	1004.1	.898561 .899403 .915875 .916874	3.2499	.844174 .869596	1.0055 <sup>-2</sup> ( 1.6239 :	6.23857 <sup>+1</sup> 6.39491 1.03277 <sup>+2</sup> 1.06294

#### EMBLISH TABLE III COMTINUED

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ALTITUDE	SOULD	SPEEC	' VISCOS	lty	l	VISCOSITY
Z,ft H,ft'	3 <sub>5</sub> , <u>ft</u>	0 <sub>8</sub> /0 <sub>80</sub>	μ, lbf sec	11/10	$\eta, \frac{\operatorname{ft}^2}{\operatorname{sec}}$	गम्भेठ
100000 110717	1041.5	.932551	3.3480-7	.895844	2.6569-2	1.67705+2
120000 119313 120695 120000	1075.6	.954018	3.3548	897650	2.7253	1.75550
130095 12000	1060.1	9-2509	3.4444	.921638	4.2060	2.67498
1%515 1%000	1061.6	9: 0854	3.4522	.923722	4.3659	2.716/1
14000 139066	1075.3	965864	3.5392		6.5980	4.19654
140946 140000	10,0.0	967596	3.5481	.949373	6.8796	4.37539
140340 140000	10.70.0	00000000	•	-5-5515		
150000 148989	1096.3	.981931	3.6324 <sup>-7</sup>	.971932	1.0191-1	6.48164*2
151087 150000	1095.2	983660	3.6424	.974617	1.0675	6.78916
155348 154199	1105.7	.990411	3.6816	.985101	1.2782	8.12933
160000 158782	1105.7	990411	3.6816	.985101	1.5133	9.62136
161237 160000	1105.7	.990411	3.6816	.985101	1.5827	1.00661+3
170000 168626	1105.7	.990411	3.6816	.985101	2.1748	1.38315
171397 170000	1105.7	.990411	3.6816	.985101	2.2877	1.45499
175346 173885	1105.7	.990411	3.6816	.985101	2.6397	1.67885
180000 178460	1095.1	.980836	3.6260	.970232	3.0229	1.92258
181567 180000	1091.4	977592	3.6072	.965196	3.1659	2.01352
190000 188285	1071.7	959953	3.5049	.937828	4.0807	2.59533
191747 190000	1067.6	.956260	3.4835	.932102	4.3058	2.73850
<b>2</b> /21/1				•		91.7
200000 198100	1047.9	.938604	3.3813 <sup>-7</sup>	.904748	5.5813 <sup>-1</sup>	3.54971 <sup>4-3</sup>
201937 200000	1043.3	.934442	3.3572	.898305	5.9373	3.77613
210000 207907	1023.6	.916824	3.2554	.871062	7.7360	4.92007
212136 210000	1018.3	.912102	3.2282	.863768	8.31130	5.28597
220000 217704	998.67	.894512	3.1268	.836634	1.0891	6.92666
222345 220000	992.74	.889200	3.0962	.828453	1.1829	7.52308
230000 227491	973.15	.871651	2.9953	.801463	1.5593	9.91689
232565 230000	966.50	<b>7</b> 865694	2.9611	.792318	1.7146	1.09047
240000 237270	946.96	.848196	2.8609	.765510	2.2743	1.44644
242794 240000	939.52	.841530	2.8229	.755320	2.5363	1.61310
249001 246063	922.8	.82654	2.737	.73245	3.251	2.0677
	-		7		. 0	+4
250000 247039	922.8	.82654	2.737 <sup>-7</sup>	73245	3.4230	2.1772+4
253033 250000	922.8	.82654	2.737	73245	4.004	2.5464
260000 256799	922.8	.82654	2.737	73245	5.737	3.6485
263282 260000	922.8	.82654	2.737	-73245	6.795	4.3218
270000 266549	922.8	.82654	2.737	.73245	9.609	6.1111
273541 270000	922.8	.82654	2.737	.75245	1.153	7.3349
280000 276291	922.8	.82654	2.737	73245	1.609	1.0231+5
285810 280000	922.8	.82654	2.73?	·732!15	1.957	1.2449
290000 286023	922.8	.82654	2.737	73245	2.692	1.7120
294089 290000	922.8	.82654	2.737	.73245	3.322	2.1128
299516 295276	922.8	.82654	2.737	.73245	4.391	2.7929
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APPENDIX A

COMPARISON OF PROMINENT AERONAUTICAL STANDARD ATMOSPHERES

#### Constants Employed

Prop- erties	Dimensions	Toussaint 1919 Prance	Gregg 1922 U.S.	ICAN 1924 Internat.	Diehl 1925 U.S.	Warfield 1947 U.S.	ICAO 1952 U.S. and Internat.	Minzner 1956 U.S.	
Po	13k			760	760		760	760	
Po	mp	1013.3		1013.2	1015.25	1013.25	1013.250	1013.250	
Po	kg m <sup>-3</sup>	1.225 1.225		1.2256	1.2255	1.2255	1.2250	1.2250	
PM <sub>O</sub>	*K kg	352.8	352.8	352.969	352 <b>.</b> 945		353.000	353.000	
R	joules K kg	2.8720		2.8705	2.87084	2.87084		2.8704	
R*	loules K kg		·				,	8.31439	
70	*c				15.	15.	15.	15.	
T <sub>1</sub>	*x					273.	273.16	273.16	
М <sub>о</sub>						28.966	28.966	28.966	
(c <sub>a</sub> ) <sub>o</sub>	R Sec				-	340.22	340.43	340.292	
r						1.4	1.401119	1.4	
8	•ĸ					120	120	110.4	
β	$\frac{kg}{\sec = (.K)^{1/2}}$					1.488,82 × 10 <sup>-5</sup>	1.495,26 × 10 <sup>-5</sup>	1.458 × 10 <sup>-5</sup>	
r	я					6,367,623		6,356,766	
	Sea Level Atmospheric Composition, Major Constituents by Per Cent								
и <sub>2</sub> 0				0.0	0.0		0.0	- 0.0	
N <sub>2</sub>			ĺ						
0 <sub>2</sub>			i	78.03			78.09	78.09	
				20.99			20.95	20.95	
A				.94			-93	-93	
co <sup>2</sup>				.94			.05	.03	

#### APPENDIX A CONTINUED

## COMPARISON OF PROMINENT AERCHAUTICAL STANDARD ATMOSPHERES

#### Temperature-Altitude Profiles

Toussaint 1919 1922 1924 1925 1947 U.S. end 1956 U.S.  Altitude km or km' t a t a t a t a t a t a t a t a t a t	r
km or km' t a t a t a t a t a t a t a	1
0 15 T 15 T 15 T 15 T 15 T 15 T	<sup>L</sup> M
	T
	1 1
4 -11 6 -11 -6.5 -6.5 -6.5 -6.5	6.5
6 -24 -24 -24 -24 -24	
8 -37 -37 -37 -37	
10 -50 -50 -50 -50 -50	
10.769,23	1
11   -57 +   -56.5 +   -55 0     -56.5 -   -56.5	Ť
20   -57 \(\frac{1}{2}\)     -56.5 \(\frac{1}{2}\)   -55.5 \(\frac{1}{2}\)   -56.5 \(\frac{1}{2}\)   -56.5	0
1 25 1 1 1 1 1 1 1 1	†
32 -55 + -35.5	+3
	<u> </u>
50   +77 +	O I
55 +9.5	-3.9
60 +77 +	-3.9
$ \begin{array}{c ccccccccccccccccccccccccccccccccccc$	十一
78	0 ·
85	
$+3\frac{24}{57}$ $-76.3$	†
120 +102 -	+3.5
49.7	+10
539.7	+
200	+5.8 1
500   2424.7	<u> </u>

Footnotes:  $t_{M}$  is in °C  $t_{M} = \frac{t}{M} \cdot M_{O}$ t is in °C  $c_{M} = \frac{t}{M} \cdot M_{O}$   $c_{M} = \frac{t}{M} \cdot M_{O}$ Below 90 km²,  $t_{M} = t$ 

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#### APPENDIX B

#### Constants

Defined Independent Physical Constants Adopted as Being Exact

mks absolute units	cgs units
$g_0 = 9.806,65 \text{ m sec}^{-1}$	980.665 cm sec <sup>-1</sup>
$M_o = 28.966$ (dimensionless)	28.966 (dimensionless)
N = $6.023,80 \times 10^{26}$ (dimensionless) (for a kg-mol)	6.023,80 x°10 <sup>23</sup> (dimensionless) (for a gm-mol)
P <sub>o</sub> = 1.013,250 x 10 <sup>5</sup> nt m <sup>-2</sup> or 0.76 m of mercury	1.013,250 x 10 <sup>6</sup> dynes cm <sup>-2</sup> or 76.0 cm of mercury
$R* = 8.314,39 \times 10^3 \text{ joules (°K)}^{-1} \text{ kg}^{-1}$	8.314,39 x 10 <sup>7</sup> ergs (°K) <sup>-1</sup> gm <sup>-1</sup>
r = 6.356,766 x 10 <sup>6</sup> m	6.356,766 x 10 <sup>8</sup> cm
s = 110.4°K	110.4°K
T <sub>1</sub> = 273.16°K	273.16°K
t <sub>o</sub> = 15°C	15°C
$\beta = 1.458 \times 10^{-6} \text{ kg sec}^{-1} \text{ m}^{-1} (\text{°K})^{-\frac{1}{2}}$	1.458 x $10^{-5}$ gm sec <sup>-1</sup> cm <sup>-1</sup> (°K) <sup>-<math>\frac{1}{2}</math></sup> [or poise (°K) <sup>-1</sup> ]
$\gamma = 1.4$ (dimensionless)	1.4 (dimensionless)
$6 = 3.65 \times 10^{-10} \text{ m}$	3.65 x 10 <sup>-8</sup> cm

#### English Units

 $t_i = 32^{\circ}F$ 

#### Numerical Constants (not exact)

Log<sub>10</sub> e = .434,294,481,9  $\pi$  = 3.141,592,654  $\sqrt{2}$  = 1.414,213,562

#### APPENDIX C

#### Conversions

Defined and Derived Conversion Factors for Transformation of Units and Scales

# Metric to English Conversions and Vice Versa

a. Defined relations

(<u>(</u>)

b. Derived relations

c. Conversion factors

$$1 = 0.304,8 \text{ m ft}^{-1}$$

$$1 = 1,852 \text{ m (i n mi)}^{-1}$$

$$1 = .453,592,3 \text{ kg lb}^{-1}$$

$$1 = \frac{1,852}{.304,8} \text{ ft (i n mi)}^{-1}$$

# 2. Geometric Altitude to Geopotential Altitude

a. Defined relations

1 standard geopotential meter = 9.806,65 joules kg<sup>-1</sup> (exact);  
geopotential altitude, 
$$H(m') = \frac{1}{G} \int g dZ$$
  
where  $G = \frac{9.806,65 \text{ joules kg}^{-1}}{1 \text{ m'}}$  (exact)  
= 9.806,65 m<sup>2</sup> sec<sup>-2</sup> m'<sup>-1</sup>

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- b Derived relations
  - 1 standard geopotential foot = 0 304,8 standard geopotential meter
  - (exact)

    1 standard geopotential meter = 3 280,839,895,013 standard geopotential feet
- c Conversion factors

$$1 = 0.304.8 \, \text{m}' \, \text{ft}'^{-1}$$

- 3 Temperature Unit and Scale Conversions
  - a Defined relations

$$t(^{\circ}C) = T(^{\circ}K) - T_{1}(^{\circ}K)$$
  
where  $T_{1}(^{\circ}K) = 273.16^{\circ}K$   
 $T(^{\circ}R) = 1 \ 8 \ T(^{\circ}K)$   
 $t(^{\circ}F) - t_{1}(^{\circ}F) = T(^{\circ}R) - T_{1}(^{\circ}R)$   
where  $t_{1}(^{\circ}F) = 32(^{\circ}F)$ 

b Derived relations

c Conversion factors

$$1 = 1.8^{\circ} R (^{\circ} K)^{-1}$$

- 4. Absolute Systems to Absolute Force. Gravitational Systems
  - a Defined relations
    - 1 kilogram (force), kgf =  $9.806,65 \text{ m sec}^{-2} \times 1 \text{ kilogram (mass), kg.}$
    - 1 round (force)  $1bf = \frac{9.805,65}{3048}$  ft  $sec^{-2} \times 1$  pound (mass), 1b

- b. Derived relations
  - $1 \text{ kgf sec}^2 \text{ m}^{-1} = 9.806,65 \times 1 \text{ kg}$
  - 1 slug = 1 lbf  $\sec^2 \text{ ft}^{-1} = \frac{9.806,65}{.504,8} \times 1 \text{ lb}$
  - 1 lbf = .453,592,3 kgf
- c. Conversion factors
  - $1 = 9.806,65 \text{ m sec}^{-2} \text{ kg kgf}^{-1}$
  - $1 = \frac{9.806,65}{.304,8}$  ft  $sec^{-2}$  lb lbf<sup>-1</sup>
  - 1 = .453,592,3 kgf lbf<sup>-1</sup>

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#### APPENDIX D

#### Assumptions

$$g = g_0 \left(\frac{r}{r+2}\right)^2$$

$$dP = -g \rho dZ$$

$$\rho = \frac{PM}{R^*T}$$

$$c^{M}\left(\frac{T}{M}\right) = M^{T}$$

$$T_{M} = (T_{M})_{b} + L_{M} (H - H_{b})$$

# where $\mathbf{L}_{\mathbf{M}}$ is given by the following table

L <sub>M</sub> in *K m'-1	Altitude	La	yer in m'
-0.0065 exact	-5,000.	to	0.
-0.0065 exact	0.	to	11,000.
0.0 exact	11,000.	to	25,000.
+0.0030 exact	25,000.	to	47,000.
0.0 exact	47,000.	to	53,000.
-0.0039 exact	53,000.	to	75,000.
0.0 exact	75,000.	to	90,000.
+0.0035 exact	90,000.	to	126,000.
+0.0100 exact	126,000.		
+0.0058 exact	175,000.		

$$H_s = \frac{R*T}{gM}$$

$$C_{\rm g} = \left(\frac{\gamma p}{\rho}\right)^{1/2}$$

$$\overline{v} = \left(\frac{8R*T}{\pi M}\right)^{1/2}$$

$$\omega = \rho \text{ g, (not } \rho \text{ g_o})$$
For -5,000.  $m' \leq H \leq +90,000. m'$ 

$$M = 28.966 \text{ (exact)}$$
For 90,000.  $m' \leq H \leq 175,000. m'$ 

$$M = \frac{23.160,126,7 \text{ H} - 1,757,856.047}{H - 78,726.253}$$
For 175,000.  $m' \leq H \leq 500,000. m'$ 

$$M = \frac{13.139,119,0 \text{ H} + 514,492.021}{H - 56,969.889}$$

$$v = \frac{M'}{\rho}$$

$$n = \frac{N}{v}$$

$$L = \frac{1}{\sqrt{2\pi6^2n}}$$

 $n = \frac{N}{v}$   $L = \frac{1}{\sqrt{2\pi\sigma^2 n}}$   $v = \frac{\overline{V}}{L}$   $T = T_M \left(\frac{M}{M_o}\right)$   $\mu = \frac{\beta}{T} \frac{T^{3/2}}{T + S}$ 

O APPENDIX E

Sea-Level Values of the Atmospheric Properties in Metric Units

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	mks units	cgs units
$(c_s)_o$	340.292,046 m sec <sup>-1</sup>	34,029.204,5 cm sec <sup>-1</sup>
RO	9.806,65 m sec <sup>-2</sup>	980.665 cm sec <sup>-2</sup>
$(H_s)_o$	8.434,413,43 x 10 <sup>3</sup> m	8.434,413,43 x 10 <sup>5</sup> cm
Lo	6.631,722,29 x 10 <sup>-8</sup> m	6.631,722,29 x 10 <sup>-6</sup> cm
$M_{O}$	28.966 (dimensionless, exact)	28.966 (dimensionless, exact)
M'o	28.966 kg (exact)	28.966 gm (exact)
n <sub>o</sub>	2.547,552,07 x 10 <sup>25</sup> m <sup>-3</sup>	2.547,552,07 x 10 <sup>19</sup> cm <sup>-3</sup>
Po	101,325 nt m <sup>-2</sup>	1,013,250. dynes cm <sup>-2</sup>
Po	.76 m Hg	.76 cm Hg
Po	10,532.274,5 kgf m <sup>-2</sup>	
To	288.16°K	288.16°K
$(T_M)_o$	288.16°K (exact)	288.16°K (exact)
$\overline{v}_{o}$	458.942,035 m sec-1	45,894.203,5 cm sec <sup>-1</sup>
$v_{c}$	23.645,444,1 m <sup>3</sup> for a kg-mol	23,645.444,1 cm <sup>3</sup> for a gm-mol
$\eta_{o}$	$1.460,741,29 \times 10^{-5} \text{ m}^2 \text{ sec}^{-1}$	$1.460,741,29 \times 10^{-1} \text{ cm}^2 \text{ sec}^{-1}$
$\mu_{o}$	1.789,428,53 x 10 <sup>-5</sup> kg m <sup>-1</sup> sec <sup>-1</sup>	$1.789,428,53 \times 10^{-4} \text{ gm cm}^{-1} \text{ sec}^{-1}$
$\mu_{0}$	$1.824,709,28 \times 10^{-6} \text{ kgf sec m}^{-2}$	

#### Sea-Level Values of the Atmospheric Properties in Metric Units

	mks units	cgs units
۲o	6.920,1004,89 x 10 <sup>9</sup> sec <sup>-1</sup>	6.920,404,89 x 10 <sup>9</sup> sec <sup>-1</sup>
ρο	1.225,013,99 kg m <sup>-3</sup>	1.225,013,99 x 10 <sup>-3</sup> gm cm <sup>-3</sup>
ρο	.124,916,663 kgf sec <sup>2</sup> m <sup>-4</sup>	
ωο	12.013,283.5 kg m <sup>-2</sup> sec <sup>-2</sup>	1.201,328,35 gm cm <sup>-2</sup> sec <sup>-2</sup>
ω <sub>o</sub>	1.225,014,00 kgf m <sup>-3</sup>	

### Ice-Point Values of Same Atmospheric Properties

	mks units	cgs units
ni	$2.687,445,47 \times 10^{25} \text{ m}^{-3}$	2.687,445,47 x 10 <sup>19</sup> cm <sup>-3</sup>
v <sub>i</sub>	22.414,594,4 m <sup>3</sup> for a kg-mol	22,414.596,4 cm <sup>3</sup> for a gm-mol
ρ,	1.292,283,037 kg m <sup>-3</sup>	1.292.283.037 x 10 <sup>-3</sup> cm cm <sup>-3</sup>

# APPENDIX F Sea-Level Values of the Atmospheric Properties in English Units

(
$$C_8$$
)<sub>0</sub> 1.116,443,72 x 10<sup>3</sup> ft sec<sup>-1</sup>
 $E_0$  32.174,048,5 ft sec<sup>-2</sup>
( $H_8$ )<sub>0</sub> 2.767,196,00 x 10<sup>4</sup> ft

 $L_0$  2.175,761,91 x 10<sup>-7</sup> ft

 $M_0$  28.966
 $M'_0$  28.966 lbs

 $n_0$  7.213,864,1 x 10<sup>23</sup> ft<sup>-3</sup>
 $P_0$  68,087.267 lb ft<sup>-1</sup> sec<sup>-2</sup>
 $P_0$  29.921,259,8 in Hg

 $P_0$  2,116.216,95 lbf ft<sup>-2</sup>
 $T_0$  518.688°R
( $T_M$ )<sub>0</sub> 518.688°R

 $\overline{V}_0$  1.505,715,34 x 10<sup>3</sup> ft sec<sup>-1</sup>
 $V_0$  83.503,098 ft<sup>3</sup>
 $\eta_0$  1.572,328,83 x 10<sup>-1</sup> ft<sup>2</sup> sec<sup>-1</sup>
 $\mu_0$  3.737,299,76 x 10<sup>-7</sup> lbf sec ft<sup>-2</sup>
 $\rho_0$  6.920,404,9 x 10<sup>9</sup> sec<sup>-1</sup>
 $\rho_0$  2.376,919,99 x 10<sup>-3</sup> lbf sec<sup>2</sup> ft<sup>-4</sup>
 $\rho_0$  2.460,514,77 lb ft<sup>-2</sup> sec<sup>-2</sup>
 $\rho_0$  7.647,513,7 x 10<sup>-2</sup> lbf ft<sup>-3</sup>

•	ADDIEVIEUCE MELLIC TRDIES OF the ARDC Model Atmosphere (1956) to 542,686 m	eargar, o	of the AR	DC Model At	cmosphere	1956 +0 540	באל ד
						111/20	3 2
ĦI	118	됩	ΕΉ	EH I	ΣI	А	و ا
_ E	щ	•K m'-1	'n	<b>*</b>	l	। ନ୍ଥ	ka = -3
-5,000	-4,996.070,27	0	320.66	320.66	28.966	1.777,6 × 10 <sup>3</sup>	1.931.2
0	0	6000.0-	288.16	288.16	28.966		1.225.0
000,11	11,019.067,83	-0.0065	216.66	216,66	28.966	2.263,2 x 10 <sup>2</sup>	3.679.1 × 10 <sup>-1</sup>
*20,000	20,063.123,68	0000	216.66	216.66	28.966	5.474,8 x 10 <sup>1</sup>	8.803.4 x 10"2
25,000	25,098.708,63	00000	216.66	216.66	28.966		4.001,6 × 10 <sup>-2</sup>
32,000	32,161.903,22		237.66	237.66	28.966		1.272,1 × 10 <sup>-2</sup>
000 <b>'</b> 2†	47,350.092,22	200	282.66	282.66	28.966		1.484.5 × 10 <sup>-3</sup>
53,000	53,445.606,64	00000	282,66	282.66	28.966	5.832,0 x 10 <sup>-1</sup>	79188,1 × 10-4
75,000	75,895.448,82		196.86	196.86	196.86 28.966		4.339 x 10 <sup>-5</sup>
90,000		5500.04	196.86		28.966	196.86 28.966 1.815,4 x 10 <sup>-3</sup>	3.213 × 10-6
126,000		+0.0100	322.86	273.6	. 45.42	24.54 1.451,0 x 10"5	1.566 x 10 <sup>-8</sup>
175,000		+0.0058	812.86	0.699	25.84	6.189,5 x 10 <sup>-7</sup>	2.655 x 10°10
200,000			1,537.86	973.5	18.34	1.447,3 x 10-8	3.279 x 10 <sup>-12</sup>
500,000	542,685.673,2		2,697.86 1,489.	1,489.	15.99	7.698 × 10 <sup>-9</sup>	6.819 x 10-14

\* Top of 1952 United States (ICAO) Standard Atmosphere \*\* Top of 1956 United States Standard Atmosphere

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	œ. į	slugs ft-3	3.745,7 × 10"3	2.376,9 x 10"3	7.061,1 x 10 <sup>-14</sup>	1.725,1 × 10 <sup>-4</sup>	7.764,4 × 10 <sup>-5</sup>	2.468,2 x 10 <sup>-5</sup>	2.880,3 × 10 <sup>-6</sup>	1.394,7 × 10-6	8.419,8 x 10 <sup>-8</sup>	6.233,5 × 10"9	3.038,0 × 10 <sup>-11</sup>	5.147,1 × 10 <sup>-13</sup>	6.361,8 x 10 <sup>-15</sup>	1.323,1 × 10 <sup>-16</sup>
Abbreviated English Tables of the ARDC Model Atmosphere to 1,780,465 Ft.	<u>п</u> 1	lbf ft-2	3.711,00 x 10 <sup>3</sup>	2.116,22 x 10 <sup>3</sup>	28.966 4.726,8 x 10 <sup>2</sup>	1.154,8 × 10 <sup>2</sup>	28.966 3.197,5 × 10 <sup>1</sup>	427.788 28.966 1.812,4 x 10 <sup>1</sup>	508.788 28.966 2.515,5 x 10 <sup>0</sup>	1.218,0 × 10 <sup>0</sup>		3.791,5 × 10 <sup>-3</sup>	3.030,5 x 10 <sup>-5</sup>	1.292,7 x 10 <sup>-6</sup>	3.022,8 × 10 <sup>-8</sup>	4,856.148 2,681,000 15.990 1.102,9 x 10 <sup>-9</sup> 1.323,1 x 10 <sup>-16</sup>
sphere	ΣI		28.966	28.966	28.966	28.966		28.966	28.966	28.966	28.966	28.966	24.54	23.84	18.3	15.990
Model Atmo	T.	a.	577.188	51.8.688	389.988	389.988	389.988	427.788	508.788	508.788 28.966	354.348 28.966	354.348	4.564	1,204.000	1,752.000	2,681,000
of the ARDC	F X	ਕਿ	577.188	518.688	539.988	389.988	389.988	427.788	508.788	508.788	354.348	354.348	581.148	1,463.148	2,768.148	4,856.148
English Tables	뭐	R ft-1	091,995,500.=	091 993 200 -		0 0 0	4.001 6lts 000	1,001,047,920 t	T. W.L., U47, Y.C.	909 021 600 -	7000	045 045 CC +	00'( 98'( 500 ±	911,981,500.+	פור פאר הסטי+	
Abbrevfated	118	f.	-16,391.301	0	36,151.797	65,823.897	82, 344.844	105,518.055	155,348.071	175,346.478	249,000,816	299,516.183	421,745.411	590,400.544	1,033,003.330	1,780,464.807
-	ĦI	ft.	-16,4c4.199	0	56,089.239	#65,616.798	82,020.997	104,986,877	154,159.475	173,884.514	246,062,992	295,275,590	, 413,386.826	574,146.981	**984,251,968	7,6,614,049,1

\* Top of 1952 United States (ICAO) Standard Atmosphere \*\* Top of 1956 United States Standard Atmosphere

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		6	SYSTEMS OF	HECHANICAL	CAL UKITS			
			METRIC	c	•		HHOLISH	
	<del></del>	Absolute cos	Absolute mks	Oravitational mks	al riks	theolute the	Gravitational fps	al fys
			7	^		8	•	
		•		Type I	11 *47.		Type I	n sect
			4 8 b	4 6 . 2.	£7 - **			48 - 14
Property D	Dimenstons							
Jength (altitude) (ccale height)		centimeter (fm)	meter (a)	meter (n)	E tet	foot (ft)	(t)	loot (12)
		gress (ges)	kilogram (kg)	"kgf sec2 m"1	Milogram (kg)	pound (1b)	slug	pound (%)
N S		second (sec)	second (sec)	second (sec)	second (sec)	(see) process	second (sec)	second (see)
	5-7 T *	dyne or an ca sec-2	~	*kilogram force *kilogram force (kgf)	*kilogram force (kgf)	pyundal (pdl)	**pound force (1bf)	*pount force (1bt)
	75	, ~ <sub>e</sub>		2 18 2	~4	r <sub>2</sub>	rt <sup>2</sup>	<sub>م</sub> ح
	÷	<b>^</b> E	~ <sub>*</sub>	~a	, a		દુષ	£21
(pomor)	f t-1	7,	" sec-1	m sec 1	M.000-1	ft 100-1	ft sec-1	1-98 Y
	7 1.5				2- 29 E	11 sec -2	ft me-2	ft sec <sup>-2</sup>
	" T51-1	8	, m		a zgy.	# 15d	as Jel ce	•• lbr :r
entiel	F +-5	ergs gr. 1 or cm' ser: 2	Joules kg-1	n² sec-2	*kgf p kg"l	pel ft lb"l or ft2 sec-2	1bf ft slug-1 or ft sec-2	**1bf ff 1b ****
DITERNITY	L-14-2	dyne cm-2 . 10-1 mb	nt m-2 - 10-2 mb "kgf m-2	*,kgf #*2	*kgf =-2		1bf ft-2	2-19 191.e.
	L.	6 cm -3	7g m²-3	*kgf sec <sup>2</sup> m <sup>-b</sup>	λε π <sup>-5</sup>	39 tc-2	** or lbf sec ft-	10 ft-3
specific weight	a 1 -2°-2	G C - 2 MC - 2	1g m <sup>-2</sup> sec <sup>-2</sup>	*kgf m*3	Ng m 2 sec -2	16 ft <sup>-2</sup> 1906 <sup>-2</sup>	*** 1 tr 2 *** 1 tr 2 **** *****************************	1b rt-2 sec-2
y demail	ũ	÷.	۳,	ŗ,a	٠,	r.3.	n-5	n-5
collision frequency t-1	7,	1-11	**************************************	**************************************			<b>1-1008</b>	**c-1
viscosity	1-71-F	polse -1 mc-1	1, n-1 sec-1	g_ w sec ya,	kg m 1 sec-1	15 ft <sup>-1</sup> eec <sup>-1</sup>	**slugs ft"1 sec-1	13 ft <sup>-1</sup> sec <sup>-1</sup>
kinesatic viscosity £20°1	1-127	Ca <sup>2</sup> sec <sup>-1</sup>	n <sup>2</sup> sec <sup>-1</sup>	12 sec-1	n <sup>2</sup> sec <sup>-1</sup>	ft <sup>2</sup> sec <sup>-1</sup>	rt² 44c-1	n² we¹l
		used by physicists	used by electri- cal engineers and physicists	used by Euro- pean maro- dynamicists	·		ughl by American serodynamiciats	used by some mechanical enginours
					200 27 1121	10. Fories (exect)	the following the state of the following the large as muchan associated with	associated with

APPENDIX

\*At sea level and at a latitude of 45° 52° 40° the numbers associated with these units will be only 1/9.6065 (exact) as large as numbers associated will corresponding units of system 2.4° 52° 40° the numbers associated with these units will be only 1/2.174,040,55 as large as numbers associated with corresponding units of system 5.4° 50° 40° the numbers associated with the same ratio applies at all altitudes.

\*\*For the absolute-force versions of gravitational units as used in this MCCE, the same ratio applies at all altitudes.

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APPENDIX K

## Comparison of the Magnitudes of Comparable Units in the Metric Absolute cgs and mks Systems of Mechanical Measure

 $1 m = 10^2 cm$ length  $1 \text{ kg} = 10^3 \text{ gm}$ mass

1 sec = 1 sec time

 $1 \text{ nt} = 10^5 \text{ dynes}$ force  $1 m^2 = 10^4 cm^2$ 

area  $1 \text{ m}^3 = 10^6 \text{ cm}^3$ volume

 $1 \text{ m sec}^{-1} = 10^2 \text{ cm sec}^{-1}$ speed (sound)

 $1 \text{ m sec}^{-2} = 10^2 \text{ cm sec}^{-2}$ acceleration

1 nt m = 10<sup>7</sup> dynes cm 1 joule = 10<sup>7</sup> ergs energy (work)

1 joule kg<sup>-1</sup> =  $10^{4}$  ergs gm<sup>-1</sup> 1 m<sup>2</sup> sec<sup>-2</sup> =  $10^{4}$  cm<sup>2</sup> sec<sup>-2</sup> geopotential

1 nt  $m^{-2} = 10^1$  dynes cm<sup>-2</sup> pressure  $1 \text{ kg m}^{-3} = 10^{-3} \text{ gm cm}^{-3}$ 

density

 $1 \text{ kg m}^{-2} \text{ sec}^{-2} = 10^{-1} \text{ gm cm}^{-2} \text{ sec}^{-2}$ specific weight

 $1 m^{-3} = 10^{-6} cm^{-3}$ number density  $1 \sec^{-1} = 1 \sec^{-1}$ 

collision frequency

l newton sec  $m^{-2} = 10^{1}$  dynes sec cm<sup>-2</sup> l kg m<sup>-1</sup> sec<sup>-1</sup> = 10 gm cm<sup>-1</sup> sec<sup>-1</sup> = 10 poise coefficient of viscosity

 $1 m^2 sec^{-1} = 10^4 cm^2 sec^{-1}$ kinematic viscosity

# Pressure in Terms of the Bar or Millibar (mb)

1 bar =  $10^3$  millibars (mb) =  $10^5$  nt m<sup>-2</sup> =  $10^6$  dynes cm<sup>-2</sup>

#### APPENDIX L

## Atmospheric Density Expressed as a Single Function or Altitude

At a recent Ad Hoc Conference on Units and Constants for Satellite Orbit Computations, this MODEL was adopted as a basis for initial calculations of IGY satellite orbits. Dr. Jacchia 3 who had received a prepublication copy of the MODEL, prepared and presented the following equations as closely representing the atmospheric density of this MODEL above 100 km altitude.

$$\log_{10}\rho = -10.919 - 0.004483Z + 7.321e^{-0.00685Z} + 3.400e^{-0.8} \left[\frac{Z}{100}\right]^{3}$$
 (1)

$$\log_{10}\rho = -11.019 - 0.00181H + 7.300e^{-0.0067H} + 3.700e^{-0.87 \left[\frac{H}{100}\right]^3}$$
 (2)

where  $\rho$  is the atmospheric density in kg/m³, Z is the geometric height above sea level in km, and H is the geopotential height in geopotential km. A comparison of densities computed from these equations with densities from the ARDC Model are tabulated on the next page.

Residuals  $\Delta \log_{10} \rho$  (ARDC Model Atmosphere densities minus interpolating formula) are given in the following table:

Z (km)	log <sub>10</sub> <b>/</b> (kg/m <sup>3</sup> )	Δ log <sub>]O</sub> <b>ρ</b> [(from (1)]	/ H /(km)	log <sub>10</sub> <b>/</b> (kg/m <sup>3</sup> )	Δ log <sub>10</sub> <b>ρ</b> [(from (2)]
0 25 50 75 100 125 150 175 200	+0.088 -1.378 -2.965 -4.304 -6.147 -7.629 -8.750 -9.462 -9.955	+.286 +.126 096 +.145 +.002 +.027 006 012	0 25 50 75 100 125 150 175 200	+0.088 -1.398 -2.986 -4.363 -6.258 -7.754 -8.871 -9.576 -10.068	+.107 184 267 +.037 043 +.051 +.001 010
250 300 350 400 450 500	-10.725 -11.328 -11.822 -12.241 -12.604 -12.922	006 002 .000 +.001 +.003	250 300 350 400 450 500	-10.859 -11.484 -12.001 -12.442 -12.826 -13.166	005 .000 +.002 +.001 001

Little effort was made to secure a good fit for heights smaller than  $100\ \mathrm{km}$ .

#### APPENDIX M

# Effective Radius of the Earth

The limitations of the inverse square law for determining the acceleration of gravity were discussed in Section 2.1 of this paper. A value of effective earth's radius was introduced as a means of offsetting some of these limitations.

The inverse square law for expressing the acceleration of gravity was

$$g = g_{\phi} \left( \frac{r_{\phi}}{r_{\phi} + Z} \right)^{2} . \tag{M-1}$$

The partial derivative of g with respect to Z is

$$\frac{\partial g}{\partial z} = 2g_{\phi} \left( \frac{r_{\phi}}{r_{\phi} + z} \right) \frac{(-r_{\phi})}{(r_{\phi} + z)^{2}}.$$
 (M-2)

This partial derivative evaluated at Z = 0 becomes

$$\left(\frac{\partial g}{\partial Z}\right)_{Z=0} = \frac{-2g\phi}{r\phi} . \tag{M-3}$$

Thus, if the actual sea-level value of  $g_{0}$  and the actual sea-level value of  $(\partial g/\partial Z)$  for the particular latitude are introduced into Eq. (M-3), the value of  $r_{0}$  consistent with these realistic quantities is the effective earth's radius at that latitude.

Harrison presented the following expression for  $(\partial g/\partial Z)_{Z=0}$  as a function of latitude  $\emptyset$ , without indicating its derivation.\*

$$-\frac{\partial g}{\partial Z} = 3.085,462 \times 10^{-6} + 2.27 \times 10^{-9} \cos 2 \phi$$
$$-2 \times 10^{-12} \cos 4 \phi.$$

<sup>\*</sup> This equation appears to be related to Lambert's alternating power series expression for g in terms of Ø and Z, which equation is discussed in Appendix O.

Using this expression, the effective earth's radius  $\overline{\mathbf{r}}_{\phi}$  at latitude  $\phi$  is

$$\vec{r}_{\phi} = \frac{2g\phi}{3.085,462 \times 10^{-6} + 2.27 \times 10^{-9} \cos 2 \phi - 2 \times 10^{-12} \cos 4 \phi}$$

For  $\phi = 45^{\circ}$  32' 40",

$$g_{\phi} = g_{o} = 9.806,65 \text{ m sec}^{-1}$$

and

$$\overline{r}_{0} = r = 6,356,766 m$$
.

#### APPENDIX N

# Acceleration of Gravity

#### 1. Background

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The inverse square law employed in this MODEL for the computation of the acceleration of gravity has been adjusted at sea level to account for the effective sea level value and the vertical gradient of g at that point, by means of an effective earth's radius (see Appendix M). This correction accounts for the centrifugal acceleration which a body experiences at sea level, by virtue of the earth's rotation, but it does not account for the fact that this centrifugal acceleration increases rather than decreases with altitude. Since the centrifugal acceleration is opposite in direction to the gravitational acceleration, the net or effective value of g falls off more repidly with altitude than even the adjusted inverse square law predicts. Because the actual earth's radius and the centrifugal acceleration both depend upon latitude, any general expression for a resultant or effective acceleration must be a function of both altitude Z and latitude  $\phi$ .

Lambert 38 developed such a general expression\* for g in the form of

$$g = c_1 - (a_2 + b_2 \cos 2\phi)Z + (a_3 + b_3 \cos 2\phi)Z^2$$

$$- (a_4 + b_4 \cos 2\phi)Z^3 + (a_5 + b_5 \cos 2\phi)Z^4$$

$$+ \dots \qquad (N-1)$$

where

g = the acceleration of gravity in m sec<sup>-2</sup>, Z = geometric altitude in m,  $\phi$  = latitude in degrees  $c_1$  =  $g\phi$ , see level value of g at latitude  $\phi$ ,  $a_2$  = 3.085,462 x 10<sup>-6</sup>  $b_2$  = 2.27 x 10<sup>-9</sup> •  $a_3$  = 7.254 x 10<sup>-13</sup>  $b_3$  = 1.0 x 10<sup>-15</sup> •  $a_4$  = 1.517 x 10<sup>-19</sup>  $b_4$  = 6 x 10<sup>-22</sup> •  $a_5$  = 2.97 x 10<sup>-26</sup>  $b_5$  = 2 x 10<sup>-28</sup>

<sup>\*</sup> The fifth term (in Z<sup>h</sup>) has not been published, but was provided by Col. C. Spohn, of Air Weather Service USAF, who probably obtained it from Lembert or Harrison.

For the case when  $\beta = 45^{\circ}$  32' 40", as in this MODEL, chosen to agree with  $g_0 = 9.806,65 \text{ m scc}^{-2}$ ,

$$\cos 20 = \cos 91^{\circ} 5' 20'' = -\sin 1^{\circ} 5' 20'' = -.019,003,7.$$
 (N-2)

For this value of  $\emptyset$ , Eq. (N-1) becomes

$$g = c_1 - c_2 Z + c_3 Z^2 - c_4 Z^3 + c_5 Z^4 - \dots$$
 (N-3)

where

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$$c_1 = 9.806,65 \text{ (exact)}$$
 m sec<sup>-2</sup>  
 $c_2 = .306,541,8_8 \times 10^{-5}$  m<sup>0</sup> sec<sup>-2</sup>  
 $c_3 = .007,253,8_1 \times 10^{-10}$  m<sup>-1</sup> sec<sup>-2</sup>  
 $c_4 = .000,151,6_{89} \times 10^{-15}$  m<sup>-2</sup> sec<sup>-2</sup>  
 $c_5 = .000,002,96_{96} \times 10^{-20}$  m<sup>-3</sup> sec<sup>-2</sup>

The reliability of the limit of this series in expressing the true value of g at any altitude is unknown to the authors of this report. It is assumed that this function represents the best available analytical expression for g in terms of Z and  $\emptyset$ . The small number of available terms and significant figures, however, places limitations on the evaluation of the series at high altitudes.

#### 2. Problem

It is necessary to determine the limitations which the small number of terms and the small number of significant figures place upon the evaluation of the function at various altitudes. It is further necessary to compare the results of the adjusted, inverse-square-law function for g with the values obtained from the infinite series function for g.

The extent to which the availability of only five terms limits the value of g at various altitudes has been studied for the case where  $\emptyset$  = 45° 32' 40" with the results indicated below. In the course of the analysis it was found that several additional terms were necessary to determine the value of g to the desired accuracies at altitudes above 150 km. The values of the additional terms were estimated by graphical extrapolation, and refined values of g were computed for various altitudes. These values of g were then compared with values from the inverse square law, using the effective earth's radius at 45° 32'  $\pm 0$ ° as determined in Appendix M.

## Results, Concerning Required Number of Terms in Equation (N-3) For Various Degrees of Accuracy

Equation (N-3), limited to four terms as published, provides accuracies

0

1

of one part in 9,800,000, or seven significant figures, for altitudes up to only about 60 km. The fifth term permits the equation to be used up to about 150 km with the same accuracy, provided that the coefficient of the third term has one additional significant figure. By means of extrapolation it was estimated that with five additional terms in Eq. (N-3), g could be determined to the stated accuracy for altitudes up to 1,140 km, provided a sufficient number of significant figures are added to all the terms beyond the first two. For other accuracies the maximum altitude to which g may be computed with a given number of terms in Eq. (N-3) is given in Table (N-I), neglecting significant figures in existing terms.

Number of		Number o	f Signifi	cant Figu	res Requi	red in g	
Terms Available	5_	3	4	5	6	7	· 8
2 3 4 5 6 7 8 9	260 700 1100	80 330 650 1000	25 150 370 640 950 1300	8 75 200 400 610 900 1100	60 110 250 420 610 860 1200	20 60 150 260 440 610 830 1140	35 100 180 320 480 620 800

Table N-I. Estimated maximum altitude in km for which a specified number of terms in Eq. (N-3) will yield accuracies of a specified number of significant figures in g, provided the various coefficients have a sufficient number of significant figures.

# 4. Results, Concerning Limitations Due to Available Significant Figures in Equations (N-1) and (N-5).

The number of significant figures in the coefficients of Eq. (N-3) stems directly from the number available in the coefficients of Eq. (N-1). An analysis of the limitations of these equations shows that for g accurate to four significant figures, these equations may be used up to 1,400 km.

For five-significant-figure accuracy in g, the accuracy of the coefficients limits the calculations to altitudes below 1,300 km; ror six-significant-figure accuracy in g, the calculations are restricted to altitudes below 500 km; while for seven-and eight-significant-figure accuracy in g, the maximum permissible altitudes are only 150 and 50 km, respectively. (see figure N-6)

Applying these restrictions to Table N-I, one obtains Table N-II.

Number of		Number of	Significant	Figures	Required	in g	
Terms Available	2	3	4	5	6	7	88
2 3 4 5 7 8 9	260 700 1100	80 330 650 1000	25 150 370 640 950 1300	8 75 200 400 610 900 1300	60 110 250 420 500 500 500	80 150 150 150 150 150 150	35 50 50 50 50 50 50 50 50 50 50 50 50 50

Table N-II. Estimated maximum altitude in km for which a specified number of terms of Eq. (N-2) will yield a specified number of significant figures' accuracy in the value of g, with the significant figures of existing coefficients limiting the results.

NOTE: Underlined figures are those limited by the number of significant figures in coefficients.

# 5. Results of Comparison of Values of g from Equation (N-3) with Inverse-Square-Law Values of g

The inverse-square-law values of g, for  $\emptyset=45^{\circ}$  32' 40", when the effective earth's radius is used, are in good agreement with the values of Eq. (N-3), with no differences occurring in the fifth significant figure below 100 km. Above this altitude the differences increase rather rapidly to a peak at 500 km, after which they fall off to zero somewhere between 700 and 800 km and increase negatively above that altitude. This large fall-off is due principally to the omission of term six which becomes extremely significant in the series at this altitude. Since this term is negative, its presence would reduce the value of Eq. (N-3) at these altitudes and tend to retain the increasing difference with the inverse-square-law value.

Values of g were recalculated from Eq. (N. 3) on the bases of four additional terms determined graphically, and these new values of g were then compared with the inverse-square-law values. In this latter comparison, the differences increased uniformly with altitude. Curves B and C of Fig. N-1 show the graphs of the two comparisons. Curve A in this figure shows the departure of the five-term-series value of g from the estimated nine-term-series value of g. Curves A and C are essentially the error curves of the five-term-series function and the inverse-square-law function,

respectively, assuming the nine-term-series value of g to be the most correct. At 150 km, the five-term-series function provides two more significant figures than the inverse square law. As altitude increases, however, the differential in accuracy drops proportionately to one significant figure at 330 km, and no difference at 750 km. A comparison of the maximum altitudes to which the five-term-sories function and the inverse-square-law function may each be used for various accuracies is given in Table N-III.

		Signi:	ficant F	igures	
	4	5	6	7	<u>8</u>
5 term series	640 500	1400 130	250 40	150 10	<u>50</u> 5

Table N-III. Comparison of maximum altitude to which each of two functions of g may be used for five different degrees of accuracy.

The numerical value of g by the several methods and the numerical differences between these values are given in Table N-VI.

#### 6. Method of Analysis

The analysis was performed by using twenty-one values of Z between 1 and 1,000 km, and independently evaluating each of the five terms of Eq. (N-3). The logarithms of the absolute values of each term were plotted as a function of the number of the term, and points corresponding to the same value of Z were connected to form the solid line portion of Fig. N-2. The lines were then extrapolated to regions corresponding to higher order terms. The values indicated for these terms by the extrapolations then served as estimated values for these terms.

The values of the several terms were then plotted as a function of altitude, as in Fig. N-3, with solid lines connecting the computed terms, and broken lines connecting the estimated terms. The analysis of the contribution of varying numbers of terms to the value of the total function was then made visually from this graph.

The significant figure analysis was performed on tabulated values of the several terms (Table N-IV and Table N-V) and the net results are plotted on Figs. N-4, N-5, and N-6.

Alt.		3rd Term	4th Term	5th Term
1	.003,085,418, <u>8</u>	.000,000,725,38	.000,000,000,151	.000,000,000,000,029
5	.015,427,094	.000,018,134,52	.000,000,018,965	.000,000,000,018,56
10	.030,854,18 <u>8</u>	.000,072,538,1	.000,000,151,698	.000,000,000,296, <u>96</u>
20	.061,708,37 <u>6</u>	.000,290,152,4	.000,001,213,51	.000,000,004,711
30	.092,562,56 <u>4</u>	.000,652,843	.000,004,095,60	.000,000,024,0 <u>54</u>
40	.123,416,75 <u>2</u>	.001,160,609	.000,009,708,0	.000,000,076,0 <u>22</u>
50	.154,270,9 <u>40</u>	.001,813,4 <u>52</u>	.000,018,961,1	.000,000,185,60
60	.185,125,1 <u>28</u>	.002,611, <u>372</u>	.000,032,764,8	.000,000,384, <u>86</u>
70	.215,979, <u>316</u>	.003,554, <u>367</u>	.000,052,029,3	.000,000,713, <u>00</u>
80	.246,8 <u>5</u> 3,5 <u>04</u>	.004,642, <u>44</u>	.000,077,665	.000,001,216, <u>35</u>
90	.277,687,6 <u>92</u>	.005,875, <u>59</u>	.000,110,581	.000,001,94 <u>8,3</u>
100	.308,541,880	.007,253,81	.000,151,6 <u>89</u>	.000,002,96 <u>9,6</u>
200	.617,083,7 <u>60</u>	.029,015, <u>24</u>	.001,213, <u>51</u>	.000,047,51 <u>3</u>
300	.925,625,6 <u>4</u>	.065,284, <u>3</u>	.004,095, <u>60</u>	.000,240, <u>54</u>
400	1.234,167,5 <u>2</u>	.116,060,9	.009,708, <u>1</u>	.000,760, <u>22</u>
500	1.542,709, <u>40</u>	.181,34 <u>5,2</u>	.018,9 <u>61</u> , <u>1</u>	.001,85 <u>6,0</u>
600	1.851,251,28	.261,13 <u>7,2</u>	.032,764, <u>8</u>	.003,84 <u>8,6</u>
700	2.159,793, <u>16</u>	.355,4 <u>36,7</u>	.052,02 <u>9,3</u>	.007,1 <u>30,0</u>
800	2.468,335, <u>04</u>	.464,2 <u>44</u>	.077,664	.012,16 <u>3,4</u>
900	2.776,876, <u>92</u>	.587,5 <u>59</u>	.110,5 <u>81</u>	.019,4 <u>83</u>
1000	3.085,418, <u>80</u>	.725,3 <u>81</u>	.151,6 <u>89</u>	.029,6 <u>96</u>

Table N-IV. Values of the first four variable terms of Eq. (N-3) for various altitudes from 1 km to 1,000 km.

NOTE: The underlined figures are beyond the limit of significance but are carried for smoothness.

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	6th Term	7th Term	8th Term	9th Term
100 200 300 400 500	.000,000,05 .000,001,8 .000,012 .000,055 .000,15	.000,000,001 .000,000,08 .000,0003,5 .000,0012	.000,000,000 .000,000,002 .000,000,04 .000,000,2 .000,001	.000,000,000 .000,000,000 .000,000,000 .000,000,
600 700 800 900 1000	.000,42 .000,9 .001,7 .002,6 .004,5	.000,045 .000,11 .000,24 .000,4 .000,7	.000,004 .000,014 .000,03 .000,05 .000,10	.000,000,3 .000,001,5 .000,003,5 .000,006,5 .000,013

Table N-V. Estimated values of terms 6 through 9 of Eq. (N-3) for altitudes between 100 and 1,000 km.

Alt.	$g = g_0 \left[ \frac{r}{r+2} \right]^2$	g* from 5 terms of Eq. (N-3)	g** from cstimated 9 terms of Eq. (N-3)	g - g*	g - g**	g* - g**
1 5 10 20 30 40	9.803,565,30 9.791,241,06 9.775,868,42 9.745,231,56 9.714,738,52 9.684,388,35	9.803,565,306 9.791,241,021 9.775,868,19 9.745,230,56 9.714,736,2 9.684,384,2	identical to adjacent column	.000,000,00 .000,000,04 .000,000,23 .000,001,00 .000,002,32 .000,004,1	***	the available significant ures.
50 60 70 80 90	9.654,180,19 9.624,113,15 9.594,186,36 9.564,398,93 9.534,750,01	9.654,173,7 9.624,103, <u>8</u> 9.594,173, <u>7</u> 9.564,382 9.534,729	departures from g* are underlined below	.000,006,5 .000,009,3 .000,012,6 .000,016 .000,021	88 me 88 g	zero within the annoper of significations
100 200 300 400 500	9.505,238,75 9.217,512,92 8.942,656,38 8.679,912,89 8.428,581,04	9.505,213 9.217,41 <u>5</u> 8.942,45 8.679,59 8.428,18	9.217,41 <u>4</u> 8.942,44 8.679,5 <u>4</u> 8.428, <u>04</u>	.000,026 .000,098 .000,20 .000,32 .000,40	.000,099 .000,21 .000,38 .000,54	000,001 000,01 000,05 000,14
600 700 800 900 1000	8.188,009,42 7.957,592,42 7.736,766,50 7.525,006,62 7.321,823,24	8.187,6 <u>1</u> 7.957, <u>39</u> 7.737,0 7.526,2 7.324,6	8.187,24 7.956,59 7.735,6 7.524,0 7.320,7	.000,39 .000,20 000,3 001,2 002,8	.000,76 .001,00 .001,2 .001,0	.000,37 .000,80 .001,4 .002,2 .003,9

Table N-VI. Values of the acceleration of gravity for various altitudes computed from three different equations as indicated, and the differences between these values of the acceleration of gravity.

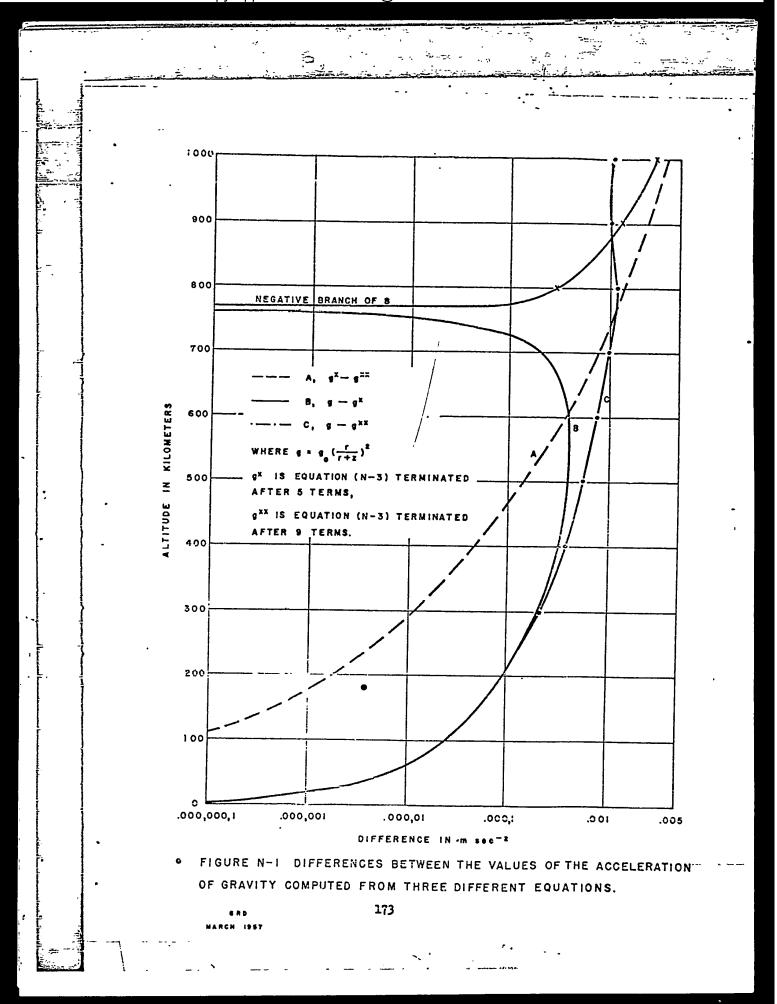
NOTE: Underlined numbers in Column  $g^*$  indicate figures of questionable significance.

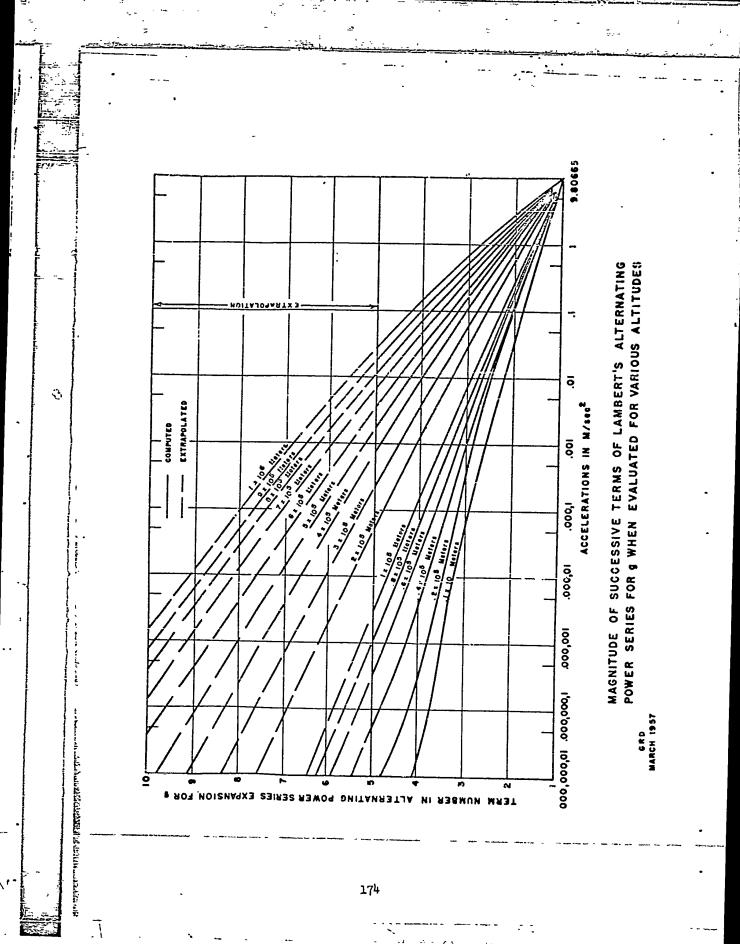
Underlined numbers in Column g\*\*indicate figures differing from Column g\*.

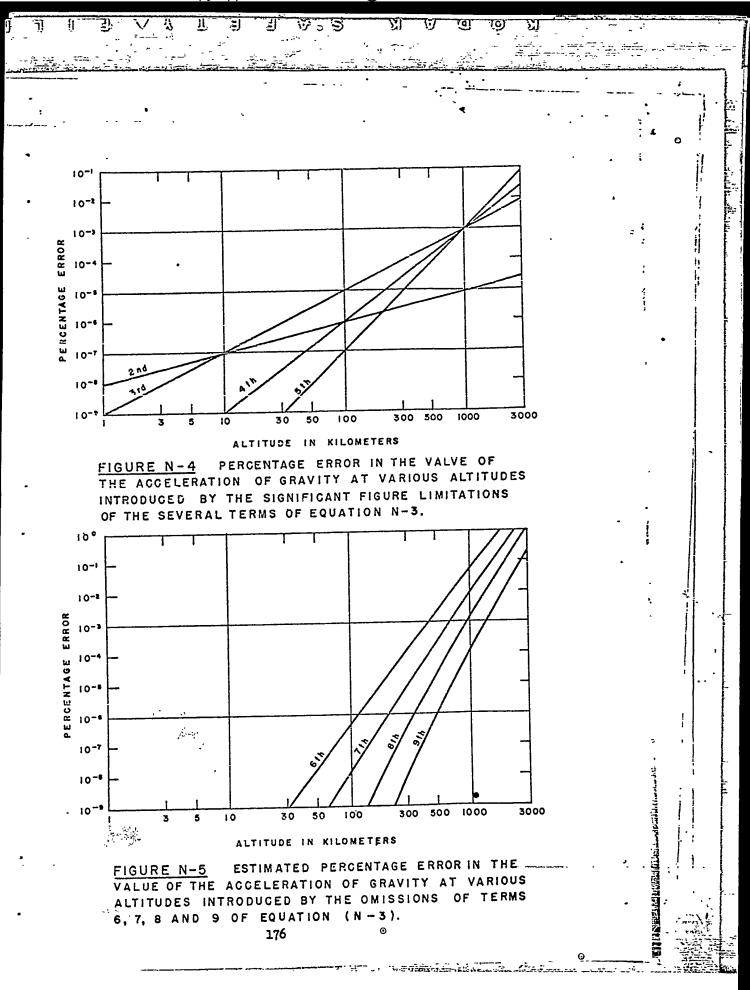
#### 7. Conclusions

- a. For most engineering purposes, the adjusted inverse-square-law function for g provides adequate accuracy.
- b. For the standard atmosphere, and for future editions of this MODEL, the values of g should be computed on the basis of an expanded version of Eq. (N-3) in which a minimum of three, and preferably five, additional terms are employed, and in which sufficient additional significant figures are provided for the various limiting coefficients, particularly coefficients of terms 3, 4, and 5.

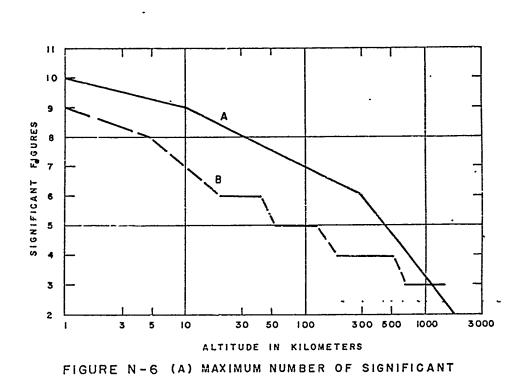
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(B) THE MAXIMUM NUMBER OF SIGNIFICANT FIGURES OF THE VALUE OF g AT VARIOUS ALTITUDES, COMPUTED FROM THE ADJUSTED INVERSE SQUARE LAW, WHICH ARE IN AGREEMENT WITH VALUES COMPUTED FROM EITHER THE 5 TERM OR 9 TERM VERSION OF EQUATION (N-3).

FIGURES AVAILABLE FROM THE EXISTING 5 TERM VERSION

OF EQUATION N-3, FOR VARIOUS ALTITUDES.

#### APPENDIX O

#### Scale Height

#### 1. Geometric Scale Height

First Concept - Scale height is equal to the height above any reference altitude at which the atmospheric pressure falls to 1/e of the pressure at the reference altitude in a constant gravity, isothermal atmosphere.

In a manner analogous to the development of Eq. (15) in terms of H (Section 3.2.1), the following equation is developed in terms of Z:

$$\hat{\mathcal{L}} = \frac{\frac{M_o}{P_b}}{Z_b} = \frac{\frac{M_o}{R^*} \int_{Z_b}^{rZ} \frac{g\dot{u}Z}{T_M}.$$
 (0-1)

For the case of an isothermal layer in a constant gravity atmosphere, Eq. (0-1) upon integration leads to

$$P = P_b \text{ exponential } -\frac{g_0 M_C^*}{R^* (T_M)_b} (Z - Z_b). \qquad (0-2)$$

It is noted that in a constant gravity atmosphere:

$$\frac{R^*(T_M)_b}{g_0M_0} = (H_g)_b , \qquad (0-3)$$

and it follows that

$$P = P_b \text{ exponential } -\frac{(z - z_b)}{(H_B)_b}.$$
 (0-4)

For the case that

$$(Z - Z_b) = (H_s)_b , \qquad (O-5)$$

Eq. (0-4) simplifies to

$$P = P_b e^{-1} = P_b/e$$
 . (0-6)

It appears, therefore, that in a constant gravity atmosphere and in a layer of constant T<sub>H</sub>, the scale height at any reference level is the increment in geometric altitude required for the pressure to fall to 1/e of the value at the reference level. Since this MODEL does not assume constant gravity, the above concept does not apply rigorously in these tables. In the special case, where sea level is the reference altitude the same concept would apply but only if the isothermal layer is assumed to extend down to there, and only for a constant gravity atmosphere.

Second Concept - In an atmosphere of constant g and constant  $T_M$ , the scale height at any altitude Z, is equal to the total mass of air in a unit column extending upward from that altitude to infinity, divided by the density at the reference altitude.

From Eq. (33) one obtains  $\frac{P}{P} = \frac{\rho}{Q} \cdot \frac{T_{M}}{(T_{CA})}. \qquad (0-7)$ 

In a constant  $T_M$  atmosphere,  $T_M = (T_M)_b$  and thus,

$$\frac{P}{\Gamma_{\rm b}} = \frac{\rho}{\rho_{\rm b}} . \tag{0-8}$$

Equation (0-2) may then be rewritten as

$$\rho = \rho_{b} \text{ exponential } - \frac{g_{o}M_{o}}{R^{*}(T_{M})_{b}}. \tag{0-9}$$

The total mass in a unit column from the reference level to infinity is:

$$\int_{Z_b}^{\infty} \rho dZ = \rho_b \int_{Z_b}^{\infty} \text{exponential} - \frac{g_0 M_0}{R^* (T_M)_b} \quad (z - Z_b)$$
 (C-10)

$$= \rho_{b} \left[ \frac{R^{*}(T_{M})_{b}}{-g_{o}M_{o}} \right] \left[ \text{exponential} - \frac{g_{o}M_{o}}{R^{*}(T_{M})_{b}} (z - z_{b}) \right]_{Z_{b}}^{\infty}$$
(0-10a)

$$= \rho_{\rm b} \left[ \frac{R^*(T_{\rm M})_{\rm b}}{-g_{\rm o} M_{\rm o}} \right] \left[ e^{-\varpi} - e^{\rm o} \right]$$
 (0-10b)

$$= \rho_{\rm b} \cdot \frac{R^*(T_{\rm M})_{\rm b}}{g_{\rm o}^{\rm M}_{\rm o}} \tag{0-l0c}$$

Since  $\frac{R^*(T_M)_b}{g_0M_c}$  = scale height at  $H_b$  in a constant gravity atmosphere, it

follows that

$$(H_g)_b = \frac{1}{\rho_b} \int_{\underline{Z}_b}^{\infty} \rho dZ. \qquad (0-11)$$

Thus the assertion of Concept 2 is demonstrated.

Third Concept - In a constant-g, constant- $T_M$ , constant-M atmosphere, the scale height at any altitude is equal to the total number of particles in a column of unit cross section extending from a reference level to infinity, divided by the number density at that altitude.

From Eqs. (26) and (27) of Sections 5.2.1 and 5.3.1, respectively, it follows that:

$$n = \rho \frac{N}{M!} \tag{0-12}$$

but

$$\frac{M'}{N} = m \tag{0-13}$$

where m = the mass of 2 single air particle.

Thus 
$$\rho = n m \qquad (0-14)$$

Thus it follows directly from Eq. (0-11) that

$$\left(\mathbf{H}_{\mathbf{s}}\right)_{\mathbf{b}} = \frac{1}{\mathbf{n}_{\mathbf{b}} n_{\mathbf{b}}} \int_{\mathbf{Z}_{\mathbf{b}}}^{\mathbf{co}} \rho d\mathbf{Z}, \qquad (0-16)$$

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$$\left(\mathbf{H}_{s}\right)_{b} = \frac{1}{\mathbf{n}_{l} m_{b}} \int_{\mathbf{Z}_{b}}^{\infty} n m d z. \tag{0-17}$$

The right-hand side of this equation would not strictly equal the total number of atmospheric particles in the column, unless the molecular weight were constant. Thus, for the assertion of the third concept to be right-ously correct, it was necessary to make the restriction of constant molecular weight in addition to the restrictions made in the first and second concepts. With this constant-M restriction, Eq. (0-17) becomes

$$\left(\mathbf{H}_{\mathbf{S}}\right)_{\mathbf{b}} = \frac{1}{\mathbf{n}_{\mathbf{b}}} \int_{\mathbf{Z}_{\mathbf{b}}}^{\mathbf{C}} \mathbf{n} d\mathbf{Z}, \qquad (0-18)$$

and the assertion is demonstrated. It is noted that a corollary to the third concept is that scale height is the length of the unit column necessary to enclose all the atmospheric particles normally present in an infinitely long unit column, extending vertically above the reference altitude, when these particles are compressed to the number density at the reference level. Hence, this quantity is the basis for computing reduced thickness of the atmosphere. Such computations are limited by the fact that constant gravity, constant T<sub>M</sub>, and constant molecular weights are assumed in the derivation of the expression.

#### 2. Geopotential Scale Height

Gcopotential scale height was defined in Section 4.1.3 of this paper as

$$\mathbf{H_g}^{\, *} \; = \frac{\mathbf{GM_o}}{\mathbf{R^*T_M}} \; .$$

In terms of this property the several concepts developed above do not have the restriction of a constant gravity atmosphere. Thus Eq. (15) of Section 3.2.2 may be rewritten as

$$P = P_b \text{ exponential } - \frac{GM_o}{R^*(T_M)_b} (H - H_b). \tag{0-19}$$

For a geopetential altitude increment equal to the geopetential scale height

$$(H - H_b) = \frac{R^*(T_M)_b}{GM_o} = H_g^*$$
,

and hence Eq. (0-19) reduces to

$$P = P_b/c. (0-19a)$$

Note that no assumption of constant gravity is made, only constant  $T_{\rm M}.$  Hence, a revision of Concept 1, eliminating the constant gravity restrictions, will apply rigorously in this MODEL in isothermal layers. For example,  $H_{\rm S}'$  at 11 km' is 6.341,615,82 x 10 $^{9}$  m'. Thus, at 17.341,615,82 km', the pressure will be  $P_{11}/e$ , where  $P_{11}$  is the pressure at 11 km. At 14 km',  $H_{\rm S}'$  has the same value; hence at 20.341,615,82 x 10 $^{9}$  m' altitude, the pressure will be  $P_{11}/e$ . The geometric altitude increment, however, will be different in the two instances, accounting for the effect of variable g on the pressure.

In geopotential form, Eq. (0-10) may be rewritten as

$$\int_{H_b}^{\infty} \rho^{dH} = \rho_b \int_{H_b}^{\infty} \exp \operatorname{exponential} - \frac{GM_o}{R^*(T_M)_b} (H - H_b). \quad (0-20)$$

By analogy this reduces to

$$(H_{s}')_{b} = \frac{1}{\rho_{b}} \int_{H_{b}}^{\infty} \rho dH.$$
 (0-21)

This equation and concept rigorously apply to isothermal layers of this MODEL.

Equation (0-16) is converted by analogy to

$$(H_{g'})_{b} = \frac{\int_{H_{b}}^{\infty} \rho dH}{n_{b}m_{b}}$$
 (0-22)

If constant molecular weight is assumed, this equation becomes:

$$(H_{g'}) = \frac{1}{n_{b}} \int_{H_{b}}^{\infty} n dH.$$
 (0-25)

This equation would provide a better basis for computing reduced thickness

#### APPENDIX P

#### More Accurate Method for Computing Geopotential in this Model

#### 1. Adjusted Classical Approach

Equation (2d) of this paper indicates the rigorous relationship between geopotential H, geometric altitude Z, and the acceleration of gravity g to be

$$H = \frac{1}{G} \int g dZ. \qquad (P-1)$$

When g is expressed by the classical, inverse-square law, adjusted for 45° 32' 40" latitude,

$$g = g_0 \left(\frac{r}{r+Z}\right)^2, \qquad (P-2)$$

the expression for geopotential becomes

$$H = \frac{g_0}{G} \left( \frac{rZ}{r+Z} \right), \tag{P-3}$$

where  $g_0$  and r have the values 9.80665 m  $\sec^{-2}$  and 6,356,766 m, respectively, as indicated in Section 2.1.

#### **€.** Lambert Series Method

In Appendix N, another expression for g in terms of Z for latitude 45° 32' 40" was developed from Lambert's general alternating power series. 38 This specific expression is

$$g = c_1 - c_2 Z + c_3 Z^2 - c_4 Z^3 + c_5 Z^4 - \dots$$
 (P-4)

where

$$c_1 = 9.806,65 \text{ (exact)}$$
 $c_2 = 30,854.18_8 \times 10^{-10}$ 
 $c_3 = 725.38_1 \times 10^{-15}$ 
 $c_4 = 15.16_{89} \times 10^{-20}$ 
 $c_5 = .296_{96} \times 10^{-25}$ 
 $m \cdot sec^{-2}$ 
 $m^{-2} sec^{-2}$ 
 $m^{-3} sec^{-2}$ 
 $m^{-3} sec^{-2}$ 

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When this expression for g is introduced into Eq. (P-1) the expression for H becomes

$$H = \frac{1}{G} \left[ c_1 \int_0^Z dZ - c_2 \int_0^Z ZdZ + c_3 \int_0^Z Z^2 dZ - c_4 \int_0^Z Z^3 dZ + c_5 \int_0^Z Z^4 dZ - \dots \right]$$
 (P-5)

where H is in standard geopotential meters.

Performing the indicated integration one obtains

$$H = \frac{c_1}{G} z - \frac{c_2}{2G} z^2 + \frac{c_3}{3G} z^3 - \frac{c_4}{4G} z^4 + \frac{c_5}{5G} z^5 - \dots,$$
 (P-6)

where the coefficients of the various powers of Z have the following numerical values:

$$\frac{c_1}{G} = \frac{9.806,65}{9.806,65}$$
= 1.0 exact
$$\frac{c_2}{2G} = \frac{30,854.188 \times 10^{-10}}{2 \times 9.806,65}$$
= 1,573.125<sub>78</sub> × 10<sup>-10</sup>

$$\frac{c_3}{3G} = \frac{725.3_{01} \times 10^{-15}}{3 \times 9.806,65}$$
= 24.65<sub>61</sub> × 10<sup>-15</sup>

$$\frac{c_4}{4G} = \frac{15.16_{89} \times 10^{-20}}{4 \times 9.806,65}$$
= .386,6<sub>99</sub> × 10<sup>-20</sup>

$$\frac{c_5}{5G} = \frac{.296_{96} \times 10^{-25}}{5 \times 9.806,65}$$
= .006,05<sub>63</sub> × 10<sup>-25</sup>

Hence one obtains

$$H = Z - 1,573.125_{78} \times 10^{-10}Z^{2} + 24.65_{61} \times 10^{-15}Z^{3}$$
$$- .386,6_{99} \times 10^{-20}Z^{4} + .006,05_{63} \times 10^{-25}Z^{5}... \qquad (P-7)$$

(where the exponents have been selected for convenience when 2° is expressed in units of 105 meters).

Evaluating the five defined terms of Eq. (P-7) for various altitudes yields the data presented in Table P-I. An examination of the logarithms of successive terms of the series evaluated for particular altitudes shows that the absolute magnitudes of successive terms fall off very nearly at a constant rate, or, in other words, the logarithmic decrement of successive terms is very nearly constant. Examples of this nearly constant logarithmic decrement,  $\Delta$  log, are given for 1,000, 300, and 100 km.

Alt.	1,000,000 m	300,000 m	100,000 m
Term #	$Log_{10}$ Term $\Delta log$	Log <sub>10</sub> Term Δlog	Log <sub>10</sub> Term Δlog
1	6.000,00	5.477,12	5.000,00
2	.803,24 5.196,76	1.326,11 4.151,01	1.803,24 3.196,76
3	.804, <u>84</u> 4.391, <u>92</u>	2.823,29	1.804.84
14	.804, <u>55</u>	1,327,43	1.391,92
•	-805.16	1.495,86	9.587,37
5	2.782,21	.164,22	1.80 <u>5,16</u>

NOTE: Underline indicates non-significant digits.

# 3. Extension of the Lambert Series

The departure of the logarithmic decrement from linearity is less than one half of one percent over the five available terms for the altitudes discussed. On the average, the differences between the logarithms of successive terms increase very slightly with increasing term number. It is not unreasonable to assume that this pettern of logarithmic decrement with slowly increasing differences might continue for a considerable number of additional terms in the series. Employing this pattern, the values of the ninth term of Eq. (P-7) for 1,000, 500, and 100 km are 3.6 x 10<sup>-1</sup>, 4.9 x 10<sup>-0</sup>, and 3.6 x 10<sup>-10</sup>, respectively, in standard geopotential meters.

Estimated values of the 6th, 7th, 6th, and 9th terms of Eq. (P-7) for various altitudes may also be determined graphically by plotting the logarithms of the various terms as functions of term number, and connecting those points corresponding to each specific altitude as in Fig. P-1. These lines are then extended linearly to higher term numbers as in the dashed line portion of Fig. P-1. The estimated values of terms 6, 7, 8, and 9 of Eq. (P-7) determined graphically on a figure three times as large as Fig. P-1 are given in Table P-II. Graphically determined values of the ninth term of Eq. (P-7) for altitudes of 1,000, 300, and 160 km differ from the three computed values given above by less than 10 per cent.

A replotting of the data of Table P-I in terms of the value of each

term of Eq. (P-7) as a function of altitude is given in Fig. P-2. The estimated values for the 6th, 7th, 8th, and 9th terms of the equation come from Fig. P-1. Figure P-2 clearly shows the contribution which each term in the series makes to the value of geopotential of a given geometric altitude. Figure P-2 demonstrates that for errors in geopotential of less than .1 m', the five term version of Eq. (P-7) may be used only to altitudes of about 280 km, neglecting the possible limitations due to significant fig-

#### 4. Comparison of the Three Methods

The values of geopotential in stendard geopotential meters for various geometric altitudes are given in Table P-III. Values designated by H are computed from the simple Eq.  $(P-\bar{\jmath})$ . Values designated by H\* are computed from the five defined terms of Eq. (P-7). Values designated by H\*\* are those resulting from the estimated nine-term version of Eq. (P-7). The values of the differences H - H\*, H - H\*\*, and H\* - H\*\* are also given in Table P-III. The difference H - H\*\* is of particular interest, since it indicates the amount of error in geopotential altitude incurred by using the simple Eq.  $(P-\bar{\jmath})$  instead of the nine-term version of Eq.  $(P-\bar{\jmath})$ . (Below 100 km altitude the error is less than 0.1 m'.)

#### 5. Limitation of the Five Term Lambert Series Due to Numbers of Terms

Because of the increase of centrifugal acceleration with altitude which is not accounted for in Eq. (P-3), the departure between the value of H from Eq. (F-5) and the value from Eq. (F-7) is expected to increase with altitude. The reversal of the trend resulting in smaller departures (i.e. smaller values in H - H\*) above 800 km suggests the inadequacy of the five-term version of Eq. (P-7). The difference H - H\*\* involving the nine-term version of Eq. (P-7) continues to increase to altitudes well over 1000 km. A graph of the various differences is given in Fig. P-3.

#### 6. Limitations of the Five Term Lambert Series Due to Significant Figures

An analysis of the values and number of significant figures of terms 2, 3, 4, and 5 of Eq. (P-7) as listed in Table P-I indicates the limitations which the number of significant figures of each term place upon the computed value of geopotential. The results of this analysis are presented in Fig. P-4. Below 10 km altitude, the number of significant figures in term number 2 is seen to limit the accuracy of Eq. (P-7). From 10 km to about 3,200 km altitude, term number 3 limits the accuracy of the equation, provided a sufficient number of terms is employed so that the number of terms does not limit the accuracy at some altitude below 3,200 km.

#### 7. Combined Limitations of the Lambert Series

The minimum numerical error obtainable with the existing five-term version of Eq. (P-7) is given as the three-segment curve A of Fig. P-5.

Segment a represents the limitation due to significant figures of term 2; segment b represents the limitation due to significant figures of term 3; while segment c represents the limitation due to the termination of the scries after term 5. Line B of that same graph represents the minimum numerical error incurred in using the simple equation for geopotential, Eq. (P-3). This error is determined from the values of H - H\*\*. The difference between these two curves (given more accurately by values of H - H\* in Table P-III) shows that for altitudes between 10 and 500 km, an improvement of only one significant figure in geopotential altitude is obtained by switching from Eq. (P-3) to the presently available form of Eq. (P-7).

### 8. Requirements Which the Extended Lambert Series Must Meet

In order to obtain the ten significant figure accuracy desirable for standard atmosphere computations at altitudes of 300, 500, and 1,000 km; three, four, and eight additional terms, respectively, must be developed for Eq. (P-7). Also, the following numbers of significant figures should be available for the several coefficients:

Alt.	300 km	500 km	1,000 km
Term #	Number of Sig. Fig.	Number of Sig. Fig.	Number of Sig. Fig.
2 3 4 5	9 7 6 5	9 8 7 6	10 9 8 7
6 7 8 9	3 2 1	5 4 2 1	7 6 5 4
10 11 12 13	•		3 2 2 1

These requirements reflect back directly upon Lambert's general expression for g as a function of Z and  $\emptyset$ ; i.e.,

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$$g = c_1 - (a_2 + b_2 \cos 2\phi)Z + (a_3 + b_3 \cos 2\phi)Z^2$$

$$- (a_4 + b_4 \cos 2\phi)Z^3 + (a_5 + b_5 \cos 2\phi)Z^4$$

$$- \dots + \dots \qquad (ref. 38)(P-8)$$

To meet the above requirements for latitude 90°, the coefficients  $a_2$ ,  $a_3$ ,  $a_4$ , ...  $a_n$  and  $b_2$ ,  $b_5$ ,  $b_4$ , ...  $b_n$  of Eq. (P-8) must have numbers of significant figures graphically estimated to be the following:

Alt.	300 km	500 km	1,000 km
n	a <sub>n</sub> b <sub>n</sub>	an bn	a <sub>n</sub> b <sub>n</sub>
2 3 1 5	9 7 7 5 6 3 5 3	9 7 8 6 7 4 6 4	10 7 9 7 8 5 7 5
6 7 8 9	3 1 2 1	5 3 4 2 2 1 1	7 5 6 4 5 4 4 3
10 11 12 13			3 3 2 2 2 2 1 1

To meet standard atmosphere requirements at latitude 45° 52' 40", the number of significant figures required for  $b_n$  would be one to two less than required for the case when  $\phi=90^\circ$ . In any case,  $b_n$  must have enough significant figures so as not to invalidate the accuracy of  $a_n$ .

#### 9. Conclusions

This analysis is strictly mathematical and does not consider whether it is physically possible to obtain the required number of terms or the necessary accuracy in Eq. (P-4) or Eq. (P.8). If no substantial improvement of Eq. (P-7) is physically possible through a better expression for the acceleration of gravity in Eq. (P-4) or Eq. (P-8) and if one must resort to arbitrary definitions as in the standard sea level pressure, then it is suggested that Eq. (P-2) for g be retained by definition, in which case geopotential is given by the simple Eq. (P-3), sufficiently accurate for most engineering purposes. Only a study of Lambert's unpublished method for the development of Eq. (P-8) will suggest the course to follow.

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•	5th Term	. οια, ααο, σοο, οαο, σ . αια, ααο, αοι, 892, <u>5</u> . αιο, ααο, οδο, 5 <u>63</u> . αιο, αιι, 978, <u>02</u> . αιο, οι 14, 71 <u>6, 8</u> . αιο, οε 2, ο <u>15</u> , 8	25.5.5.5.5.5.5.5.5.5.5.5.5.5.5.5.5.5.5.	දු		udes as terms in m <sup>l</sup> )	are
		બ	.000,159,259 .000,470,94 .001,017,88 .001,984,53	.006,056,2 .193,802 1.474,65 6.201,6 18.925,2	47.094 101.788 198.452 357.62 605. <u>63</u>	eometric Altitude: rm. (Value of ter	ignificance, but
	4th Term	.000,000,003,866, <u>9</u> .000,002,416, <u>87</u> .000,038,669, <u>9</u> .000,618,718 .003,132,26	.024,168,7 .050,116,2 .092,846,7 .158,39 <u>1.9</u> .253,71 <u>2</u>	. 386,699 6.187,18 31.322,6 98.994,9 241.687	501.162 928.4 <u>64</u> 1,587.9 <u>19</u> 2,577. <u>13</u> 3,866. <u>99</u>	the First Five Terms of Eq.(P-7)for Various Geometric Altitudes as Indicated by the Value of the First Term. (Value of terms i	lined figures are beyond the limit of significance, but are or smoothness.
	3rd Te.m	.000,024,65 <u>6</u> .003,082, <u>01</u> .024,65 <u>6,1</u> .197,24 <u>9</u> .665,7 <u>15</u> .1.577, <u>99</u>	3.082, <u>01</u> 5.325, <u>72</u> 8.457, <u>04</u> 12.623, <u>2</u> 17.97 <u>4,</u> 2	24.656,1 197.24 <u>9</u> 665.71 <u>5</u> 1,577.9 <u>9</u> 3,082.0 <u>1</u>	5, 325.72 8, 457.04 12, 623.2 17, 974.3 24, 656.1	irst Five Terms of E Indicated by the V	The underlined figures are tearried for smoothness.
	2nd Term	157, 512, 58 5.932, 814, 45 15.731, 25 <u>7, 8</u> 62.925, 03 <u>1, 2</u> 141.581, 320	393.281,445 566.325,281 770.831,632 1,006.800,499 1,274.231,88	1,573.125,78 6,892.503,12 14,158.132,0 25,170.012,5 39,328.144,5	$56,632.528,1$ $77,083.16\overline{2},2$ $100,680.04\overline{9},\overline{2}$ $127,423.18\overline{8}$ $157,312.5\overline{18}$	Values of	NOTE: The underl
	lst Term	1,000 5,000 10,000 20,000 70,000	% % % % % % % % % % % %	100,000 200,000 300,000 400,000	600,000 700,000 800,000 900,000	Table P-I.	
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9th Term		.000,000,000,000 .000,000,000,000 .000,000,	.000,000,000,36 .000,000,160 .000,005,1 .000,089	.003,4 .014 .042 .14 .36	7)for Various ig. ?-1).
8th Term	.000,000,000,000. .000,000,000,000.	.000,000,000,082 .000,000,000,33 .000,000,001,35 .000,000,003,4	.000,000,024 .000,054 .000,125 .001,43	.033 .135 .34 .30 2.4	8, and 9 of Eq.(P-'panded version of F
7th Term	.000,000,000,000. .000,000,000,018 .000,000,000,27	.000,000,011 .000,000,04 .000,000,12 .000,000,28	.000,001,5 .000,18 .002,7 .024	.4 1.2 2.8 7.0 15.	Estimated Values of Terms 6, 7, 8, and 9 of Eq.(P-7) for Various Altitudes (estimated from expanded version of Fig. P-1).
6th Term	.000,000,000,000. .000,000,000,000. .000,000,	.000,000, 4,000,000, 6,000,011,2 .002,023	.000,095 .006 .064 .40	4.4 11.2 23. 83. 95.	Table P-II. Estimated Va Altitudes
Alt.	1 ~ 0 8 8 9	8888	100 200 500 500 500	2,000 1,000 1,000 1,000	Table

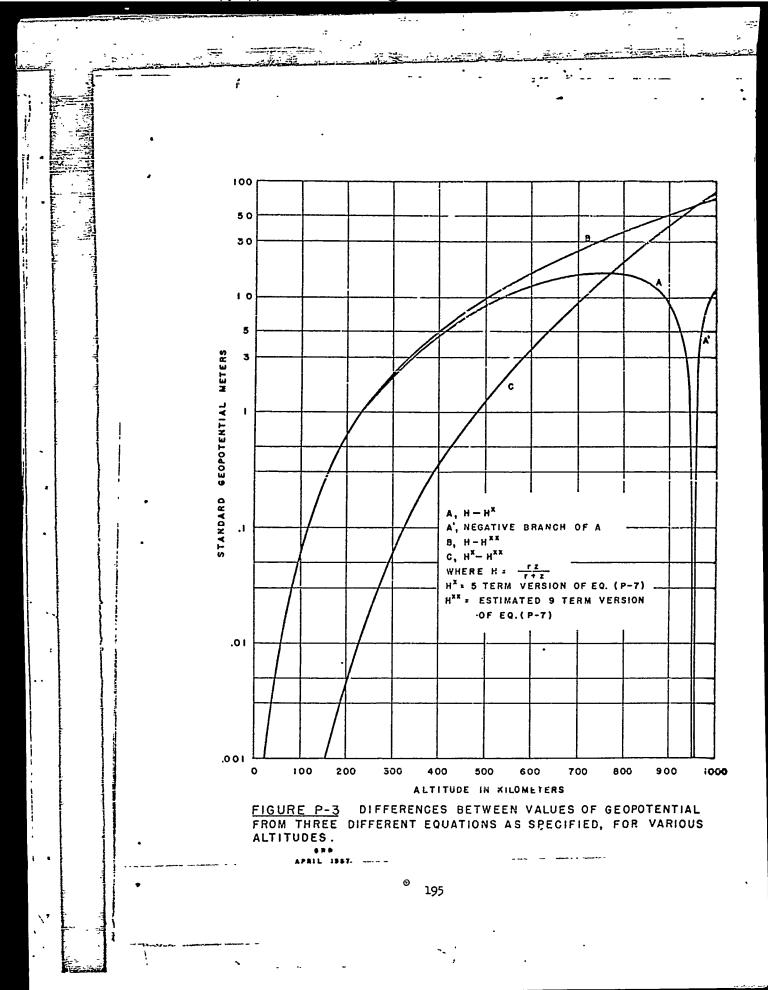
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**H - H	.000,000 .000,008 .000,078 .000,68 .000,60	.010,9 .018,9 .030,1 .04,5	.087,4 .685 2.26 5.22 9.85	16.7 26.0 50.0 50.0 50.0 50.0	for Various Different Equ Geopotentia	the degree of deparare depressed.
*H + H	.000,000 .000,008 .000,078 .000,68	.010,9 .018,9 .050,1 .045	889. 448 89. 448 89. 448	12.7 15.9 15.7 8.4 -1.5	ntial Meters from Three Bee Velues of	cates the de   H** are dep   not nore th
11**, 9 Terms of Eq. (P-7)	4	Game as H*	\$65,477.6 376,314.8 463,525.4	548,235.1 630,537.1 710,53.19 788,325.1 864,00 <sub>0</sub> .1	andard Geopote ! 40" Computed ea Between Tho	llues of H indi Lluen of H* and are reliable to
H*, 5 Terms of Eq. (P-7)	999.842,712,0 4,996.070,265 9,984.293,36 19,937.271,60 29,859.081,27 39,749.868,0	49,609.776,6 59,438.950,8 69,237.533,6 79,005,667 88,74,5.492	98, 451.149 195, 898.75 286, 477.72 376, 315.2 463, 531.1	548,239.1 630,547.2 710,558.4 788,371.6 864,082.2	Values of Geopotential in Standard Geopotential Meters for Various Geometric Altitudus at Latitude 45° 32' 40" Computed from Three Different Equations an Indicated, and the Differences Batween Those Velues of Geopotential, also in Standard Geopotential, Meters.	The underlined portion of values of H indicates the degree of departure from values of H* and H**.  Nonsignificant figures in values of H* and H** are depressed.  The difference tabulations are reliable to not more than three significant figures.
$H = \frac{rZ}{r+Z}$	999.842,712,0 4,996.070,273, <u>5</u> 9,984.293,4 <u>38</u> 19,937.272,27 <u>8</u> 29,859.08 <u>3,61</u> 39,749.8 <u>73,</u> 50	149,609.787.52 59,438.969,72 69,237.563,65 79,005.723,87 88,743.556,01	98,451. <u>237,0</u> 193,89 <u>9.431,5</u> 286,47 <u>9.921,3</u> 376,320.032,3 463,53 <u>9.662,8</u>	548,251.817,0 630,565.093,6 710,574.123,6 788,380.030,4 864,070.701	Table P-III. Values of Altitudes Indicated, Standard	NOTE: The underline values of H* Nonsignifican The difference
Alt.	42883		100 B 60 B	1, 800 1, 900 1, 900 1,	Table	
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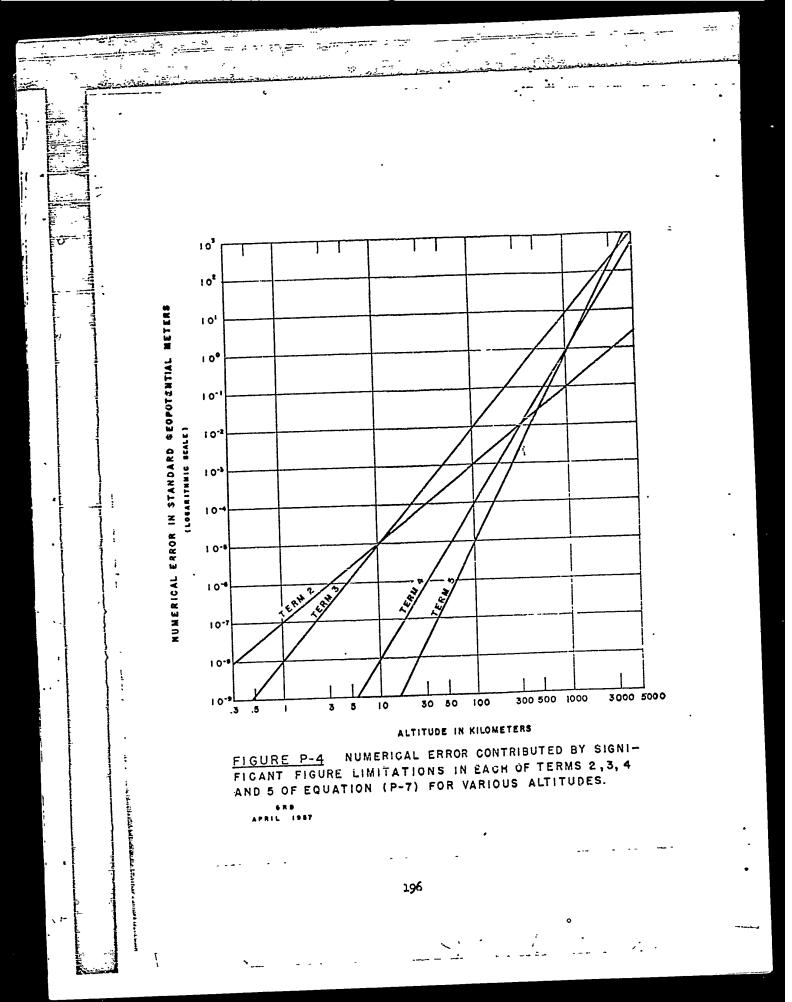
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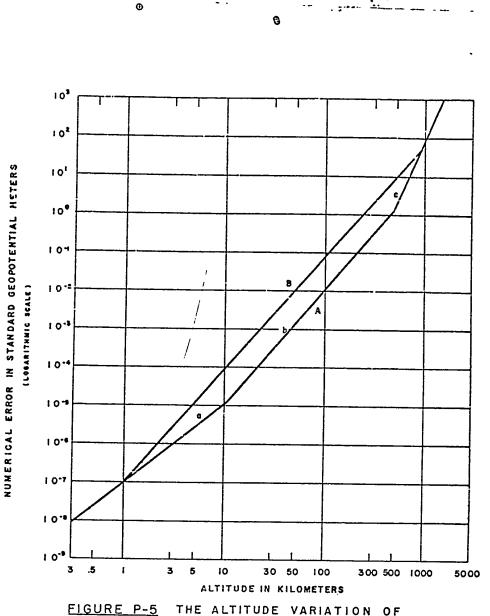


FIGURE P-5 THE ALTITUDE VARIATION OF (A), MINIMUM NUMERICAL ERROR ASSOCIATED WITH THE EXISTING 5 TERM VERSION OF EQUATION P-7 FROM BOTH SIGNIFICANT FIGURE CONSIDERATIONS, AND A LACK OF SUFFICIENT NUMBER OF TERMS.

(B), MINIMUM NUMERICAL ERROR ASSOCIATED WITH THE USE OF THE ADJUSTED VERSION OF  $H = \frac{rz}{r+z}$  AT VARIOUS ALTITUDES AT 45° 32' 40" L.

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## LIST OF AIR FORCE SURVEYS IN GEOPHYSICS (Unclassified)

	(2.12.11.22.11)		ο.	Security
Number	Title	Author	Date	Class.
1		W. K. Widger, Jr.	Mar 52	S-RD
2	Methods of Weather Presentation for Air Defense Operations	W. K. Widger, Jr.	Jun 52	С
3	Some Aspects of Thermal Radiation From the Atomic Bomb	R. M. Chapman	Jun 52	S
4	Final Report on Project 8-52M-1 Tropopause	S. Coroniti	Jul 52	S
5	Infrared as a Means of Identification	N. Oliver J. W. Chamberlain	Jul 52	S
6	Heights of Atomic Bomb Results Relative to Basic Thermal Effects Produced on the Ground	R. M. Chapman G. W. Wares	Jul 52	S-RD
7	Peak Over-Pressure at Ground Zero From High Altitude Bursts	N. A. Haskell	Jul 52	S
8	Preliminary Data From Parachute Pressure Gauges. Operation Snapper. Project 1.1 Shots No. 5 and 6	N. A. Haskell	Jul 52	S-RD
9	Determination of the Horizontal	R. M. Chapman M. II. Seavey	Sep 52	S
10	Soil Stabilization Report	C. Moline x	Sep 52	υ
n	Geodesy and Gravimetry, Preliminary Report	R. J. Ford, Maj., USAF	Sep 52	S
12	The Application of Weather Medification Techniques to Problems of Special Interest to the Strategic Air Command	C. E. Anderson	Sep 52	S
13	Efficiency of Precipitation as a Scavenger	C. E. Anderson	Aug 52	S-RD
14	Forecasting Diffusion in the Lower Layers of the Atmosphere	B. Davidson	Sep 52	С
15	Forecasting the Mountain Wave	C. F. Jenkins	Sep 52	U
16	A Preliminary Estimate of the Effect of Fog and Rain on the Peak Shock Pressure From an Atomic Bomb	J. H. Healy H. P. Gauvin	Sep 52	S-RD

<sup>\*</sup>Titles that are omitted are classified.

	Numbe	Title	Author	Date	Securi
ì	17	Operation Tumbler-Snapper Project 1.1A. Therma Radiation Measurements With a Vacuum Capacitor Microphone	l woo	Sep 52	Class C-RD
	18	Operation Snapper Project 1.1, The Measurement of Free Air Atomic Blast Pressures	J. O. Vann, Lt Col., USAF N. A. Haskell	Sep 52	S-RD
	19	The Construction and Application of Contingency Tables in Weather Forecasting	E. W. Wahl R. M. White H. A. Salmela	Nov 52	U
	20	Peak Overpressure in Air Due to a Deep Under-	N. A. Haskell	Nov 52	s
	21	Slant Visibility	R. Penndorf B. Goldberg D. Lufkin	Dec 52	Ü
	22	Geodesy and Gravimetry	R. J. Ford, Maj., USAF	Dec 52	S
	23	Weather Effect on Radar	D. Atlas V. G. Plank W. H. Paulsen A. C. Chmela J. S. Marshall T. W. R. East K. L. S. Gunn	Dec 52	U
	24	A Survey of Available Information on Winds Above 30,000 Ft.	C. F. Jenkins	Dec 52	Ü
	25	A Survey of Available Information on the Wind Fields Between the Surface and the Lower Stratosphere	W. K. Widger, Jr.	Dec 52	U
	26		A. L. Aden	D **	_
	0.7	•	L. Katz	Dec 52	S
	27		N. A. Haskell	Dec 52	s
	•	1	R. H. Chapman G. W. Wares M. H. Scavey	Dec 52	S-RD
	29 <i>l</i>	A Note on High I I TO I I	J. Kuettner	Dec 52	Ū

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y	Number	Title	Author	Date	Security Class.
	30	Results of Controlled-Altitude Balloon Flights at 50,000 to 70,000 Feet During September 1952	T. O. Haig Maj., USAF R. A. Craig	Feb 53	<b>U</b> -
	31	Conference: Weather Effects on Nuclear Detenation	B. Grossman, Ed.	Feb 53	S-RD
	32	Operation IVY Project 6.11. Free Air Atomic Blast Pressure and Thermal Measurements	N. A. Haskell P. R. Gast	Mar 53	S-RD
	33	Variability of Subjective Cloud Observations - 1	A. M. Galligan	Mar 53	U
	34	Feasibility of Detecting Atmospheric Inversions by Electromagnetic Probing	A. L. Aden	Mar 53	U
	35	Flight Aspects of the Mountain Wave	C. F. Jenkins J. Kuettner	Apr 53	U
	36	Report on Particle Precipitation Measurements Performed During the Buster Tests at Nevada	A. J. Parziale	Apr 53	S-RD
*	37	Critical Envelope Study for the XB-63, 13-52A, and F-89	N. A. Haskeil M. H. Scavey R. M. Chapman	Apr 53	S
,	38	Notes on the Prediction of Overpressures From Very Large Thermo-Nuclear Bombs	N. A. Hoskell	Apr 53	S
	39	Atmospheric Attenuation of Infrared Oxygen Afterglow Emission	N. J. Oliver J. W. Chamberlain	Apr 53	S
	40		R. E. Hanson, Capt, USAF	May 53	S
	41	The Silent Area Forecasting Problem	W. K. Widger, Jr.	May 53	S
	42	An Analysis of the Contrail Problem	R. A. Craig	Jun 53	C
	43	Sodium in the Upper Atmosphere	L. E. Miller	Jun 53	U
	44	Silver Iodide Diffusion Experiments Conducted at Camp Wellfleet, Mass., During July-August 1952	P. Goldberg A. J. Parziale G. Faucher B. Manning H. Lettau	Jun 53	U
	45	The Vertical Distribution of Water Vapor in the Stratosphere and the Upper Atmosphere	L. E. Miller	Sep 53	U
	46	Operation IVY Project 6.11. Free Air Atomic Blast l'ressure and Thermal Measurements — Final Report	N. A. Haskell J. O. Vann, Lt Col, USAF P. R. Gast	Sep 53	S-RD

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Num	Title				
47	Critical Envelope Study for the B61-A	Author	1	Date	Security Class.
40		N. A. Has R. M. Cho M. H. Seav	oman	P 53	S-RD
48 49	Operation Upshot-Knothole Project 1.3. Fre Air Atomic Blast Pressure Measurements. Revised Report	e N. A. Hask R. M. Brube Maj., USAF	ell No	v 53.	S-RD
50	Maximum Hamidity in Engineering Design	N. Sissenwi			
30	Probable Ice Island Locations in the Arctic				U
51	,,	A. P. Crury I. Browne	May	54	U
52	. Investigation of TRAC for Active Air Defense Purposes	G. W. Wares R. Penndorf V. G. Plank B. H. Grossm	Dec	53	S-RD
	Radio Noise Emissions During Thermonuclear Reactions	T. J. Keneshe		<b>34</b>	С
53	A Method of Correcting Tabulated Rawinsonde Wind Speeds for Curvature of the Earth	R. Leviton	Jun 5		ŭ
54	A Proposed Radar Storm Warning Service for Army Combat Operations	M. G. H. Ligde	Aug 54	<b>,</b>	ប
	A Comparison of Altitude Corrections for Blast Overpressure	N. A. Haskell	Sep 54	s	
	Attenuating Effects of Atmospheric Liquid Vater on Peak Overpressures from Blast Waves	H. P. Gauvin J. H. Healy M. A. Bennet	Oct 54	s	
D A	indspeed Profile, Windshear, and Gusts for esign of Guidance Systems for Vertical Rising ir Vehicles	N. Sissenwine	Nov 54	บ	
Pr	ne Suppression of Aircraft Exhaust Trails	C. E. Anderson	Nov 54	บ	
	eliminary Report on the Attenuation of Thermal diation From Atomic or Thermonuclear Weapons	R. M. Chapman M. H. Scavey	Nov 54	S-RD	
Met	ght Errors in a Rawin System  corological Aspects of Constant Level	R. Leviton	Dec 54	บ	
Ball	, 334	W. K. Widger, Jr. M. L. Haas E. A. Doty, L. Col. E. M. Darling, Jr.	Dec 54	s 	ange and a second

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Number	Title	Author	Date	Security Class.
62	Variations in Geometric Height of 30 to 60,000 Ft. Pressure Altitudes	N. Siesenwine A. E. Cole W. Baginsky	Dec 54	C-MA
63	Review of Time and Space Wind Fluctuations Applicable to Conventional Ballistic Determinations	W. Baginsky N. Siesenwine B. Davidson H. Lettau	Dec 54	U
64	Cloudiness Above 20,000 Feet for Certain Stellar Navigation	A. E. Cole	Jan 55	s
65	The Fensibility of the Identification of Hail and Severe Storms	D. Atlas R. Donaldson	Jan 55	ប
66	The Rate of Rainfall Frequencies Over Selected Air Houtes and Destinations	A. E. Cole N. Sissenwine	Mar 55	s
67	Some Considerations on the Modelling of Cratering Phenomena in Earth	N. A. Haskell	Apr 55	S-RD
68	The Preparation of Extended Forecasts of the Pressure Height Distribution in the Free Atmosphere Over North America by Use of Empirical Influence Functions	R. M. White	May 55	U
69	Cold Weather Effects on B-62 Launching Personnel	N. Sissenwine	Jun 55	s
70	Atmospheric Pressure Pulse Measurements: Operation Castle	E. Flauraud	Aug 55	S-RD
71	Refraction of Shock Waves in the Atmosphere	N. A. Haskell	Aug 55	s
72	Wind Variability as a Function of Time at Muroc, California	B. Singer	Sep 55	U
73	The Atmosphere	N. C. Gerson	Sep 55	ń
74	Areal Variation of Ceiling Height	W. Baginsky A. E. Cole	Oct 55	C-MA
75	An Objective System for Preparing Operational Weather Forecasts	I. A. Lund	Nov 55	U
76	The Practical Aspect of Tropical Meteorology	C. E. Palmer C. W. Wise L. J. Stempson G. H. Duncan	Sep. 55	U
77	Remote Determination of Soil Trafficability by Aerial Penetrometer	C. Molineux	Oct 55	U
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Number	Title	Author	Date	Security Class.
78	Effects of the Primary Cosmic Radiation on Matter	H. O. Curtis	Jan 56	U
79	Tropospheric Variations of Refractive Index at Microwave Frequencies	C. F. Campen A. E. Cole	Oct 55	U
80	A Program to Test Skill in Terminal Forecasting	I. I. Gringorten I. A. Lund M. A. Miller	Jun 55	U
81	Extrone Atniespheres and Ballistic Densities	N. Sissenwine A. E. Cole	Jul 55 .	U
82	Rotational Frequencies and Absorption Coefficients of Atmospheric Gases	S. N. Ghosh H. D. Edwards	Mar 56	U
83	lonospheric Effects on Positioning of Vehicles at High Altitudes	W. Pfister T.J. Keneshea	Mar 56	U
84	Pre-Trough Vinter Precipitation Forecasting	P. W. Funke	Feb 57	U
85	Geomagnetic Field Extrapolation Techniques - An Evaluation of the Poisson Integral for a Plane	J. F. McClay P. Fougere	Feb 57	s

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